

Terrestrial carbon and nutrient fluxes and the biogeochemical responses in the Mississippi River plume and Northern Gulf of Mexico

Wei-Jun Cai

The University of Georgia
Department of Marine Sciences
Athens, GA, USA



Ocean Carbon & Biogeochemistry

*Studying marine biogeochemical cycles
and associated ecosystems
in the face of environmental change*

Science Workshop, July 21-24, 2008, Woods Hole, MA, USA

Acknowledgement

Co-authors & contributors

- Xianghui Guo, Wei-Jen Huang, Yongchen Wang
- Steve Lohrenz---USM
- Michael Murrell---EPA
- Rik Wanninkhof and T.H. Peng---NOAA-AOML
- Minhan Dai---Xiamen Univ., China

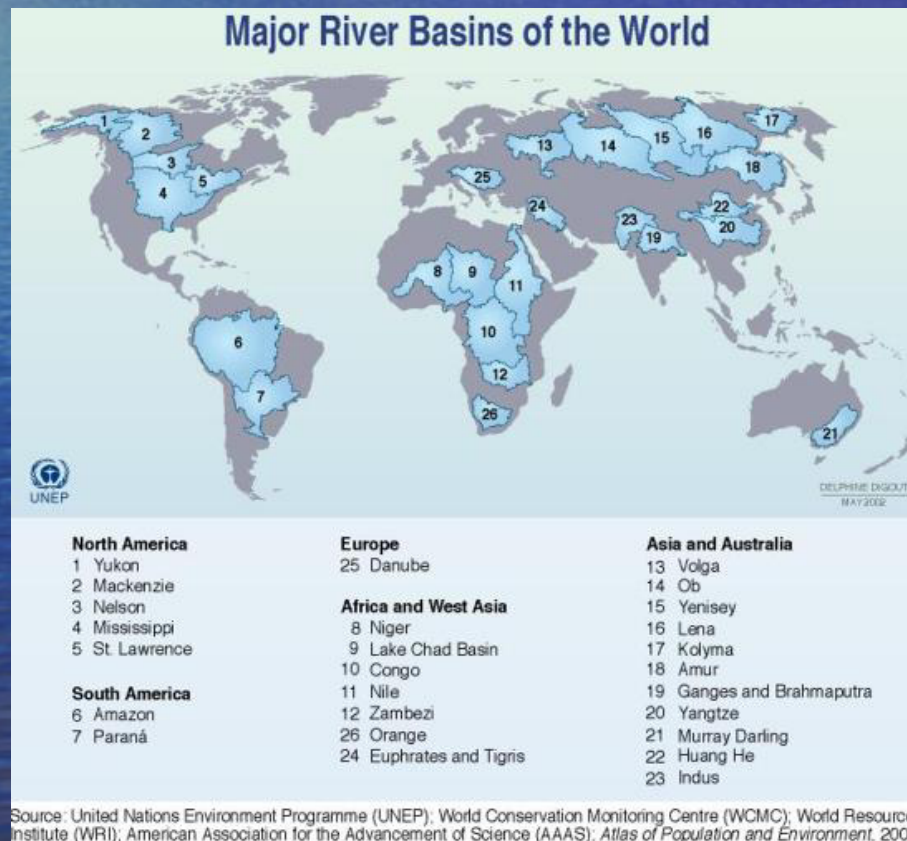
Funding

- NASA— (2005-2008)
- NOAA— (2005-2008)
- NSF—Chemical Oceanography (2008-2010)

outline

- River delivery of C and nutrients to coastal oceans
 - Global flux
 - Mississippi River flux
 - The effect of river loading on coastal metabolism
- Biogeochemical responses in the river plume and nearby areas
 - Nutrient behavior
 - Biological production
 - Control on the CO₂ system
- Carbon budget in the Northern Gulf of Mexico

Global river delivery of C and nutrients to coastal oceans



Global riverine C, N, and P flux

C, N and P inputs from land to the global coastal ocean (10^{12} mol/yr)					
Item	Meybeck 1982&1993	NEWS-model 2005	Smith et al. 2003	Adopted here	ratio
DIC	32		LOICZ	33.9	C:N 21
DIN	0.52	1.8	1.35	1.5	N:P 30
DIP	0.03	0.03	0.07	0.05	C:P 640
(Smith & Hollibaugh, 1993)					
DOC	20.5	14	17	18	C:N 26
DON	0.85	0.71		0.7	N:P 35
DOP	0.027	0.019		0.02	C:P 900
PIC	14			14	
POC	18.2		17	18	C:N 6
PON	2.9			2.9	N:P 5
PP(Inorg)	0.6			0.6	C:P 30
Total C	84.7			83.9	C:N 16
Total N	4.3			5.1	N:P 8
Total P	0.65			0.67	C:P 121
autotrophic loading based on DIN input and Redfield ratio					10
assuming 50% of DIN will be denitrified in coastal zone					5
assuming 80% of DIN will be denitrified					2
heterotrophic loading					
assuming xx% of land OC is respired in coastal ocean				0.70	25
N recycled					3
net heterotrophic loading					23

- DIC flux \sim TOC flux

- Heterotrophic loading \gg auto...

Table 2.1 Carbon Fluxes and Reservoirs Relevant to RiOMar Systems

	C flux (Tg/yr)	Range (Tg C yr⁻¹)	Reference
<i>Total Riverine Carbon Input to the Ocean</i>		700-1000	
Dissolved Inorganic Carbon (DIC)	400	381-410	1,2,3
Total Organic Carbon (TOC)	450	200-530	2,3,4,5
Particulate Organic Carbon (POC)	200	138-288	4,6,7
Dissolved Organic Carbon (DOC)	250	214-360	2,4,8,9,10
<i>Terrestrial Storage</i>		600-1500	11
<i>Burial in Marine Sediments</i>	130	98-138	5,12,13
Terrestrial		43-104	5,12
Marine		55	5,14

¹Meybeck, 1993

²Meybeck and Vörösmarty, 1999

³Degens et al., 1991

⁴Spitzky and Ittekkot, 1991

⁵Schlunz and Schneider, 2000

⁶Lyons et al., 2002 (and references within)

⁷Ittekkot and Laane, 1991

⁸Spitzky and Leenheer, 1991

⁹Aitkenhead and McDowell, 2000

¹⁰Hedges et al., 1997

¹¹Stallard, 1998

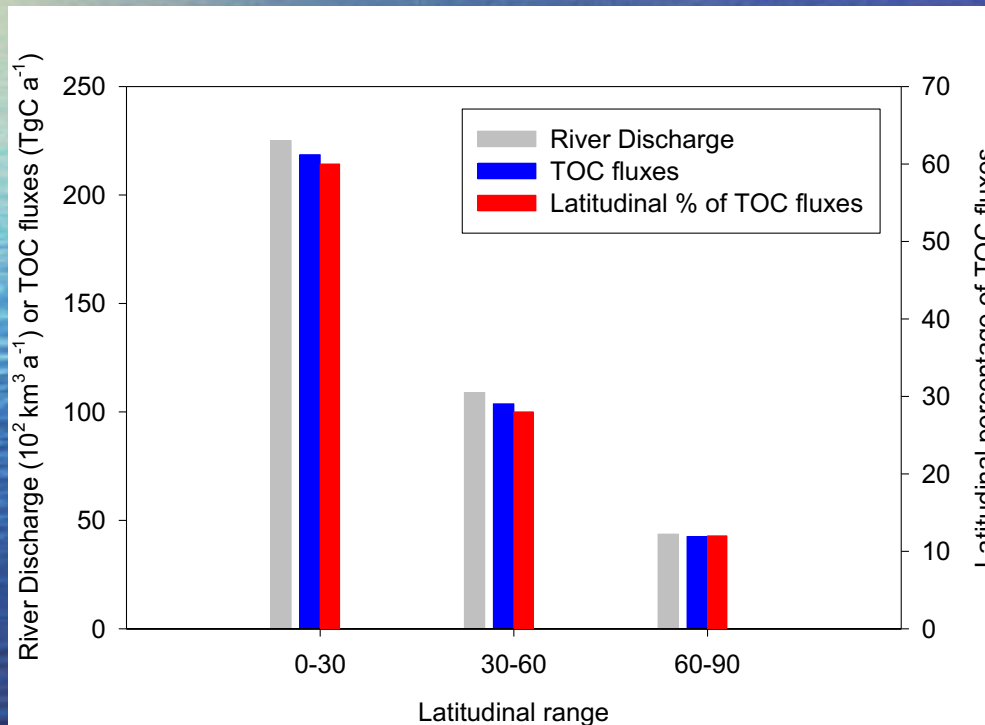
¹²Berner, 1982

¹³Hedges and Keil, 1995

¹⁴Berger et al., 1989

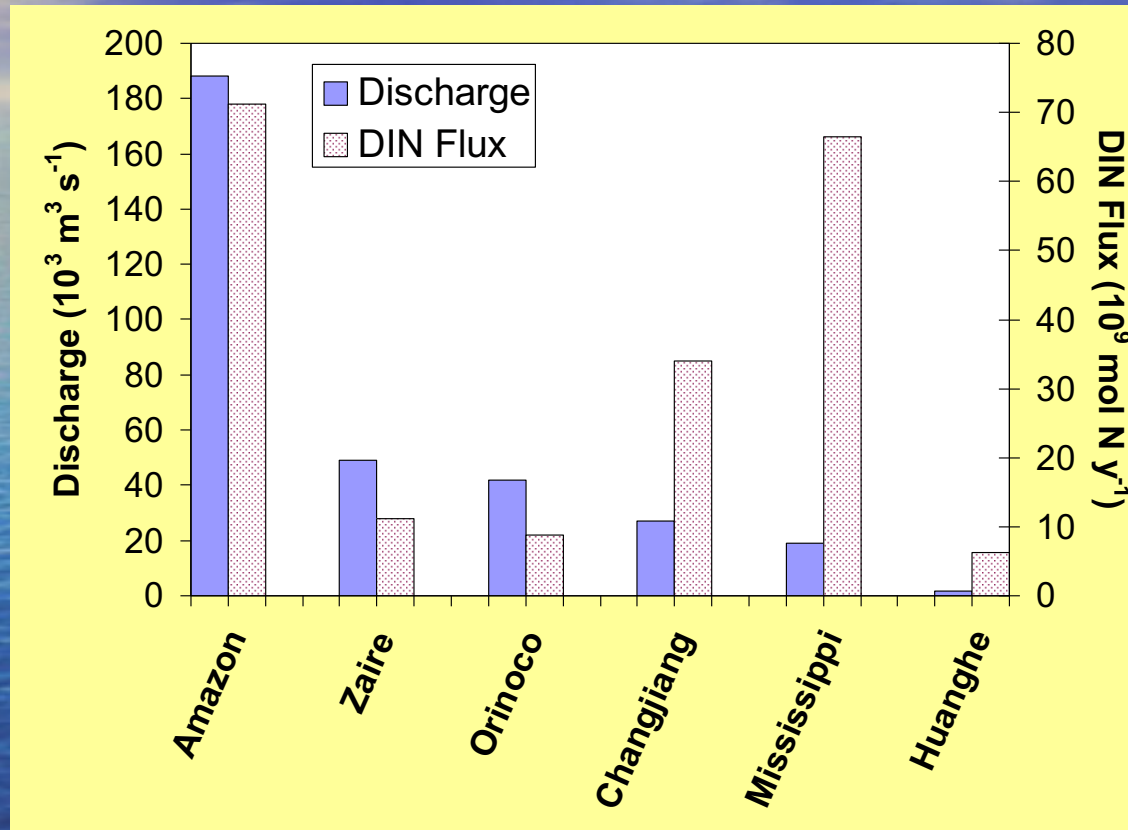
Contrast in organic C input

- deliver more heterotrophic loading to low-lat margins



Compiled by M. Dai and W-J. Cai;
also see Borges papers

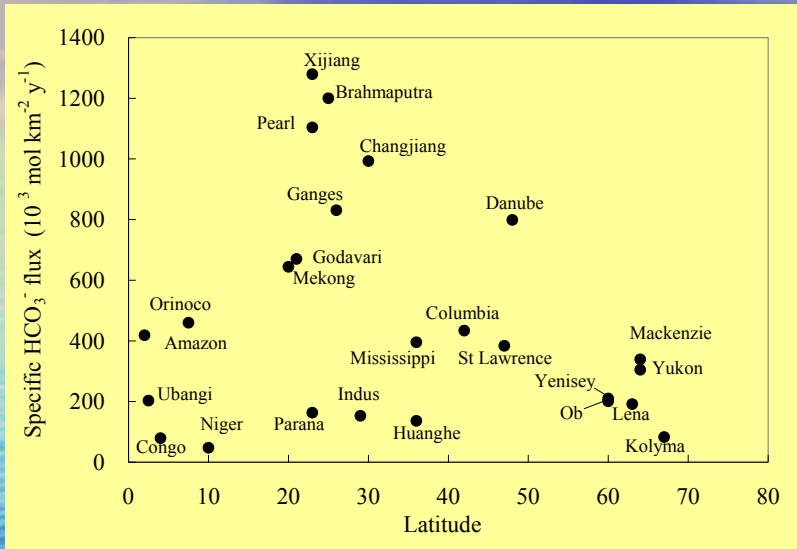
DIN flux in world's major rivers



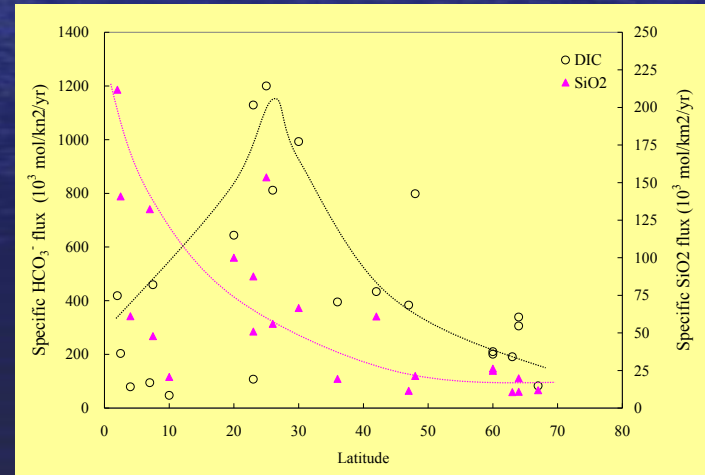
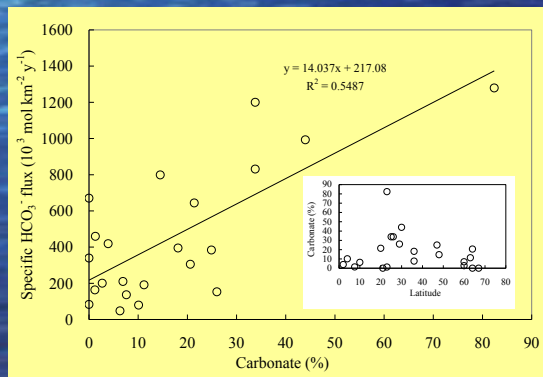
Source: Dagg et al. 2004

Speculate: River loading drives tropical margins towards heterotrophic, and mid lat margins to autotrophic.

Distribution of riverine HCO_3^- & Si flux



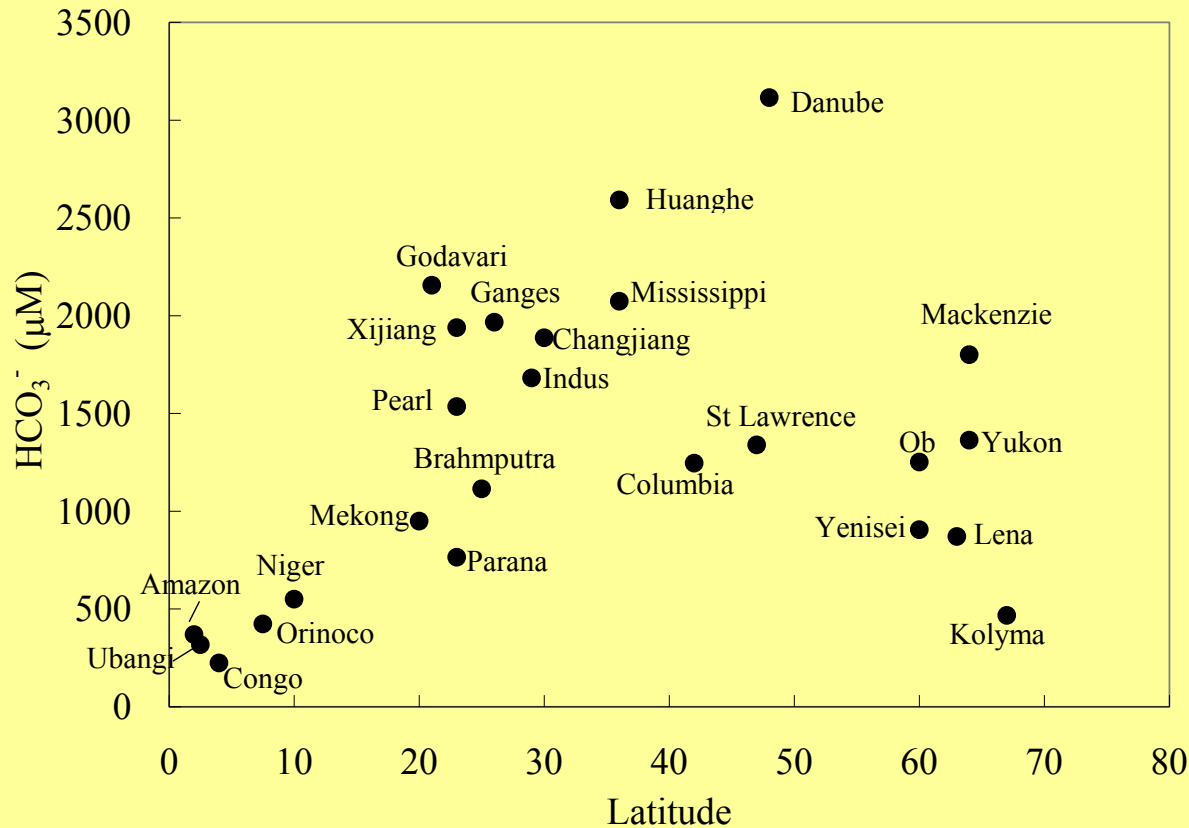
1. Latitudinal distribution in HCO_3^- flux
2. CaCO_3 minerals in drainage basin
3. Contrast to silicate flux



From Cai et al. 2008
Continental Shelf Research 28:1538-1549.

Cai unpublished

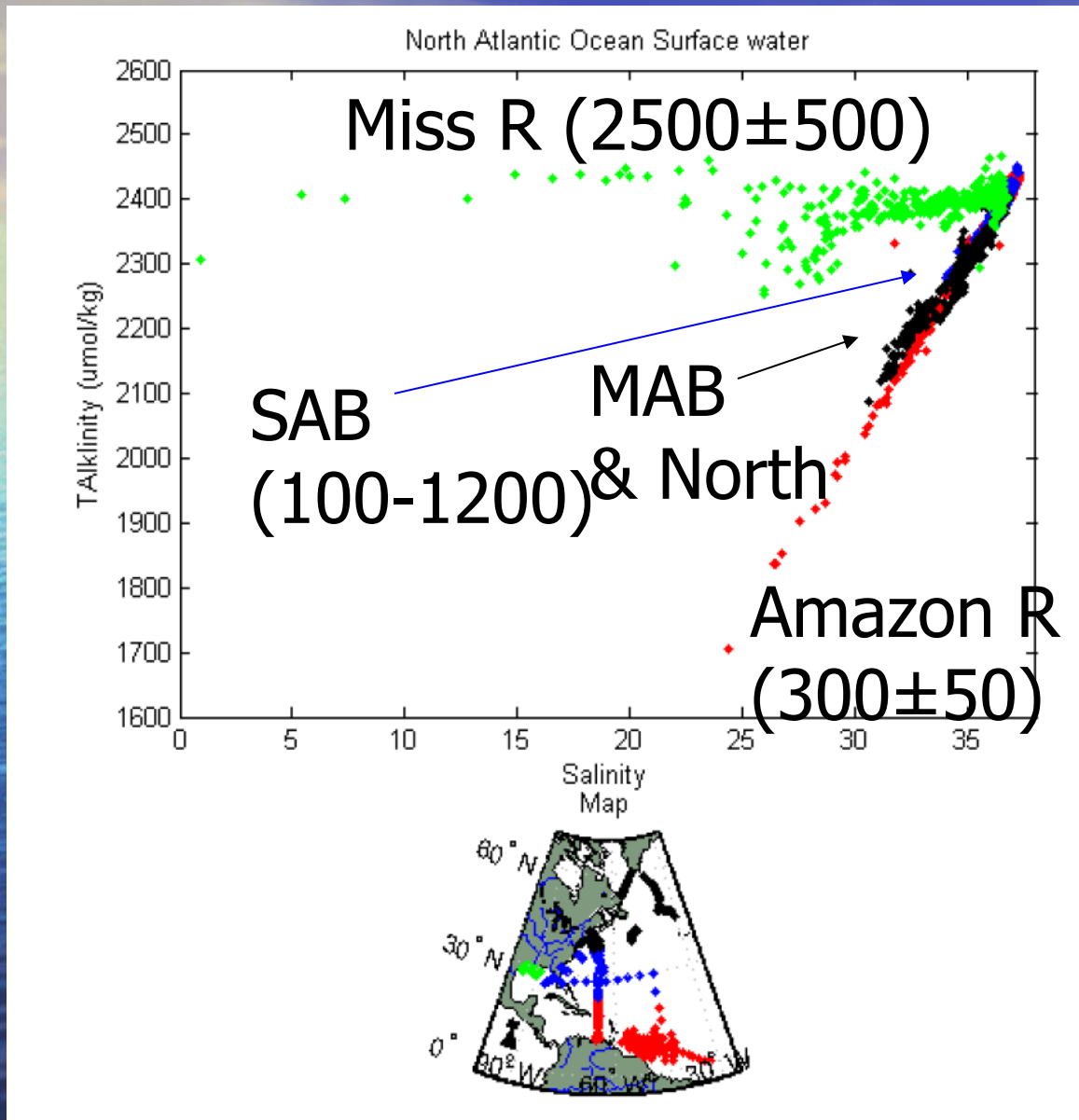
Distribution of riverine HCO_3^- concentration



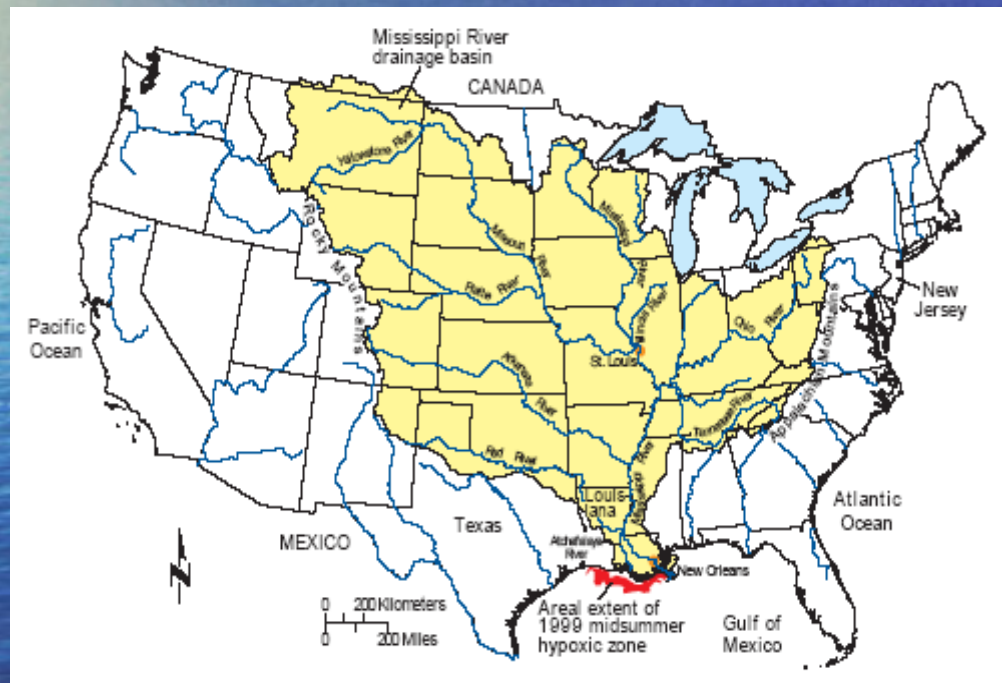
Maximum shifted to higher latitude as a result of precipitation pattern

From Cai et al. 2008
Continental Shelf Research 28:1538-1549.

Influence on TAlk distr in Ocean Margins



Mississippi River flux



- Largest river basin in North America and third largest in the world
- Drainage basin encompasses 41% of the lower 48 United States

Table 7.8.1. Average concentrations and annual fluxes of organic and inorganic C, N, and P delivered to the Gulf of Mexico by the Mississippi-Atchafalaya River System (MARS) (in 10^{12} g y^{-1} or Tg y^{-1}). Total water discharge is 530×10^9 m³ y^{-1} . TSM stands for total suspended matter.

Chemical	Concentration (mM)	Annual Flux Tg y^{-1}	Reference
TSM	---	210 ^s	Meade and Parker (1985)
POC	1.6% of TSM	3.4	Trefry et al. (1994)
		1.2 ^{&}	Duan and Bianchi 2006
DOC	0.28	1.8	Trefry et al. (1994)
	0.33	2.1	Benner and Opsahl 2001
	0.49	3.1	Bianchi et al. 2004
PIC	0.15% of TSM	0.31	Trefry et al. (1994)
DIC	0.219	21 [*]	Cai (2003)
TAlk	0.216	21 [*]	Cai (2003); Raymond and Cole (2003)
Total Nitrogen (N)		1.57	Goolsby et al. (1999)
NO ₃ + NO ₂		0.95	Goolsby et al. (1999); Howarth et al. (1996)
Ammonium		0.03	Goolsby et al. (1999)
Dissolved Org. N		0.38	Goolsby et al. (1999)
Particulate Org. N		0.20	Goolsby et al. (1999)
Particulate Org. N		0.45 ⁺	Trefry et al. (1994)
Total Phosphorus (P)		0.136	Goolsby et al. (1999)
PO ₄		0.042	Goolsby et al. (1999)
Particulate P		0.095	Goolsby et al. (1999)
Si-dissolved		2.32	Goolsby et al. (1999)

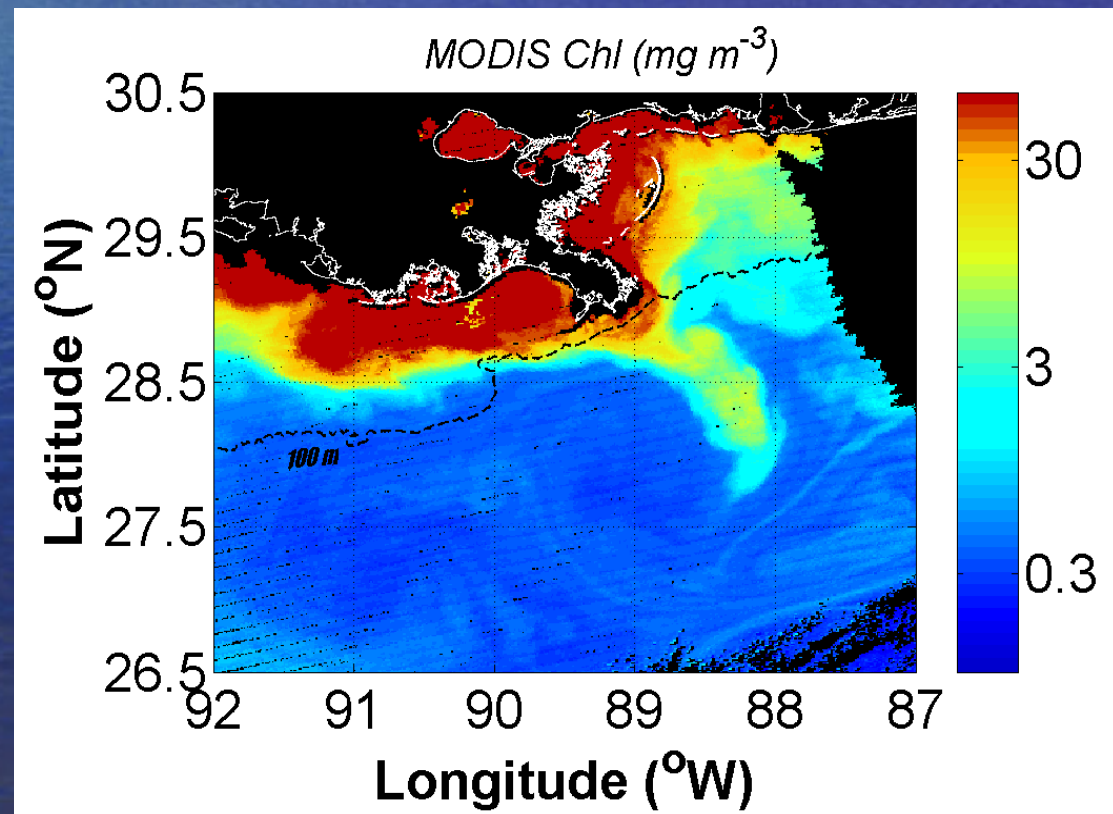
1. DIC >> TOC

2. Heterotrophic loading ~ autotrophic loading

3. Disparate time scales

Biogeochemical responses in the river plume and nearby area — biological production

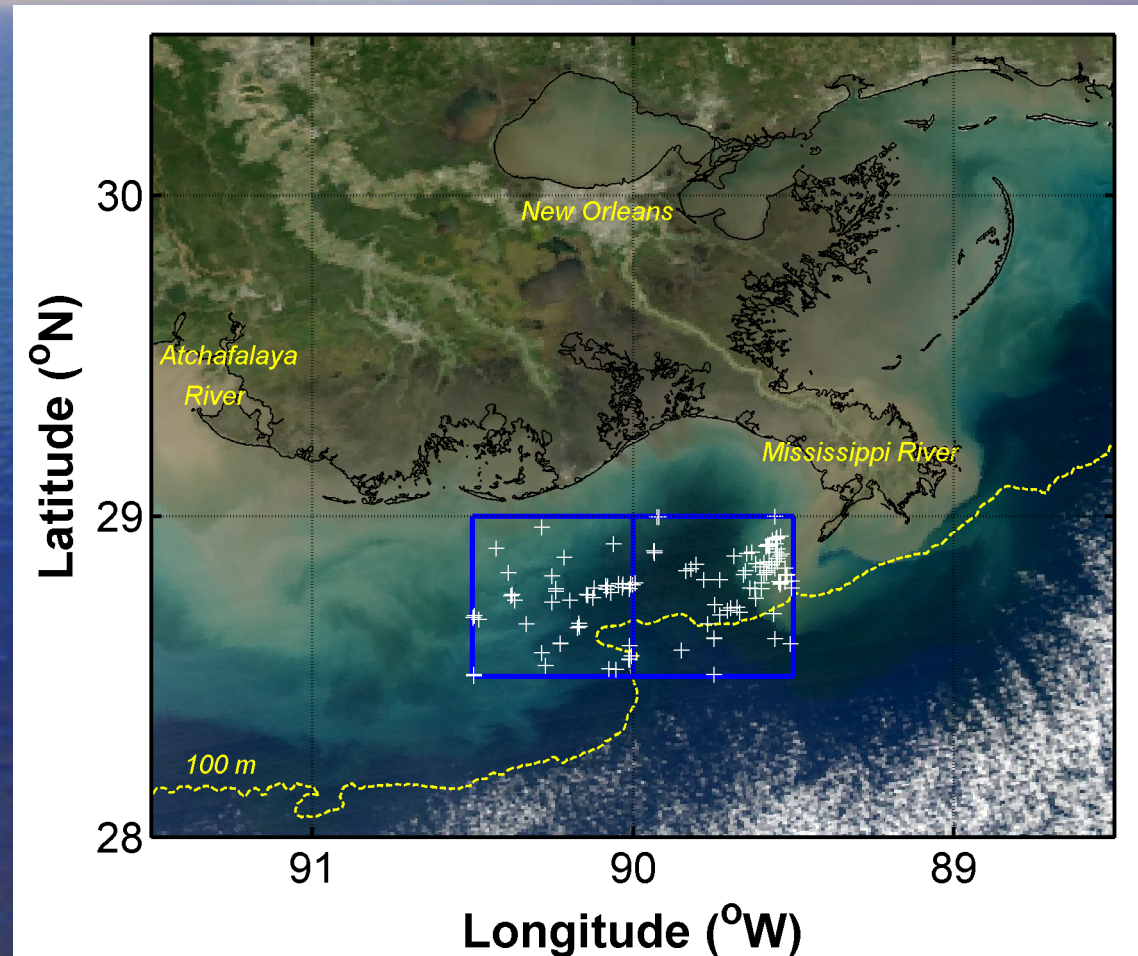
- Satellite evidence points towards linkages between high chlorophyll and river outflow



Relationship Between River Inputs and Coastal Ecosystem Properties

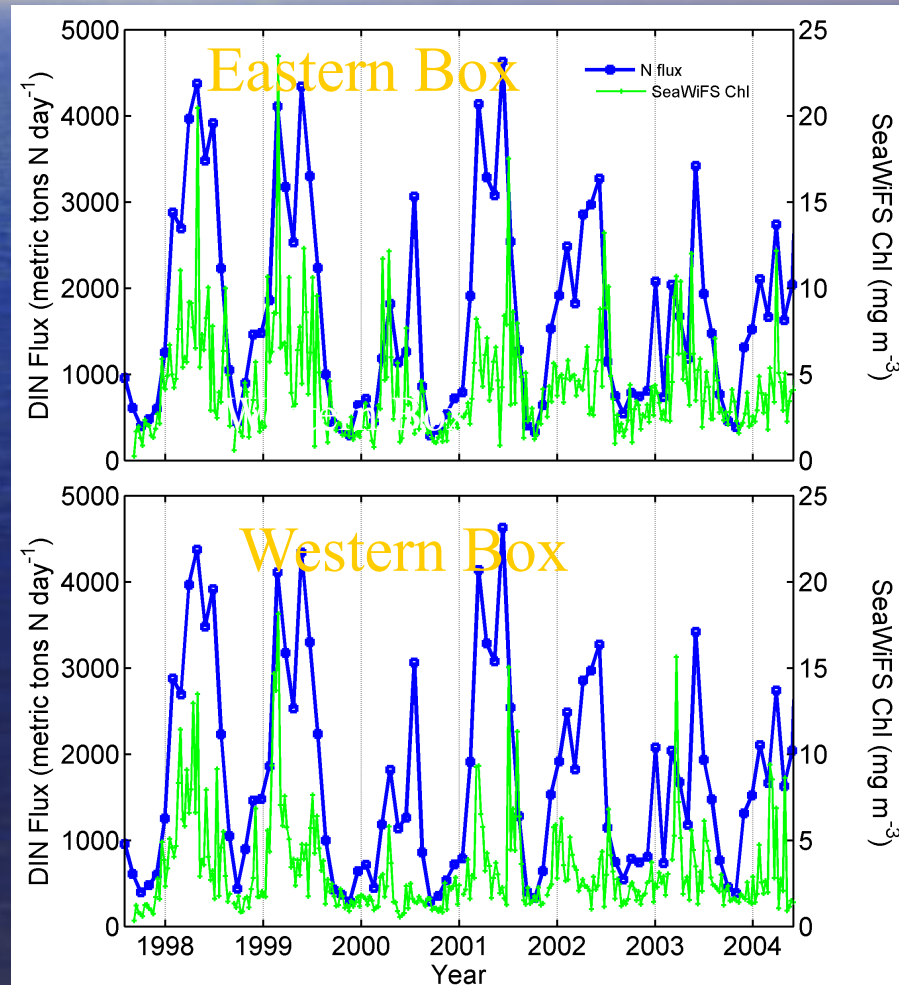
3 May 2004

- Satellite evidence points towards linkages between high chlorophyll and river outflow



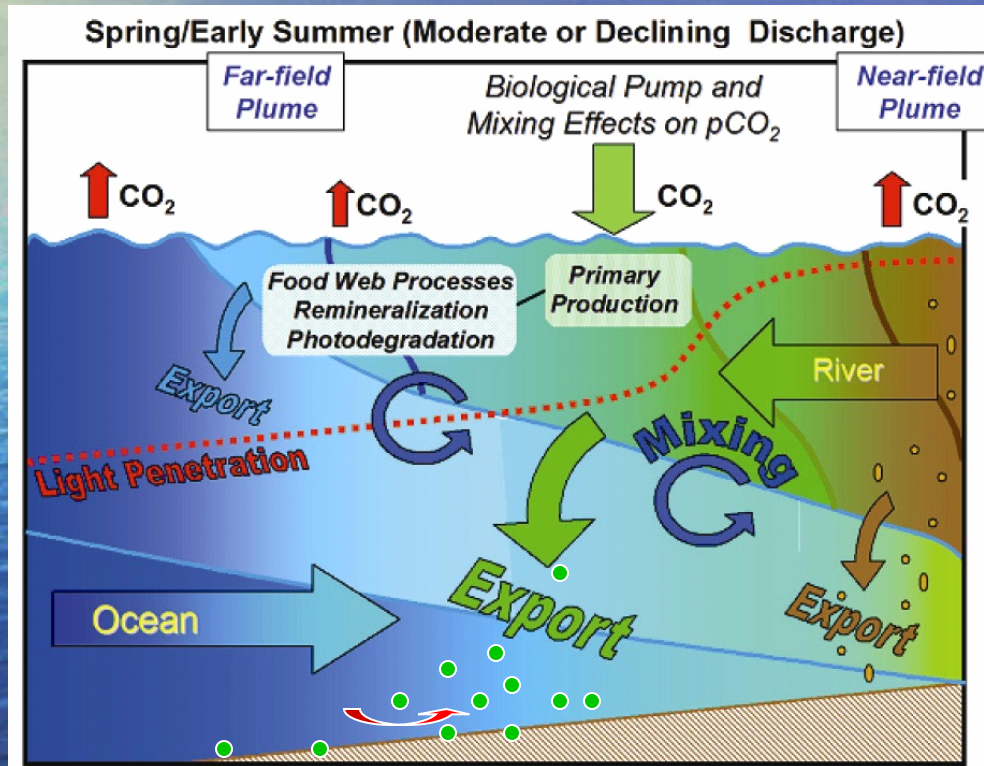
Relationship Between River Inputs and Coastal Ecosystem Properties

- Relationship between river DIN flux and satellite-derived chlorophyll



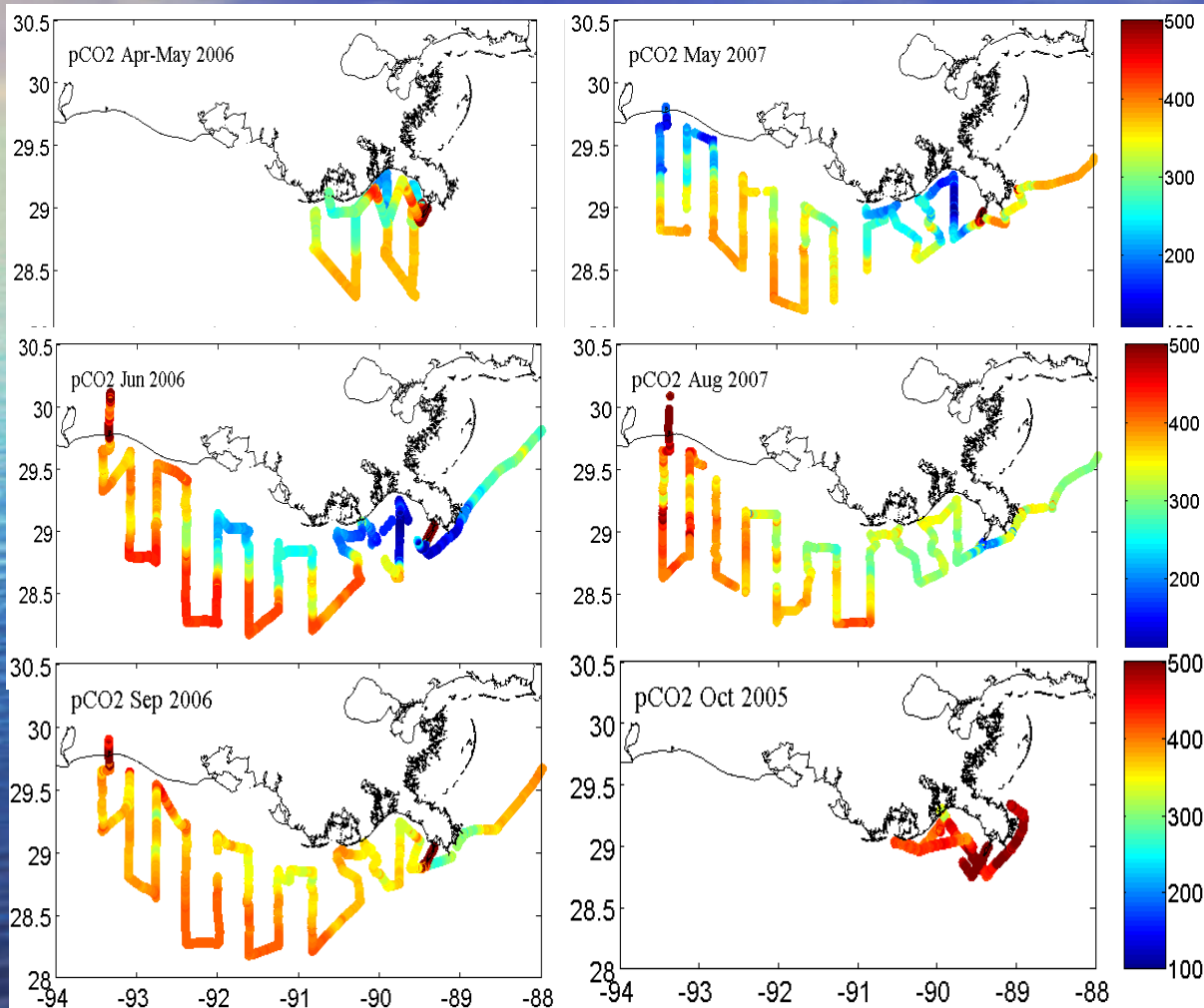
Source: Lohrenz et al. (2008)

Biogeochemical responses — nutrients and CO₂ system



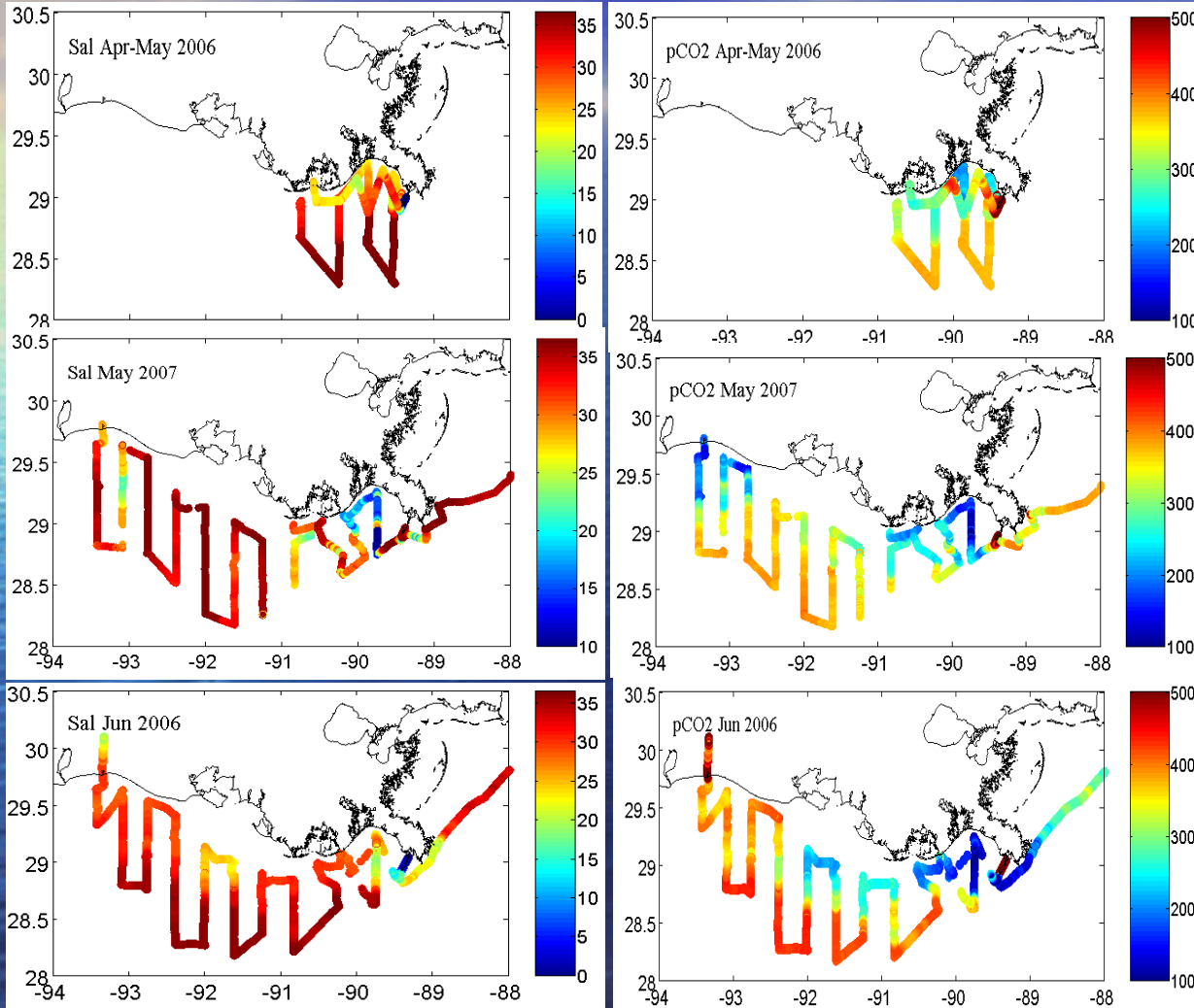
The high nutrient loading and the short turnover time contribute to a strong biological pump for carbon uptake in the mid-field

$p\text{CO}_2$ distribution



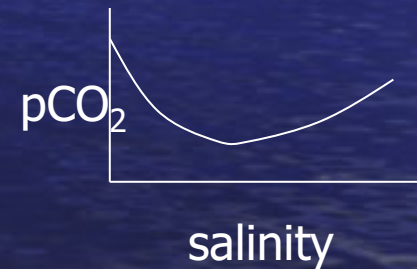
Salinity

$p\text{CO}_2$ (spring & summer)



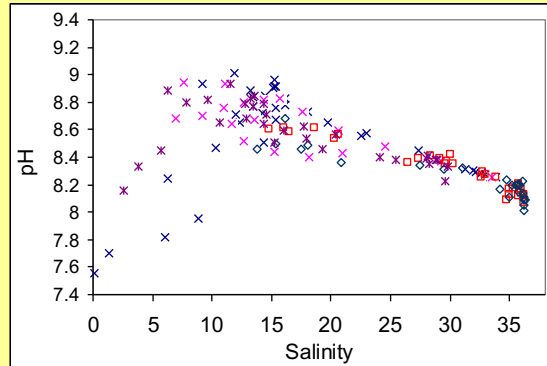
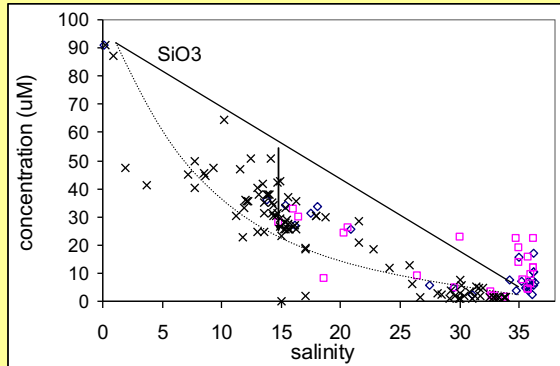
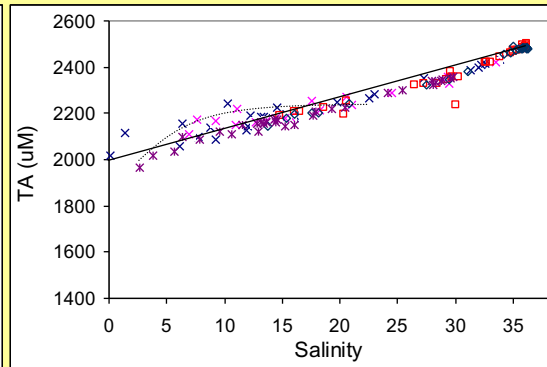
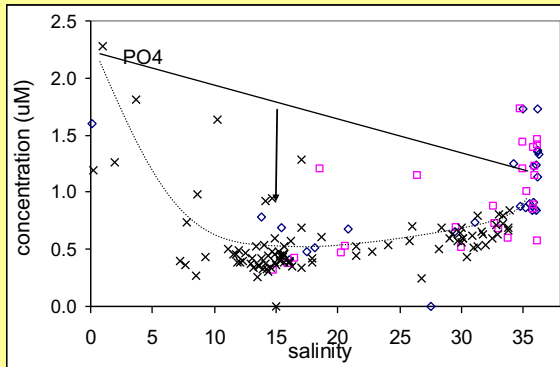
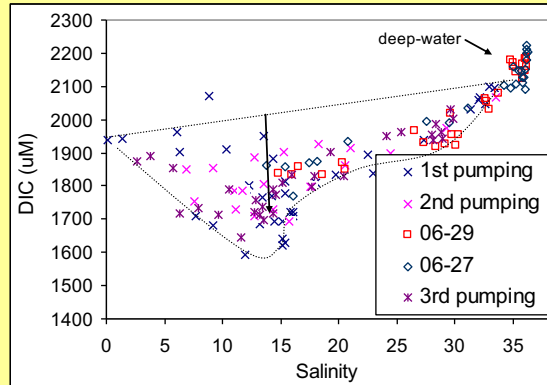
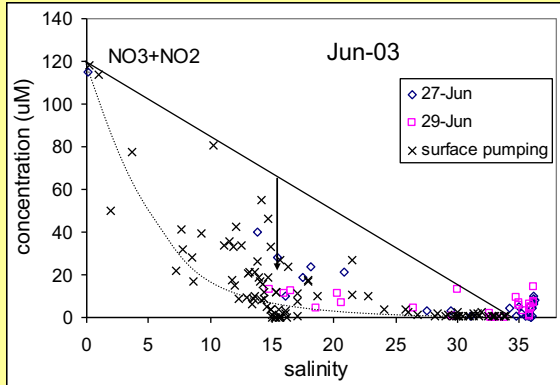
Salinity explains some but not all features

1. Inside river channels & embayment, low S high $p\text{CO}_2$ (turbidity)
2. In the plume, low S low $p\text{CO}_2$ (bio uptake)



3. Variability

Mississippi River plume, June 2003

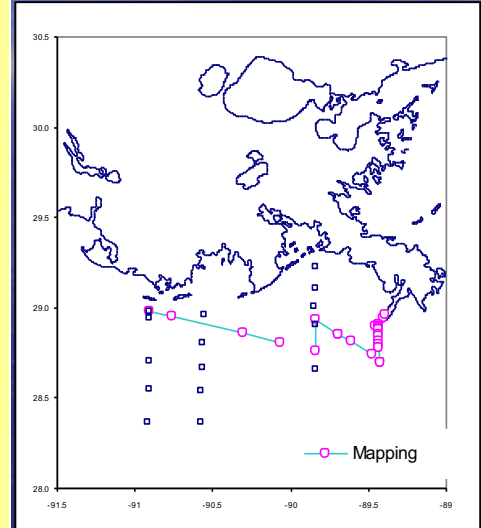
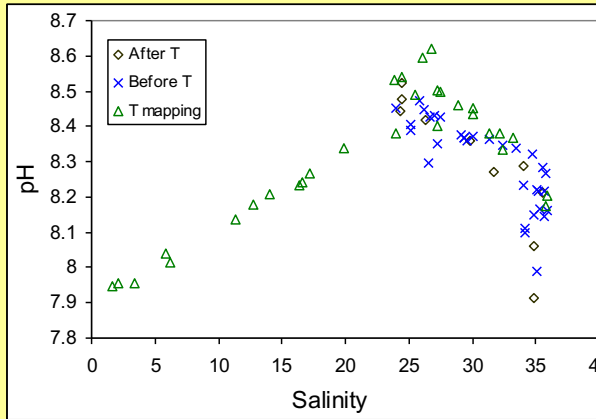
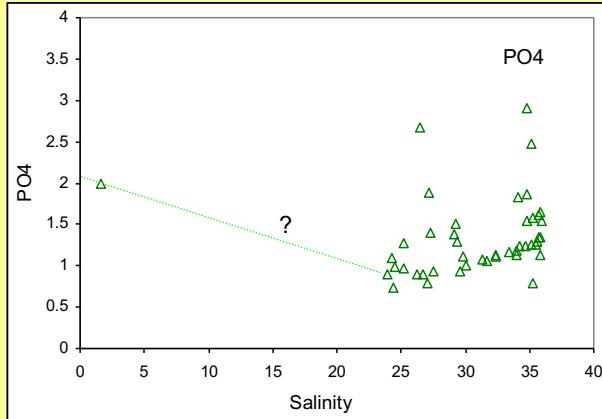
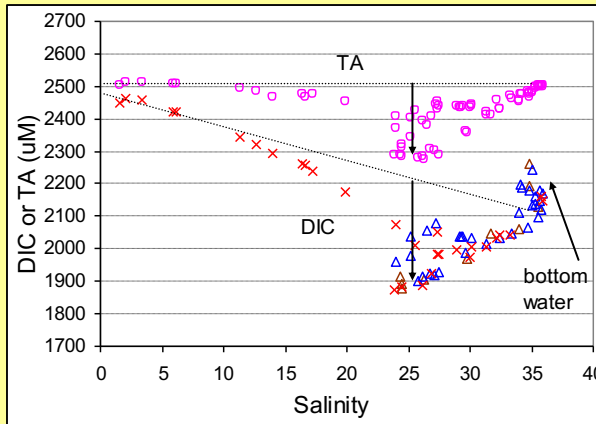
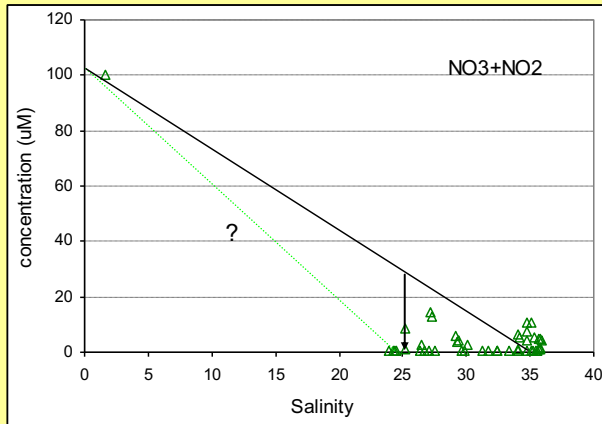


(Rodney Powell)

(W.-J Cai)

1. Great DIC removal & nutrient removal at S=15.
2. No TA removal.
3. At S=15, max Δ DIC \sim 430 μ M. Applying a Redfield ratio of 6.6, we would predict a max NO₃ removal of 65 μ M.

Mississippi River plume, August 2004

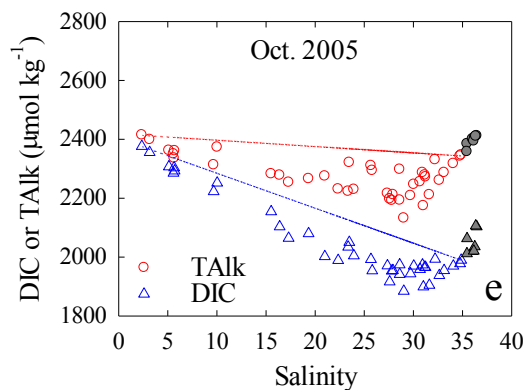
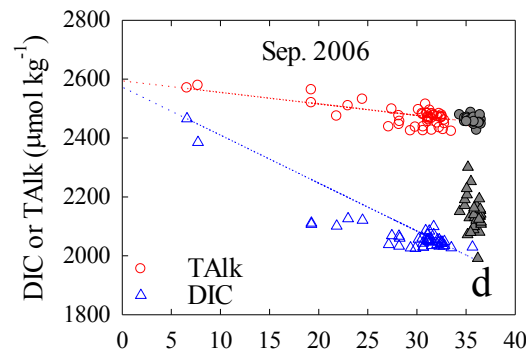
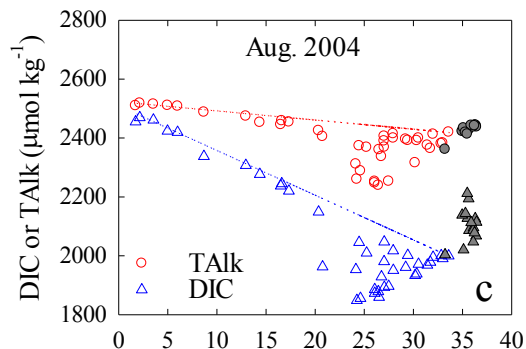
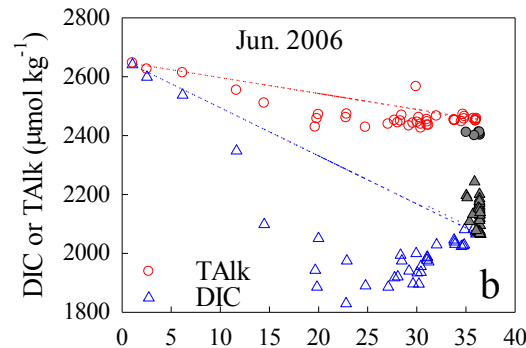
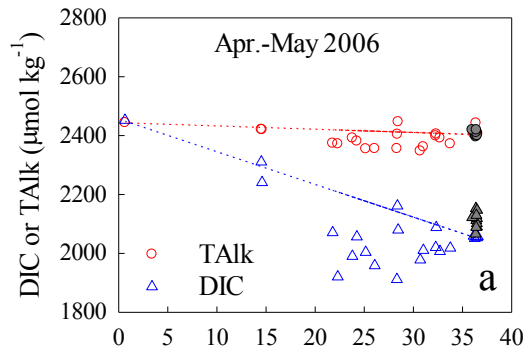


$$\Delta \text{DIC} = 320 \text{ uM}, \Delta \text{TA} = 210 \text{ uM}$$

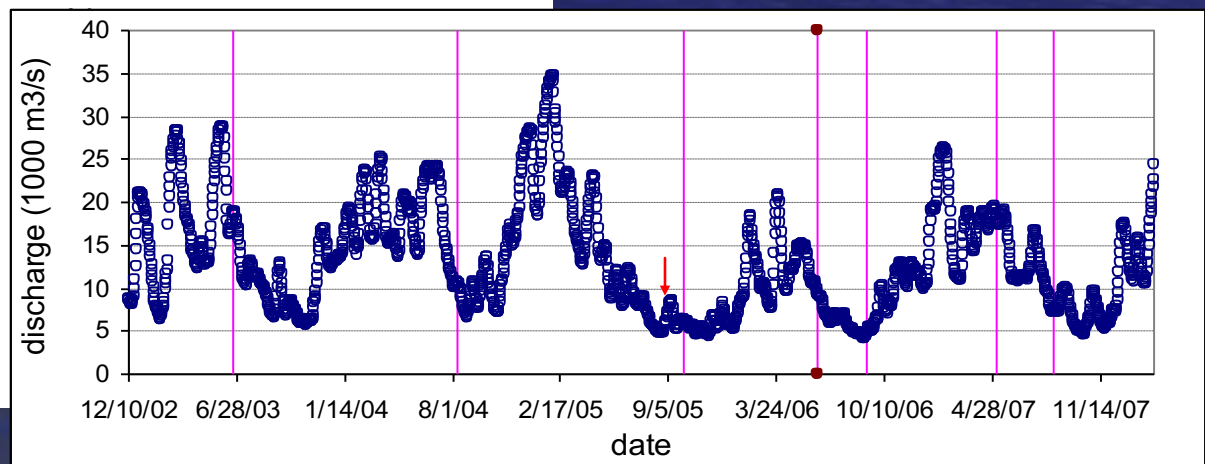
$$\text{DIC removal due to OC} = 320 - 210/2 = 215 \text{ uM}$$

$$\text{Predicted NO}_3 \text{ removal} = 33 \text{ uM}$$

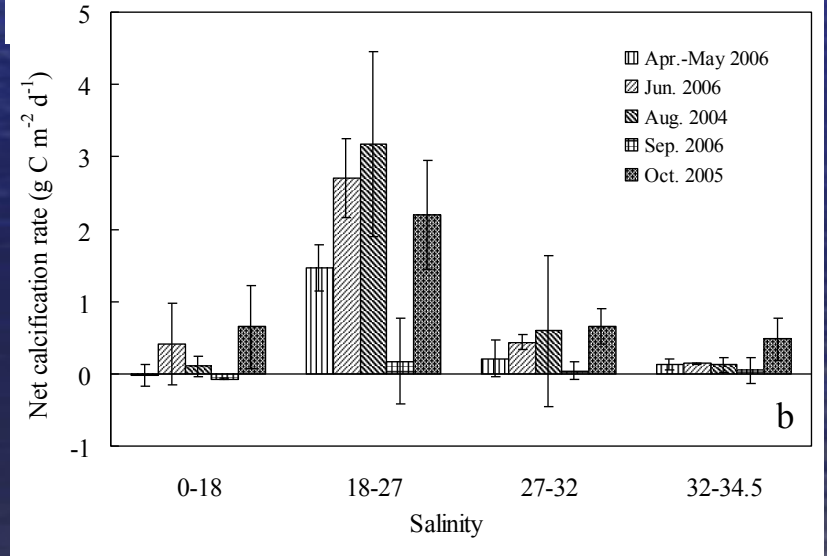
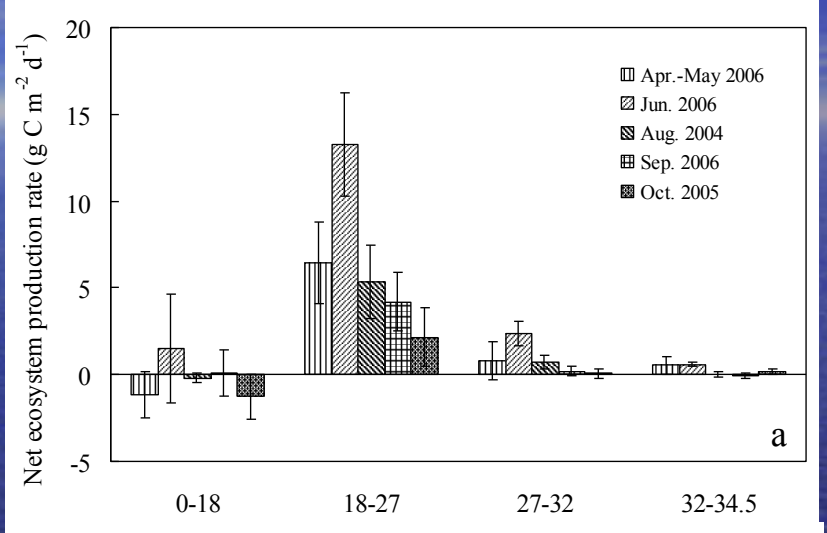
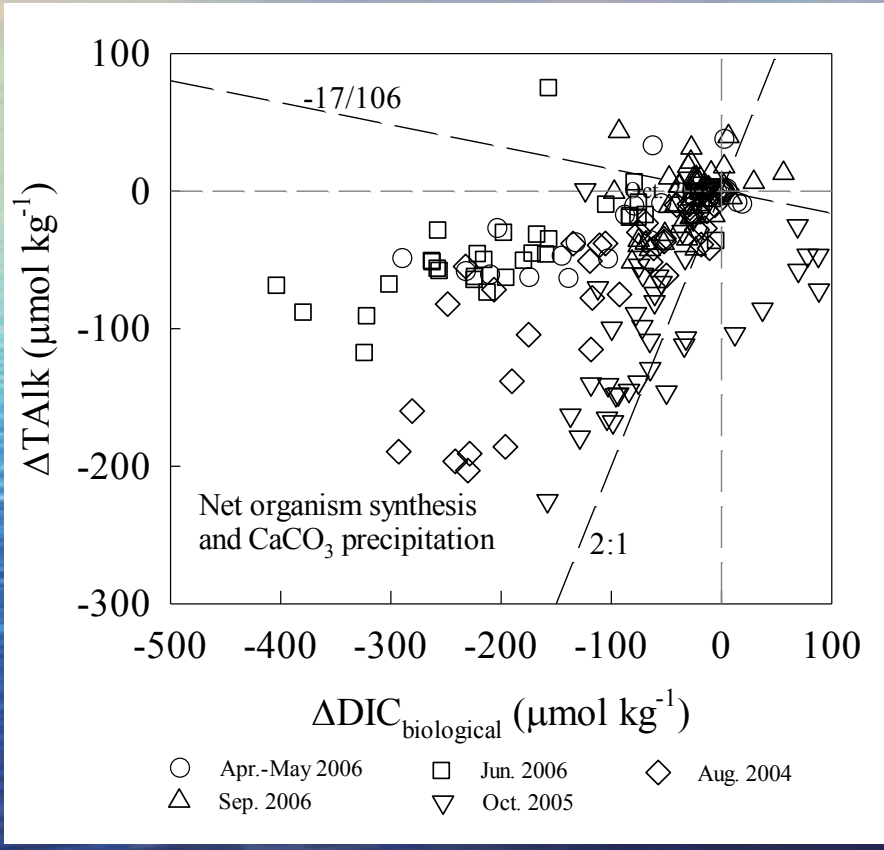
DIC and TAlk in the Mississippi River plume



DIC and TA removal

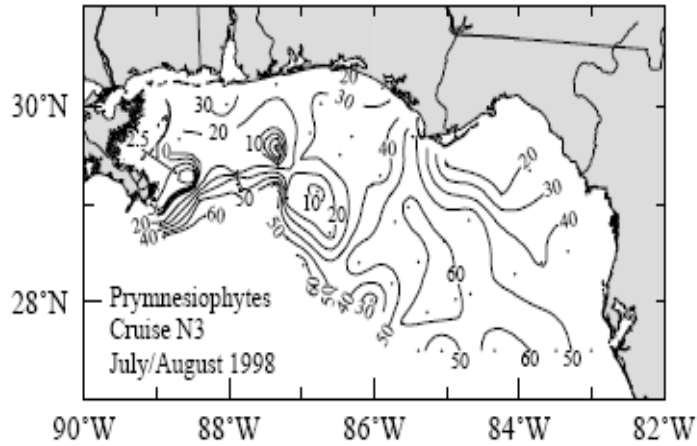


Net community production (NCP) & net calcification

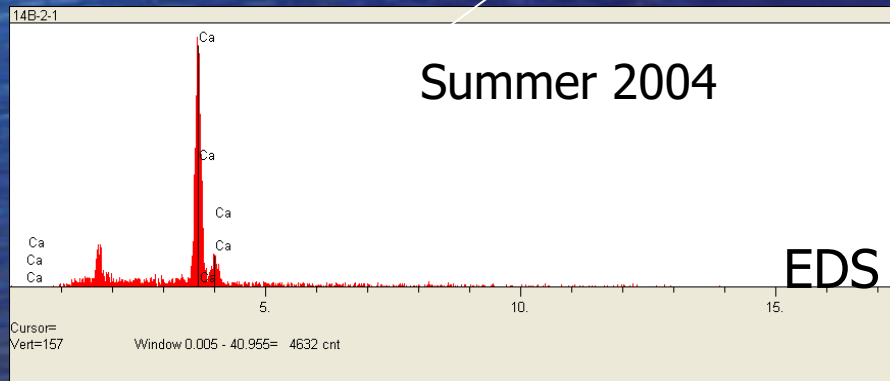
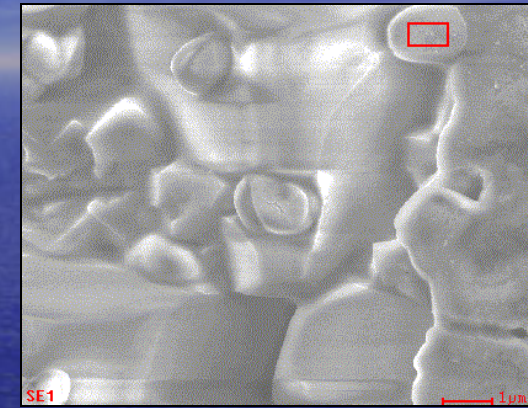
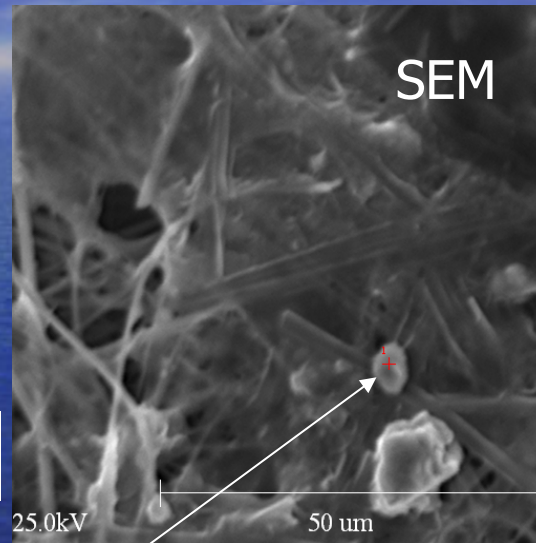


Mixing layer depth and water residence time taken from Green et al. (2006)

evidence of coccolithophores

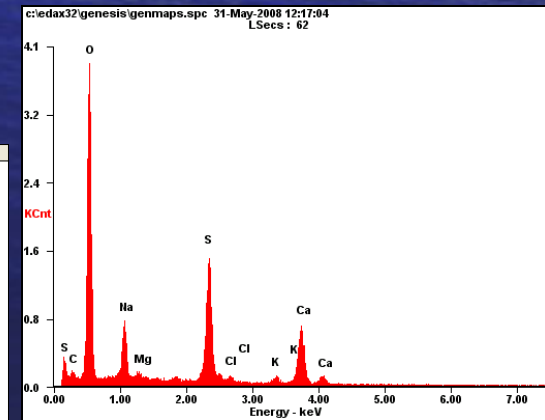


Y. Qian et al. / *Continental Shelf Research* 23 (2003) 1-17



Summer 2004

EDS



Summer 2008
W-J Cai (unpub)

Robert Stavn, Naval Research Laboratory

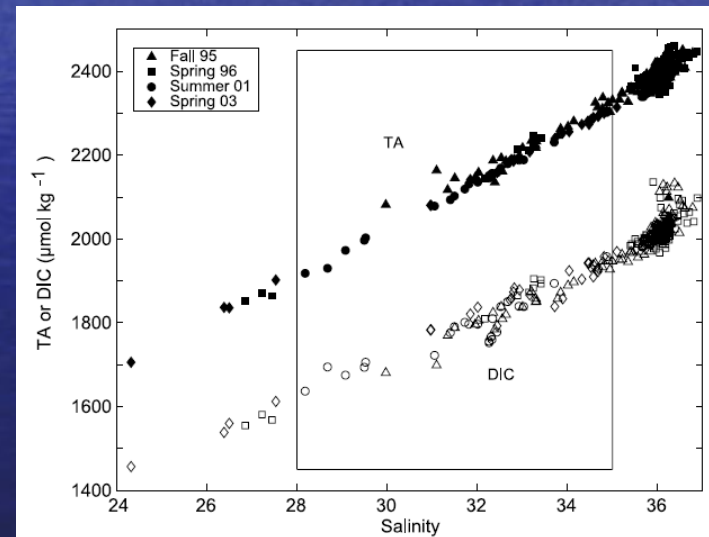
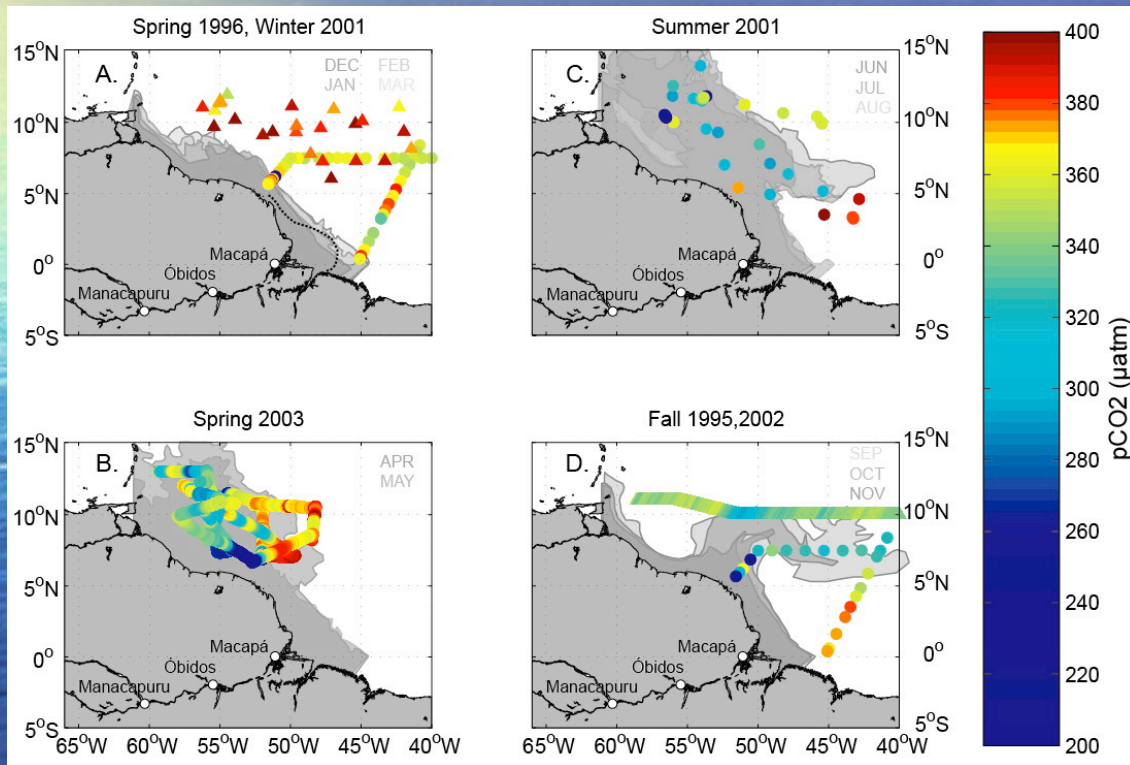
Significance of potential coccolith bloom in the Miss R plume

- It increases $p\text{CO}_2$ (opposite to diatom)

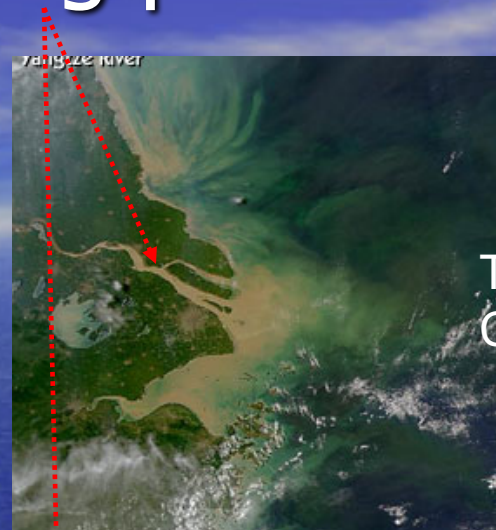
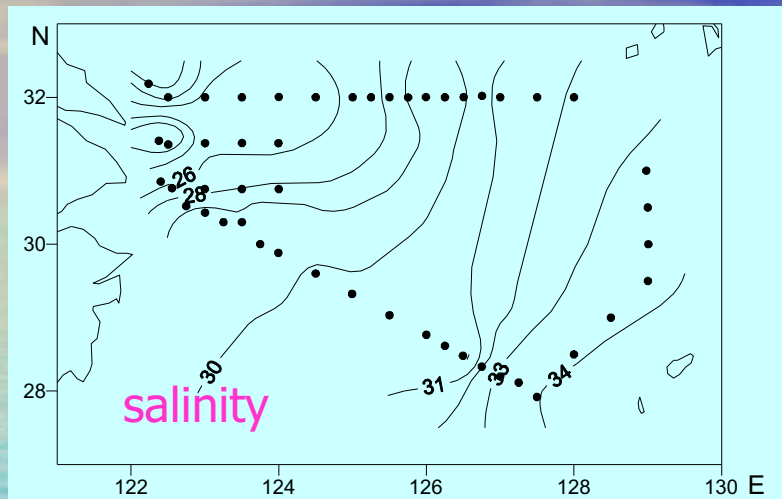


- Diatom bloom creates a low $p\text{CO}_2$, high pH & high CO_3^{2-} condition for coccolith bloom (ecosystem succession)
- No such response reported in the Amazon plume. Why? (ecosystem shift under anthropogenic pressure)

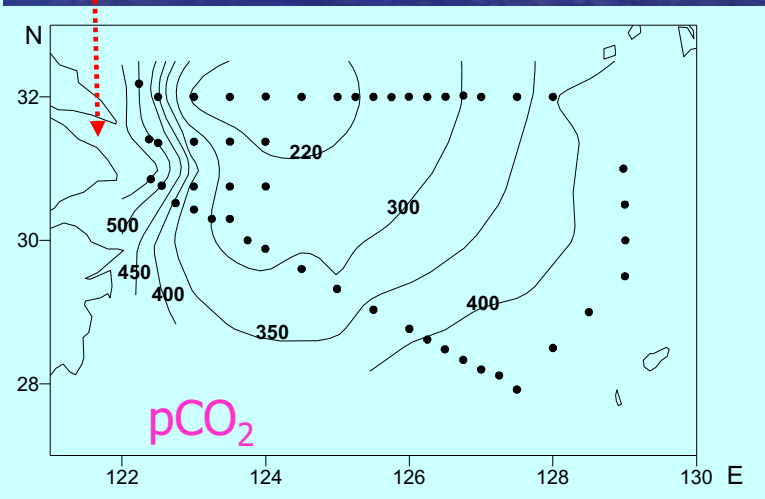
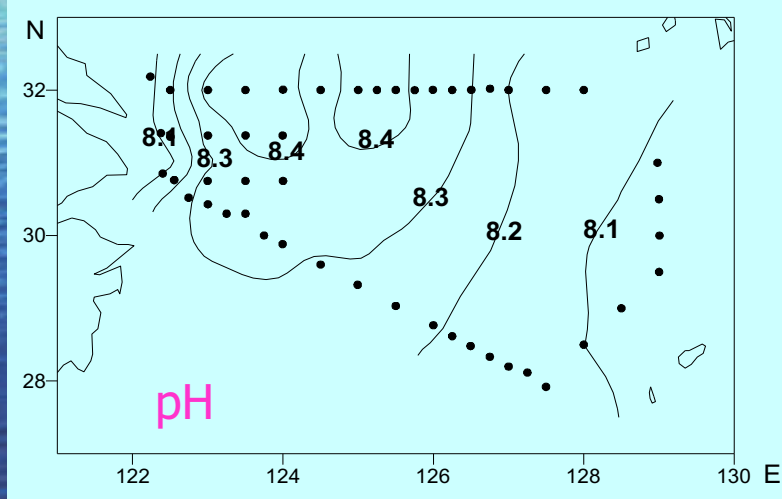
$p\text{CO}_2$ in the Amazon plume



$p\text{CO}_2$ in the Changjiang plume

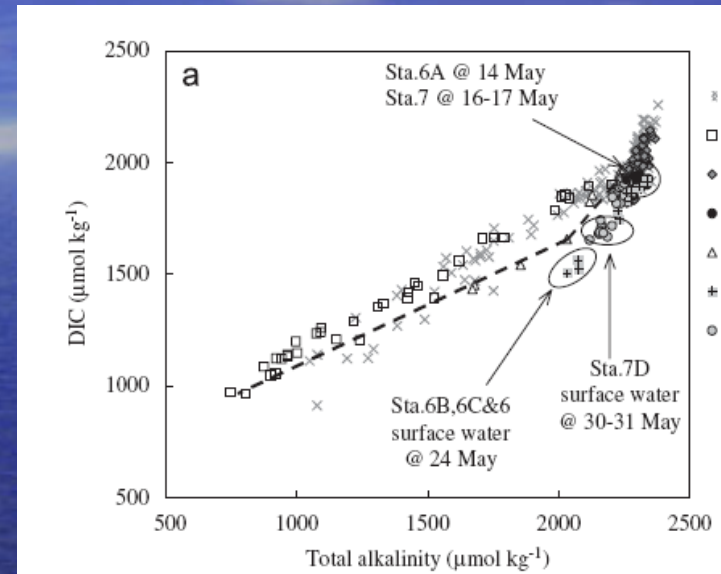
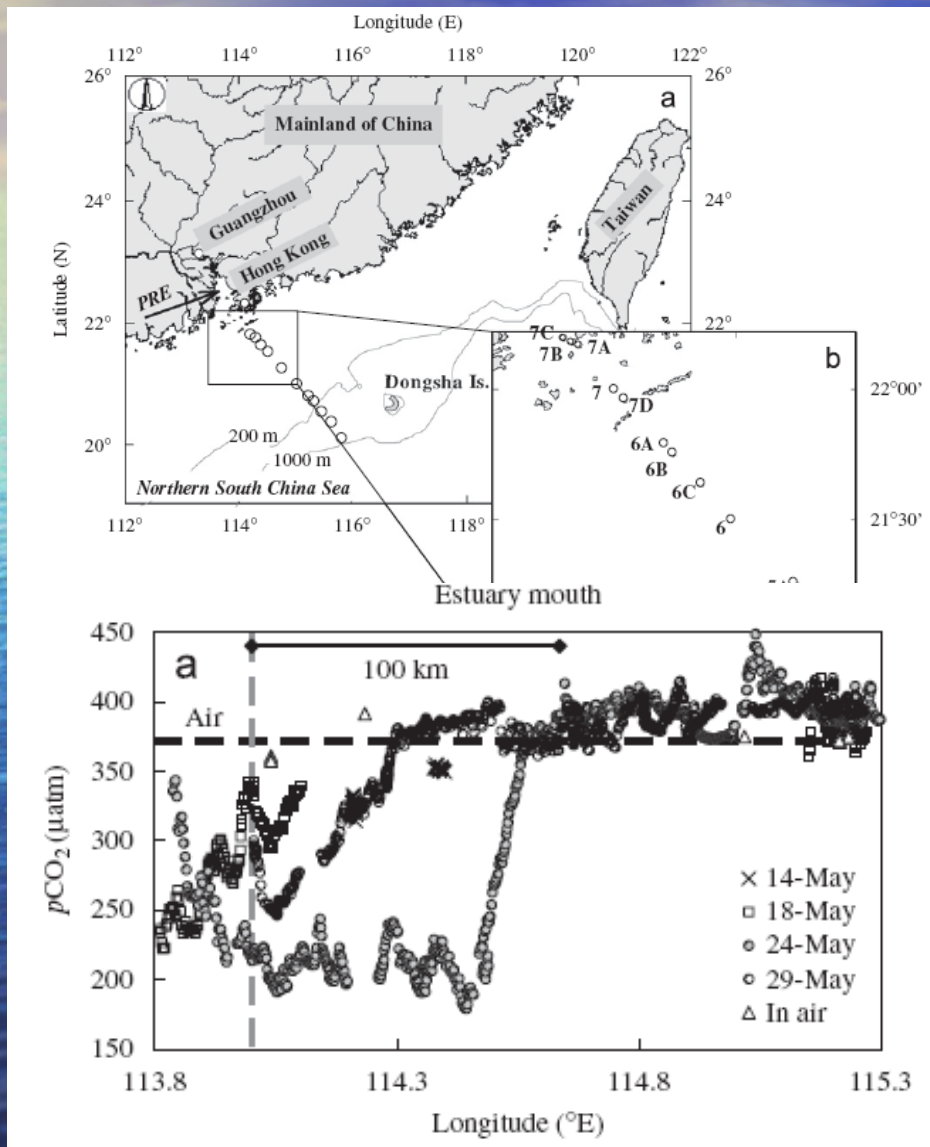


The East
China Sea



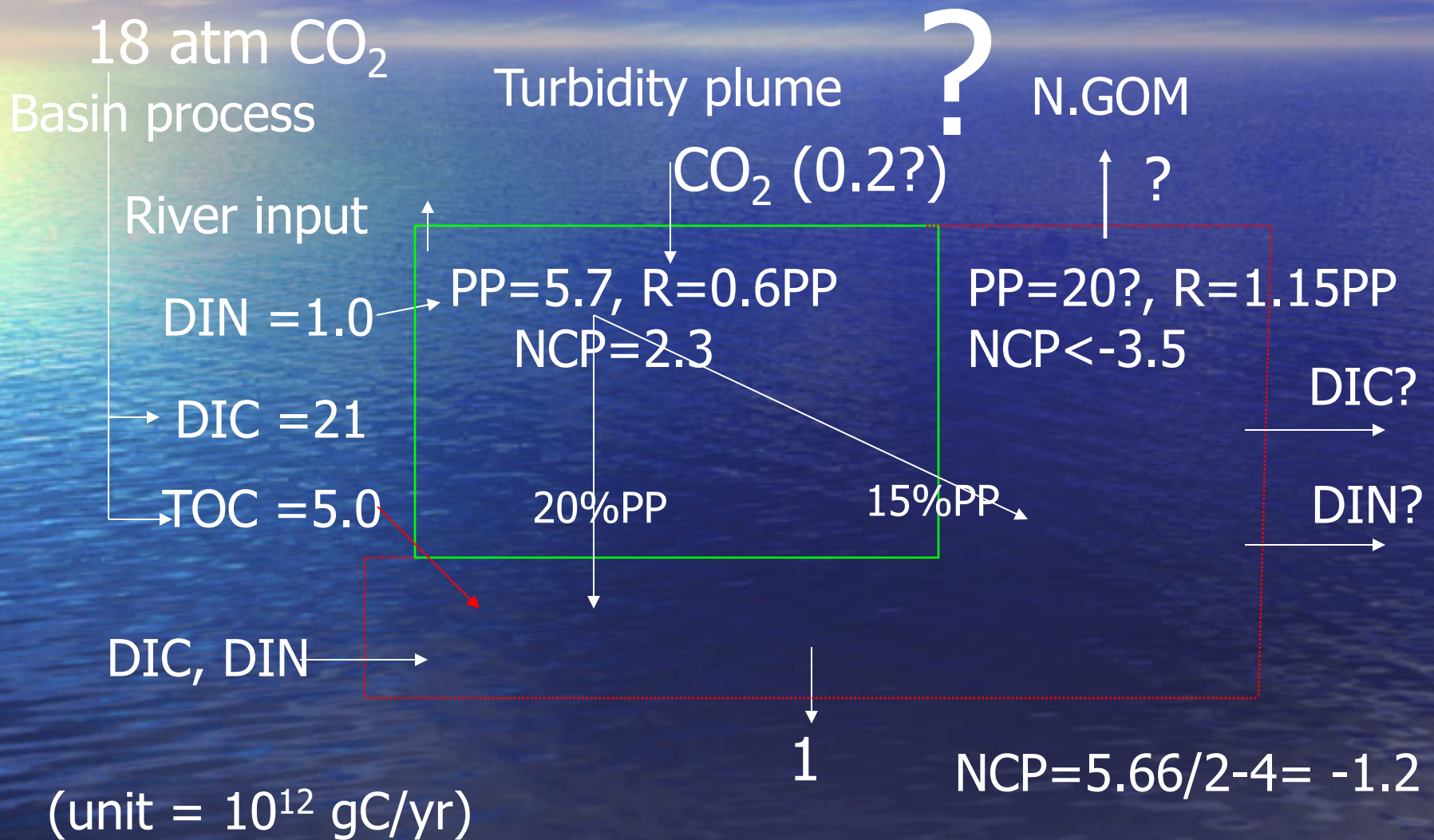
$p\text{CO}_2$ in the East China Sea in summer 1998 during the great flood period
(Results from the Chinese JGOFS, L. Zhang pers. comm.)

$p\text{CO}_2$ in the Pearl River plume

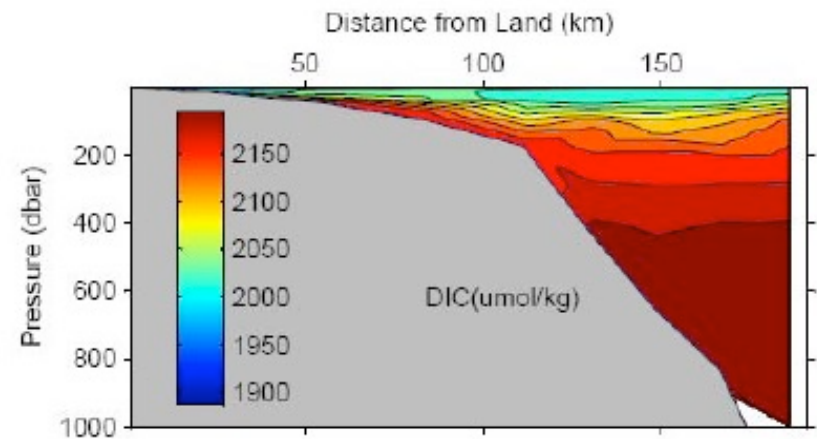
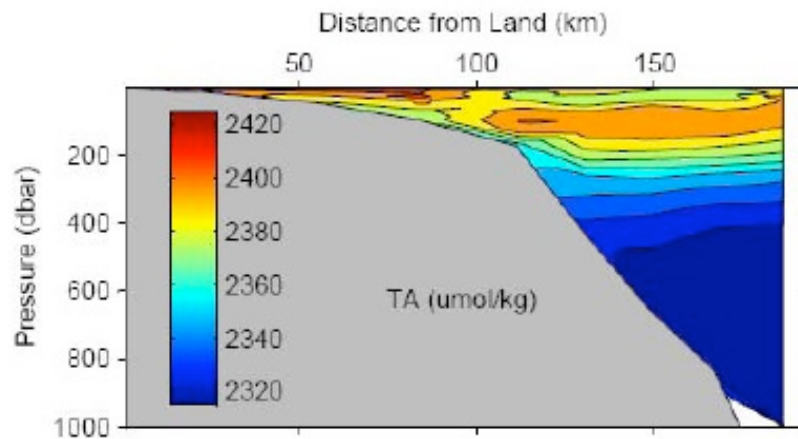


Dai et al. CSR (2008)

Carbon budget in the Northern GOM



GOMECC-Mississippi Transect



- Great DIC enrichment in the bottom water

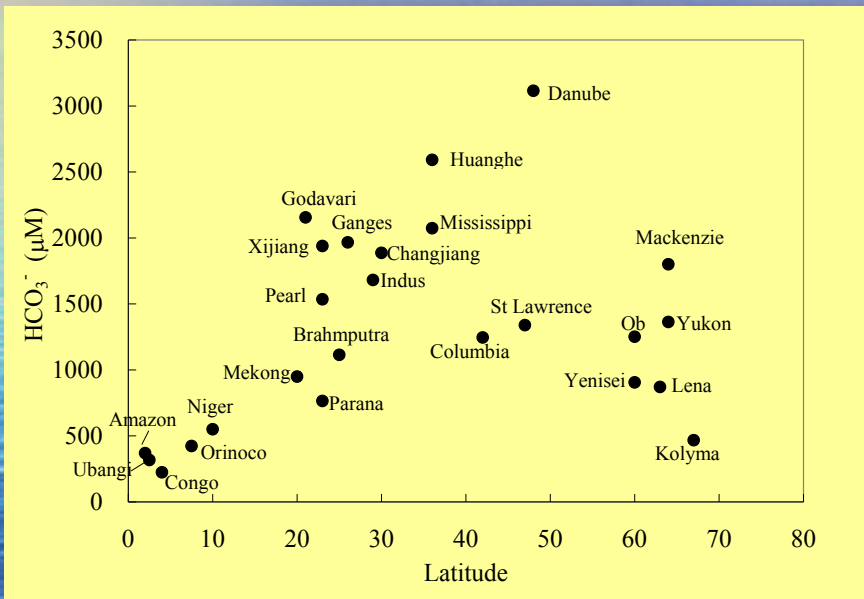
Summary

- Globally, river loadings are heterotrophic, in particular, this is the case for tropical rivers.
- Large river plumes are autotrophic and a sink of CO_2 . This is particularly true for rivers with high anthropogenic DIN loading.
- River HCO_3^- affects coastal TAlk distribution.
- pCO_2 in large river plumes has a great spatial and seasonal variability.
- Potential coccolith bloom may be an important mechanism controlling surface water pCO_2 for some large river plumes.

Thank you!

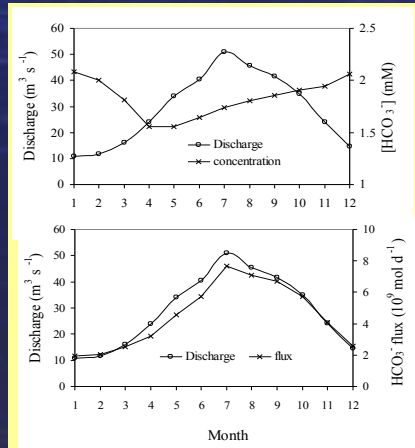
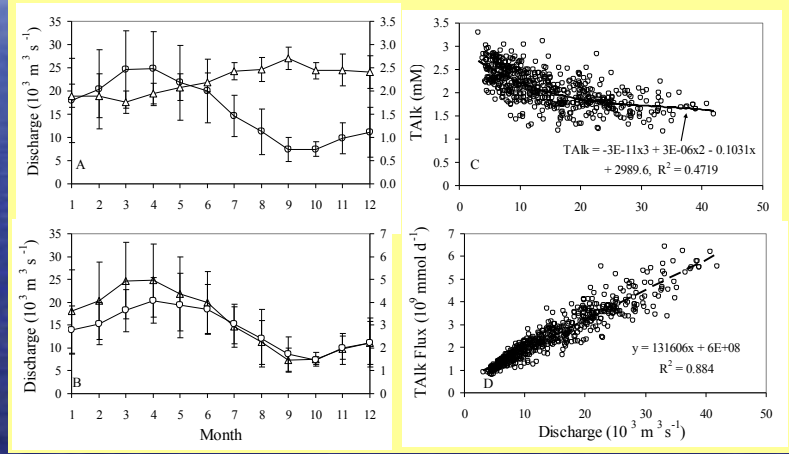
Distribution of riverine HCO_3^- concentration

The Mississippi River



Maximum shifted to north as a result
Of precipitation pattern

From Cai et al. 2008
Continental Shelf Research 28:1538-1549.



The Changjiang (Yangtze River)

Mixing in the Amazon River plume

vs. in the MR plume

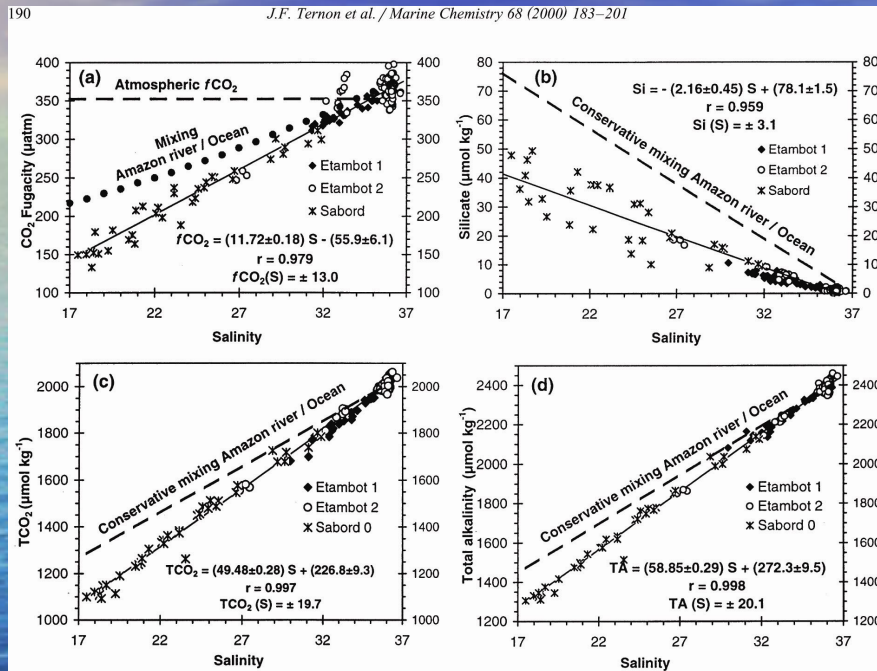
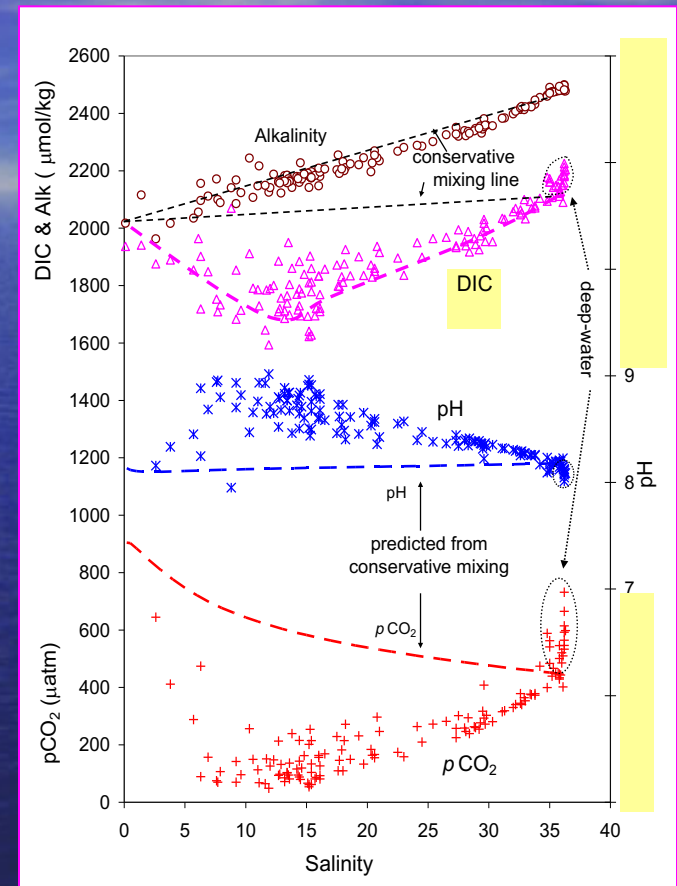


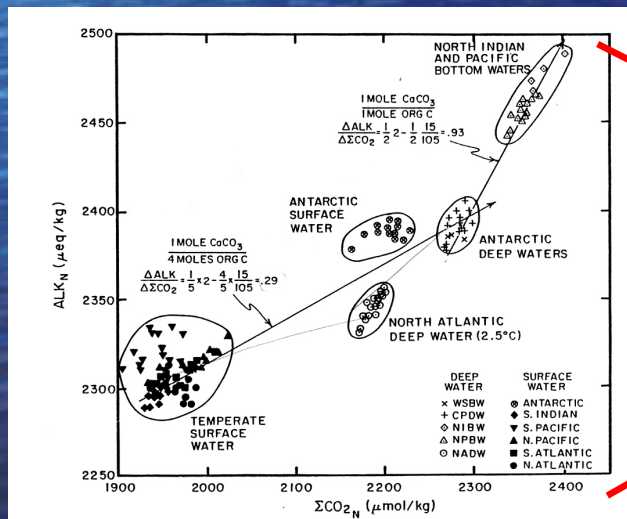
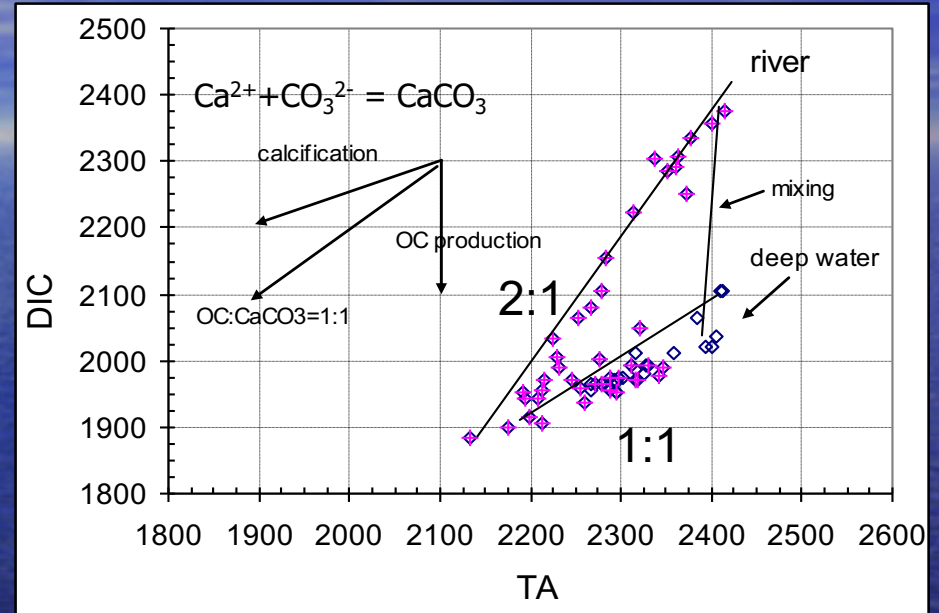
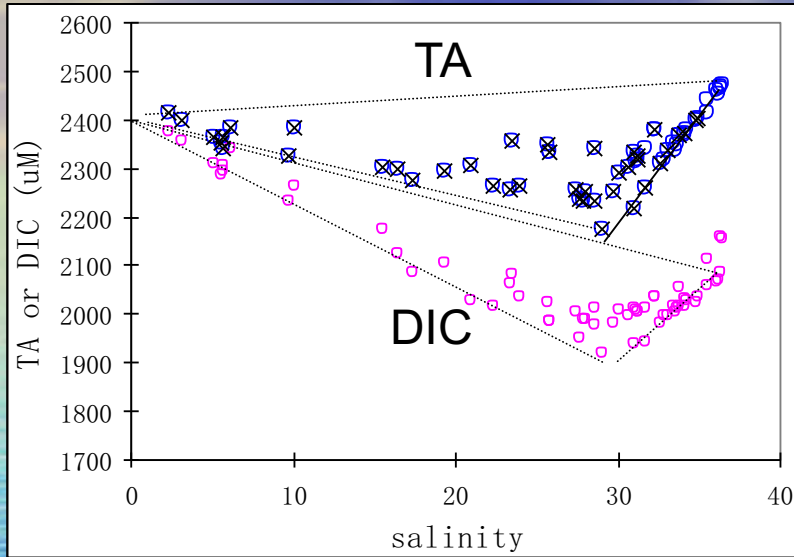
Fig. 5. (a) Relationship between CO_2 fugacity (μatm) and salinity at the sea surface during ETAMBOT 1, ETAMBOT 2 and SABORD. The equation and correlation coefficient, r , are for the regression line (solid line) fitted to the data. $f\text{CO}_2(S)$ indicates the uncertainty in



Amazon River water has low TA & low buffer capacity; in contrast, MR water has high TA and high buffer capacity.

October 2005

mixing



$$\frac{\Delta Alk}{\Delta TCO_2} = \frac{1}{2} \times 2 - \frac{1}{2} \times \frac{15}{105} = 0.93$$

$$\frac{\Delta Alk}{\Delta TCO_2} = \frac{1}{5} \times 2 - \frac{4}{5} \times \frac{15}{105} = 0.28$$

Satellite image of CaCO_3

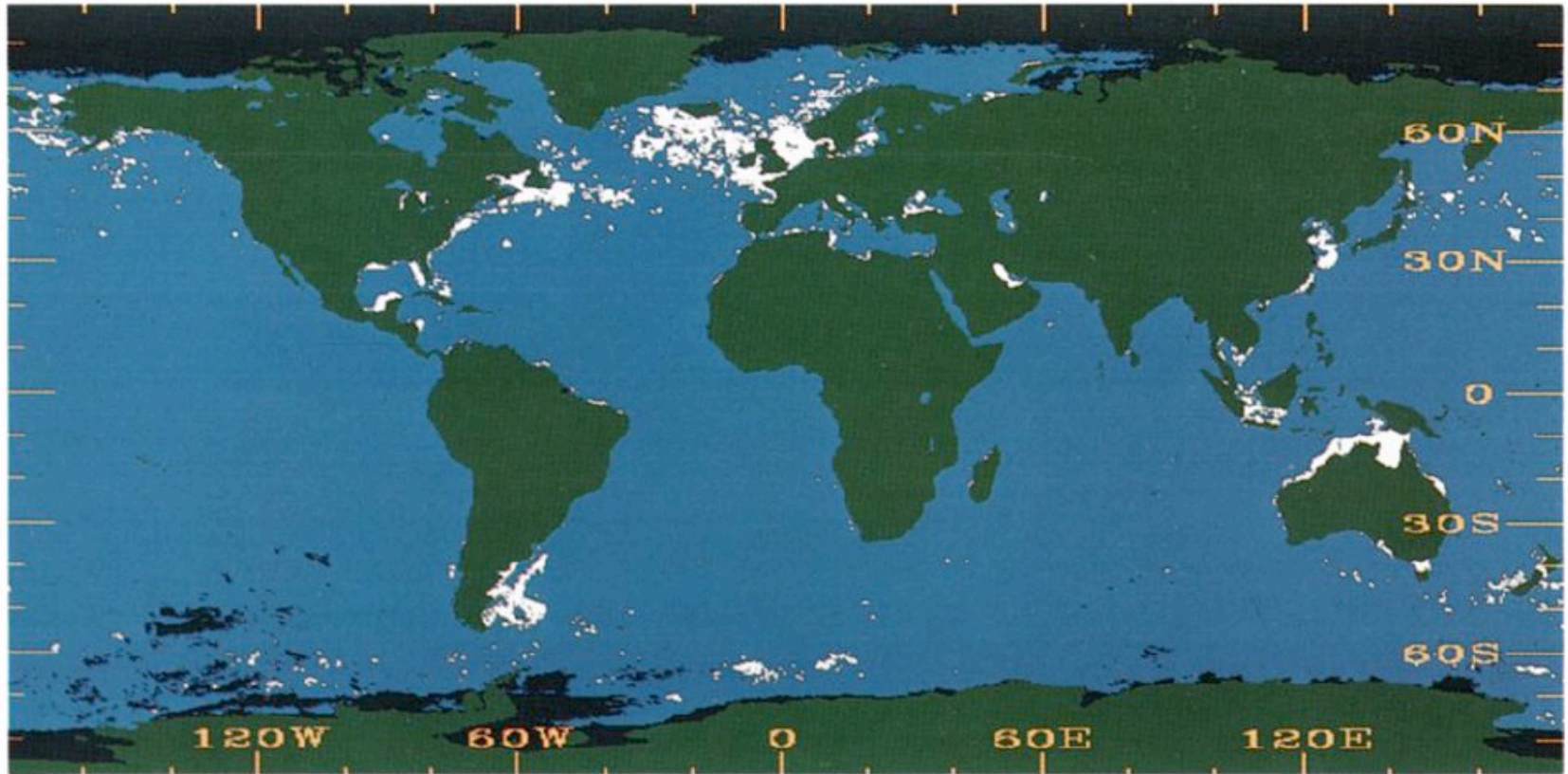
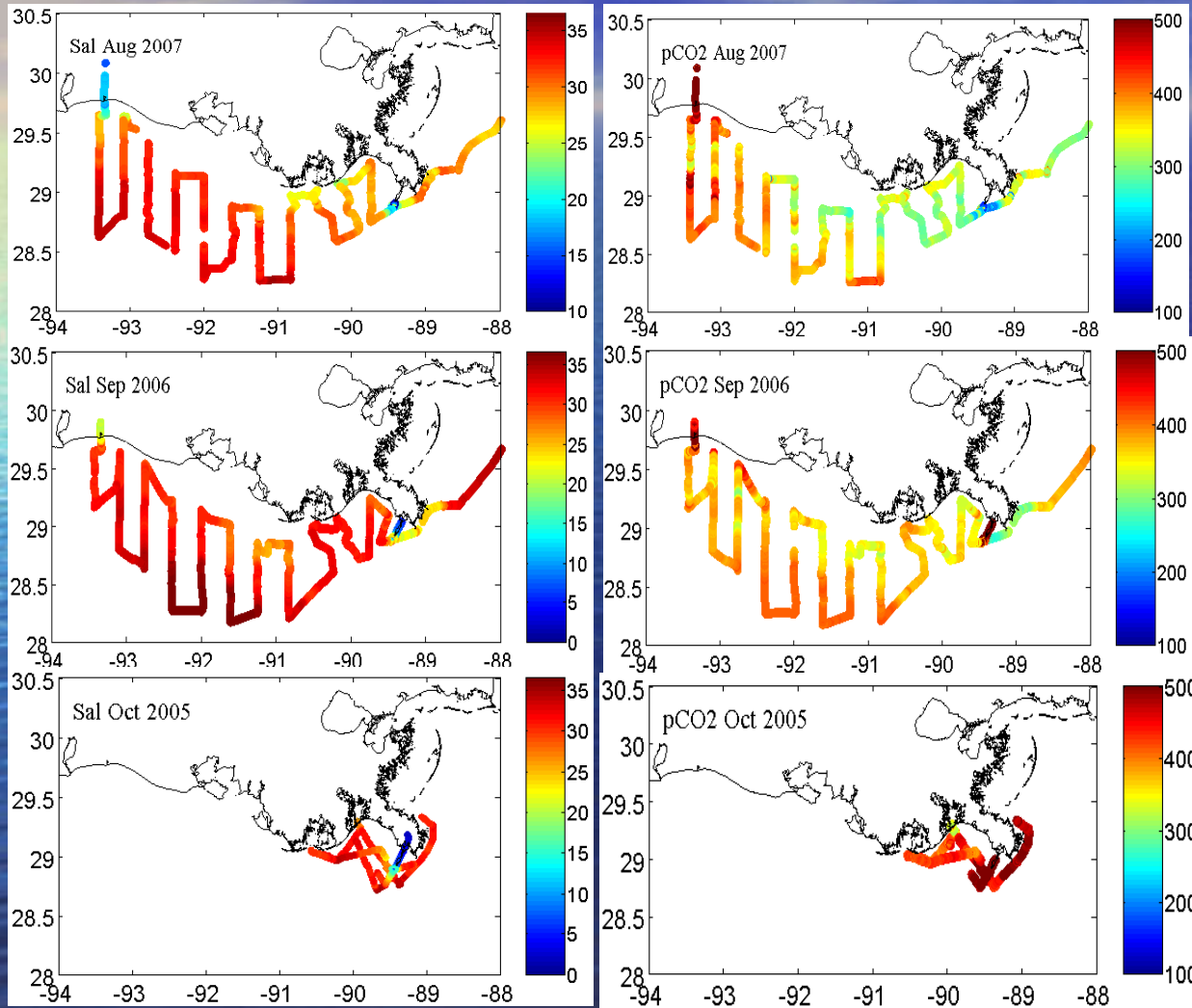


Plate 1. Climatology of classified coccolithophorid blooms (measuring $>4800 \text{ km}^2$) for the world's oceans in CZCS imagery dating from November 1978 to June 1986. The maximum spatial extents of blooms detected during this period are displayed. The coccolithophorid bloom class is white, the noncoccolithophorid bloom class is blue, and the land is green. Black indicates areas lacking image coverage.

salinity

$p\text{CO}_2$ (late summer & fall)



Seasonal variation

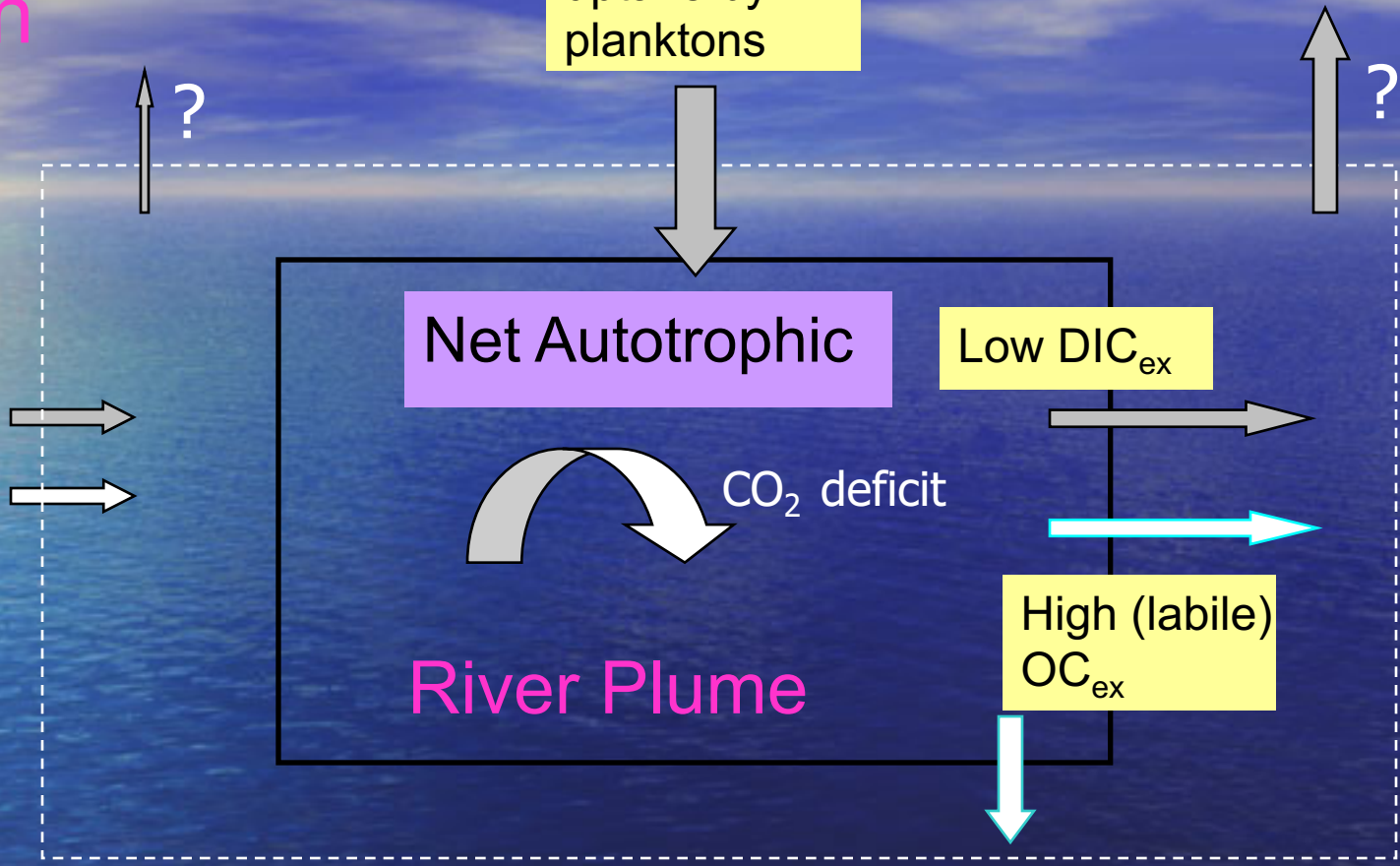
1. Low $p\text{CO}_2$ in spring and early summer
2. High CO_2 in late summer and fall

Large River
Margin

CO_2 sink

uptake by
planktons

River Inputs
TOC, DIC, high DIN



$\text{NCP} > 0$ (net autotrophic) or $\text{GPP} > \text{R}$

$\text{OC exports} > \text{OC imports}$ & $\text{DIC import} > \text{DIC export}$

Results: low DIC, & low $p\text{CO}_2$