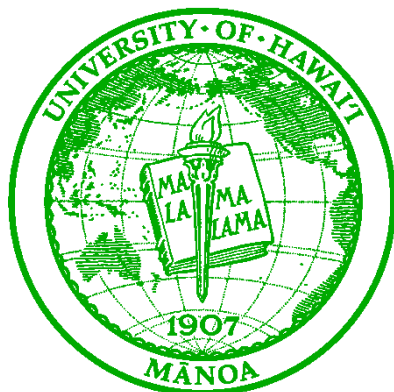


25 years of Hawaii Ocean Time-series carbon flux determinations: Insights into productivity, export, and nutrient supply in the oligotrophic ocean



MATTHEW CHURCH, ROBERT BIDIGARE, JOHN DORE, DAVID KARL,
MICHAEL LANDRY, RICARDO LETELIER, ROGER LUKAS
OCEAN CARBON BIOGEOCHEMISTRY SUMMER WORKSHOP

JULY 2013



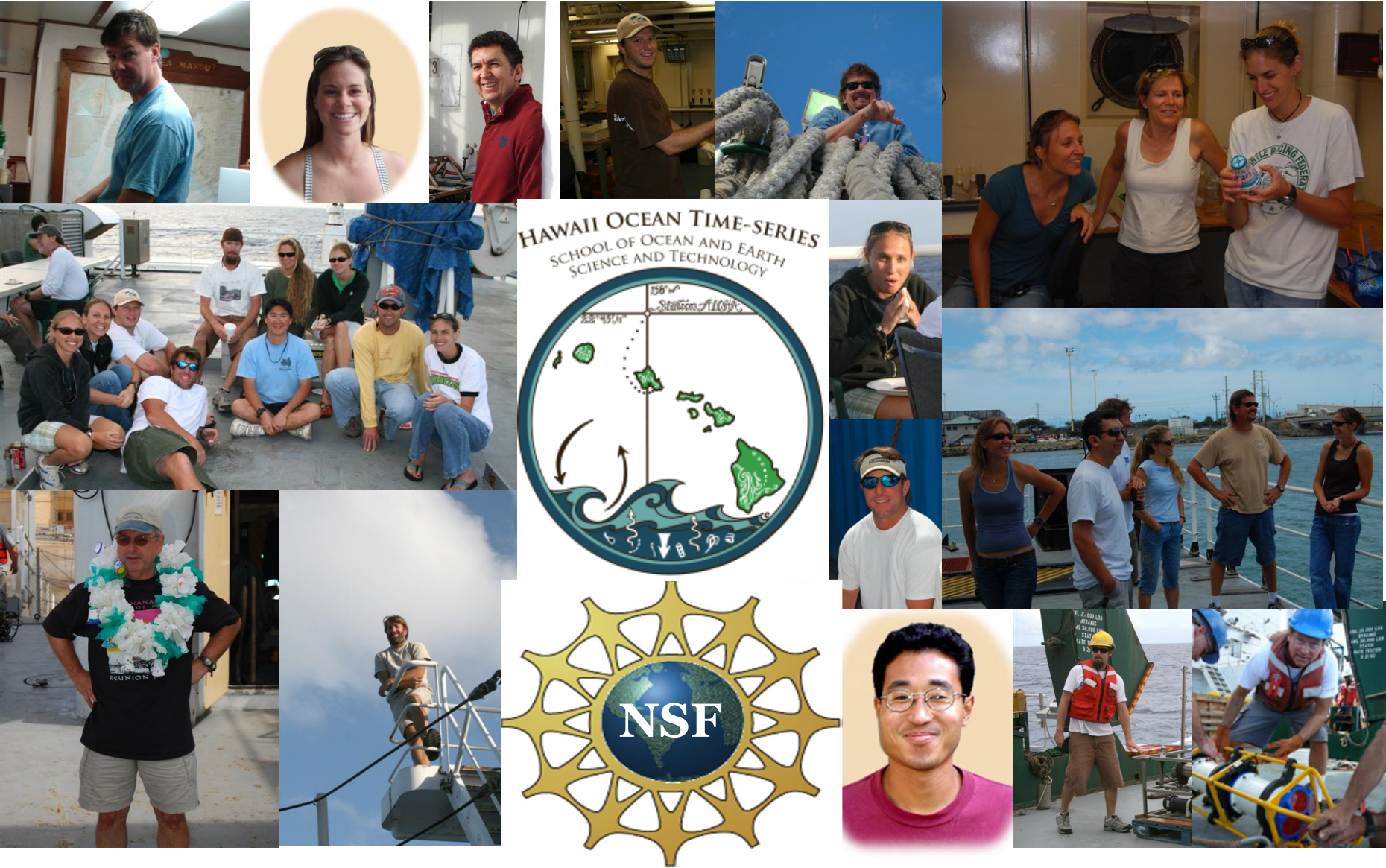
Thank you

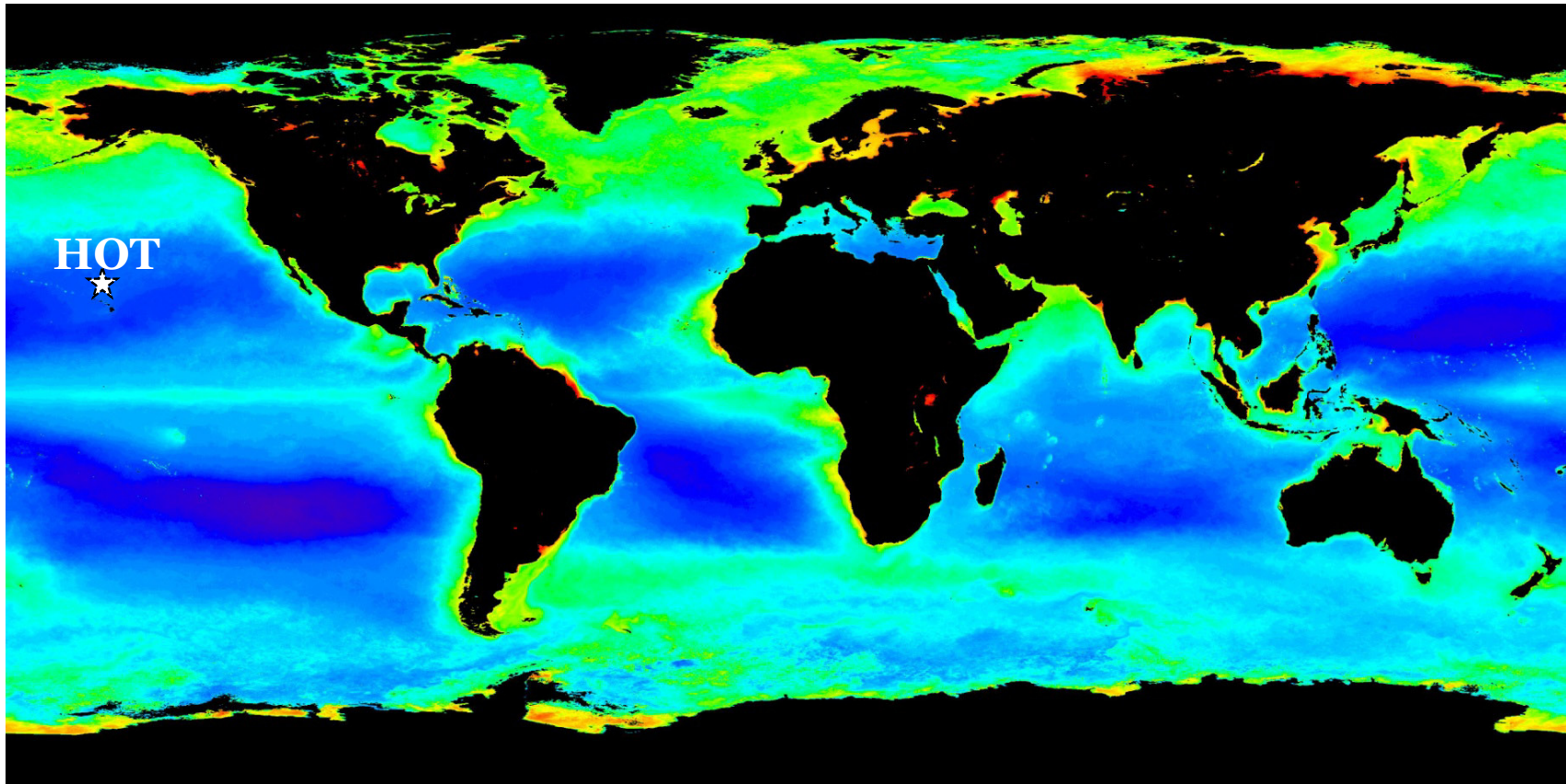


- Craig Carlson (UCSB) and Ricardo Letelier (OSU)
- Ken Johnson (MBARI)
- Heather Benway (WHOI)
- Mary Z. (WHOI)
- The National Science Foundation



A Dedicated HOT Team





- **Subtropical gyres comprise some the largest habitats on this planet**
- **Constraining carbon production and sequestration in these regions is critical to global carbon budgets**
- **Time series programs afford unique opportunities to define the magnitude and pathways of carbon fluxes in the open sea**

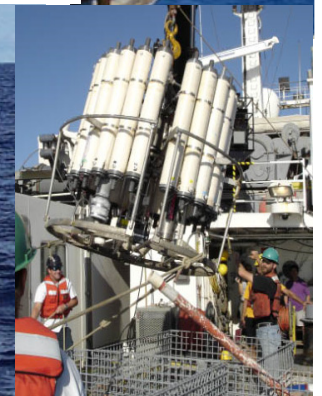
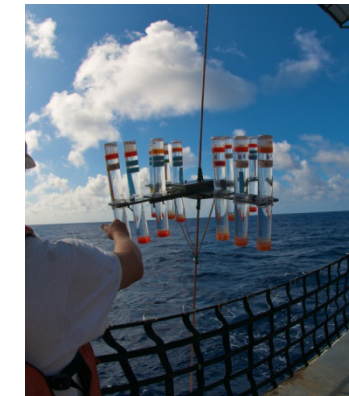
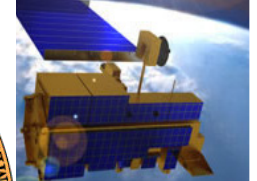
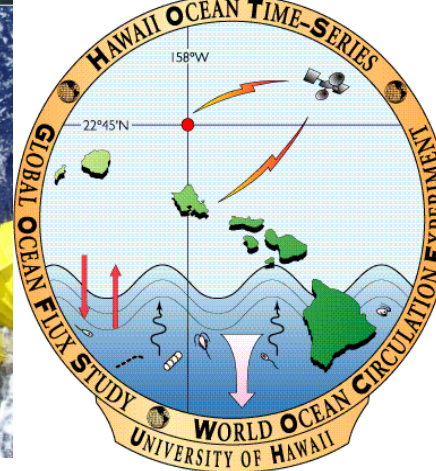
Time, water, and change

- **The complexity of ecosystem dynamics, even in “stable” systems, demands multi-disciplinary, sustained observations.**
- **Implementation and leveraging of remote and autonomous sampling platforms at ocean time series sites is providing new insights into bioelemental cycling in these ecosystems.**

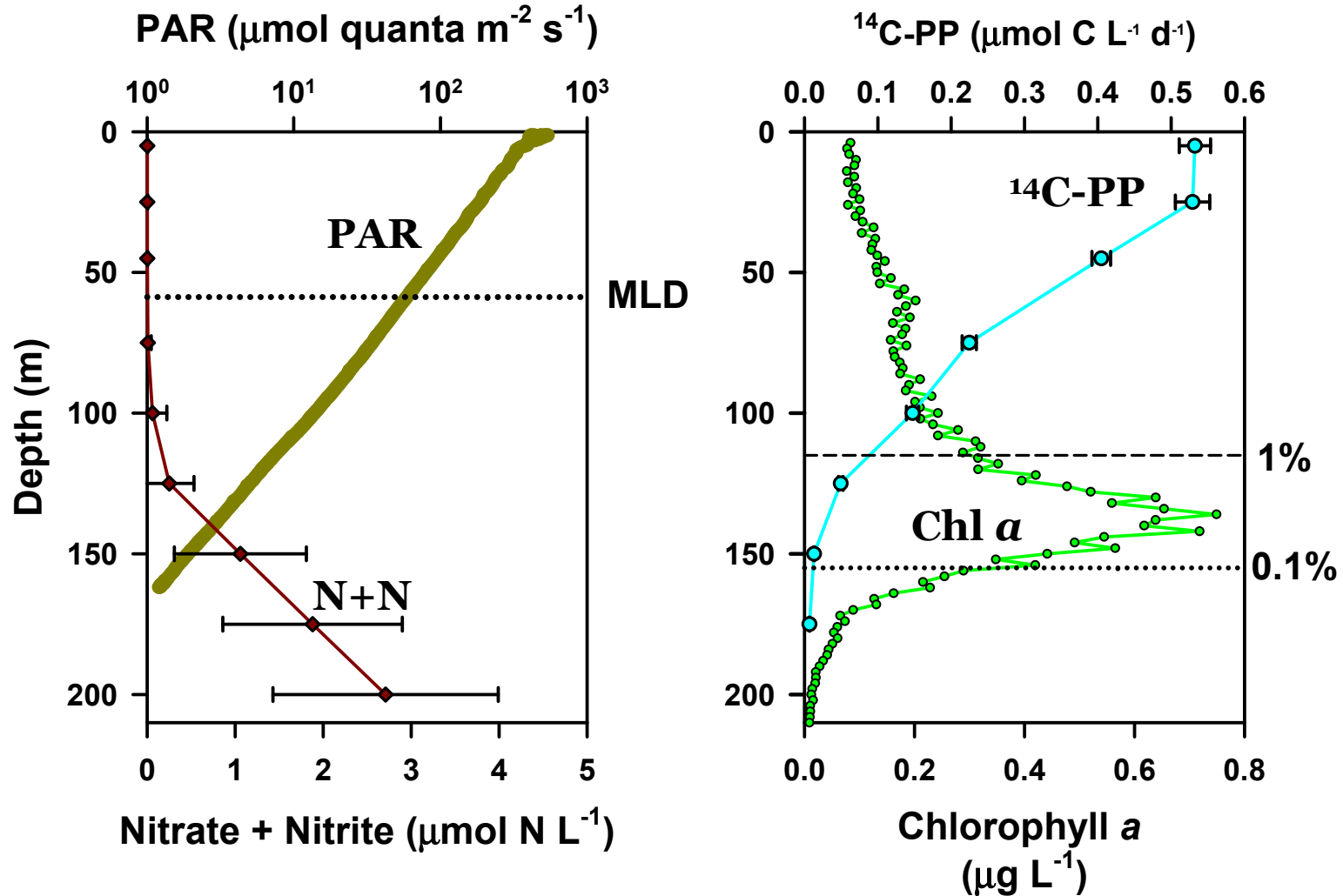


The Hawaii Ocean Time-series (HOT)

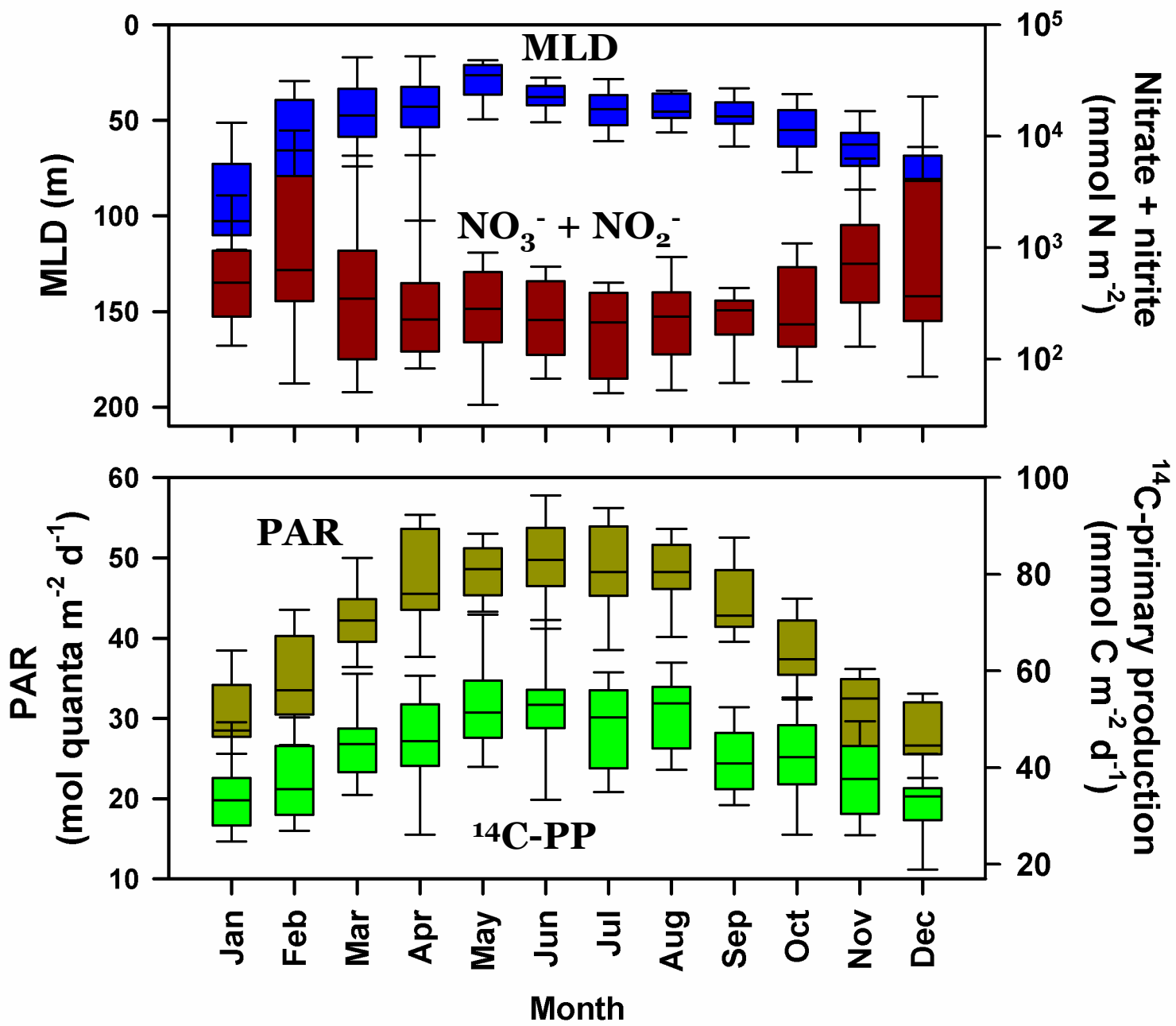
- Near monthly cruises to Station ALOHA since October 1988
- ALOHA is a deep, open ocean (~4800 m) site
- Shipboard and remote (moorings, gliders, floats, and satellites) measurements of ocean biogeochemistry, physics, and plankton ecology
- 4-day cruises, intensive sampling to 1000 m



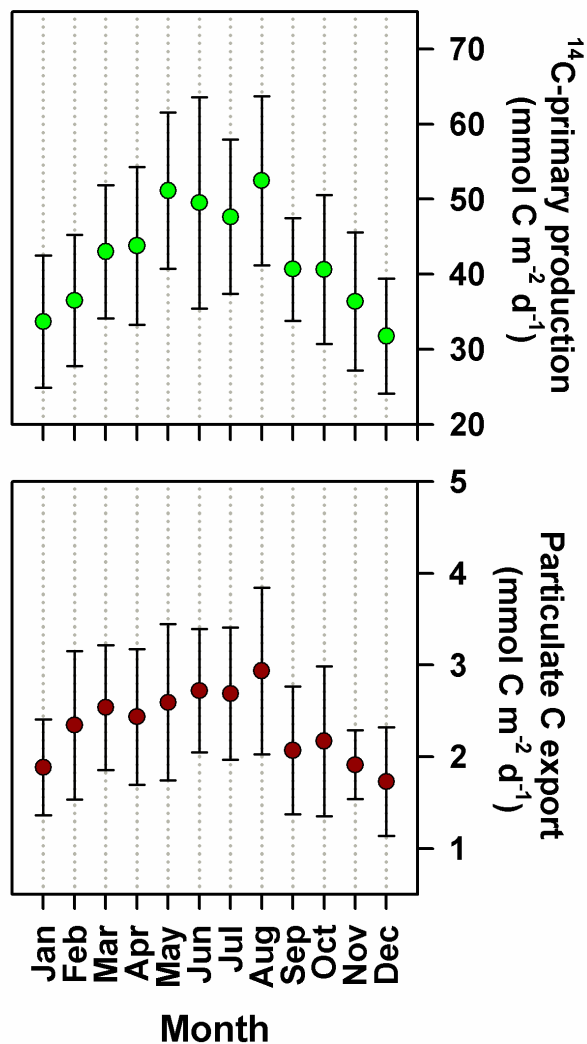
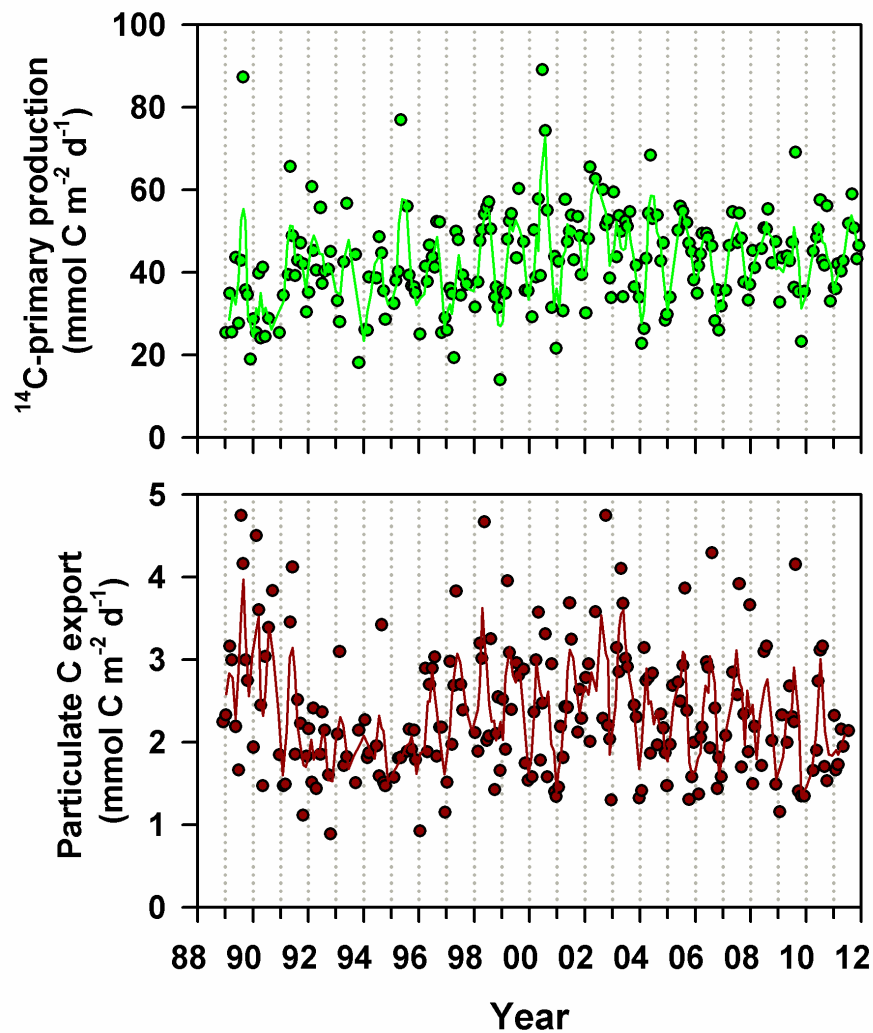
The upper ocean habitat



- Mixed layer ~30-100 m, euphotic zone ~100-125 m
- >65% of the daily carbon fixation occurs in the nutrient-deplete mixed layer



Carbon fixation and particulate carbon export

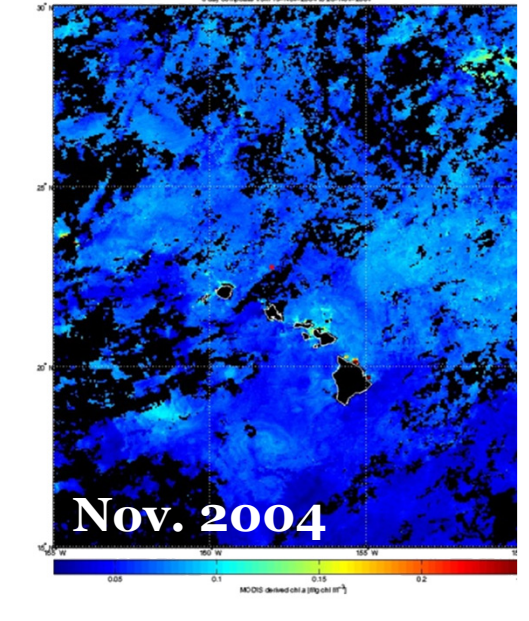
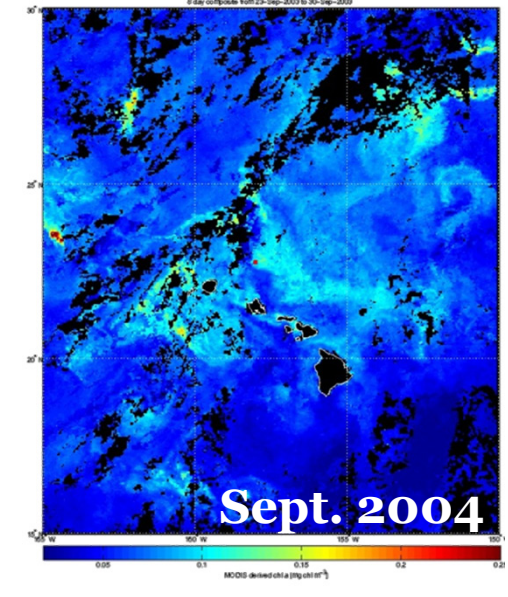
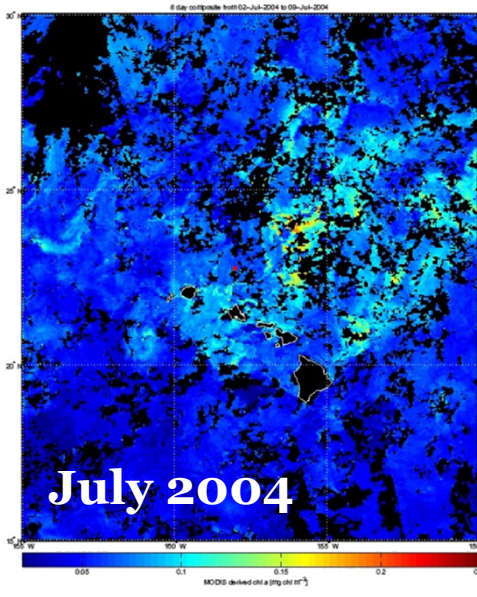
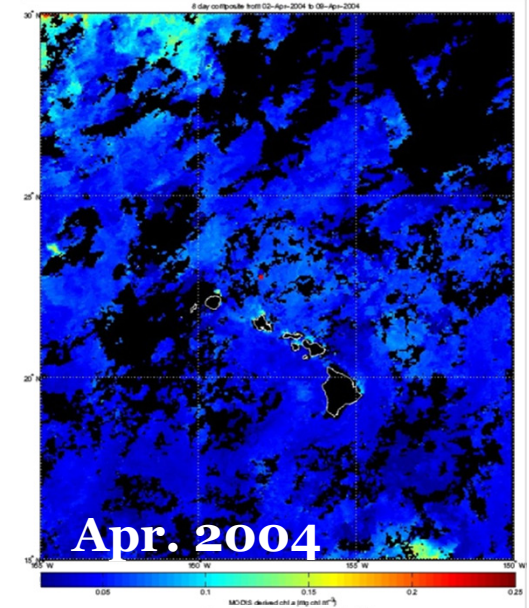
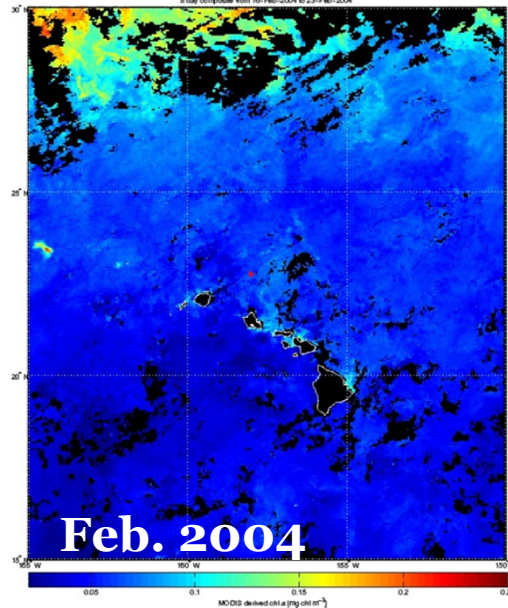
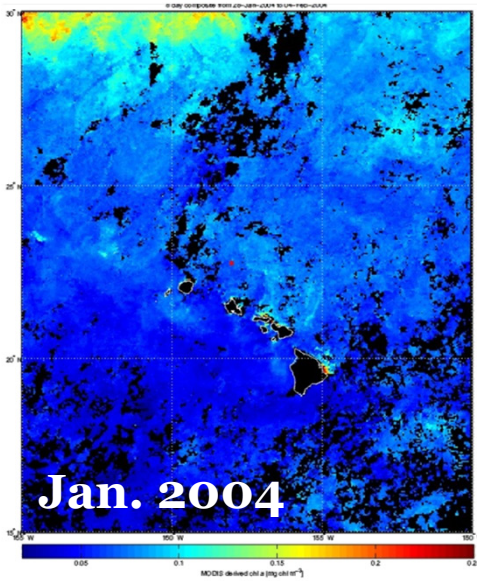


Annually averaged
 ^{14}C -PP:
 $\sim 15 \text{ mol C m}^{-2} \text{ yr}^{-1}$

Annually averaged
PC_{150m}
export:
 $\sim 0.9 \text{ mol C m}^{-2} \text{ yr}^{-1}$

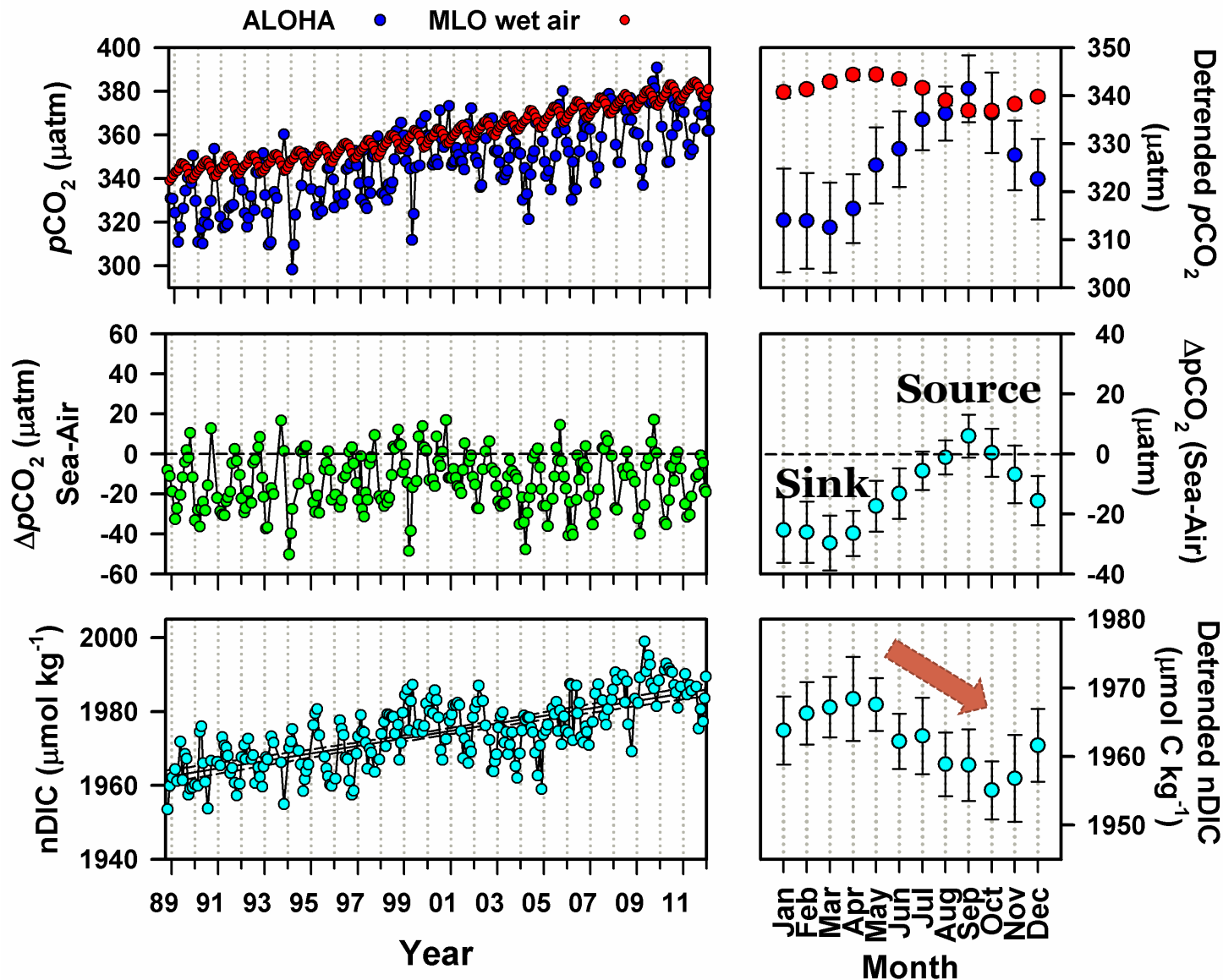
$$\text{PC}_{150 \text{ m}} : ^{14}\text{C}\text{-PP} = 0.04\text{-}0.12$$

The many faces of Station ALOHA



Ricardo Letelier and Angel White (OSU)

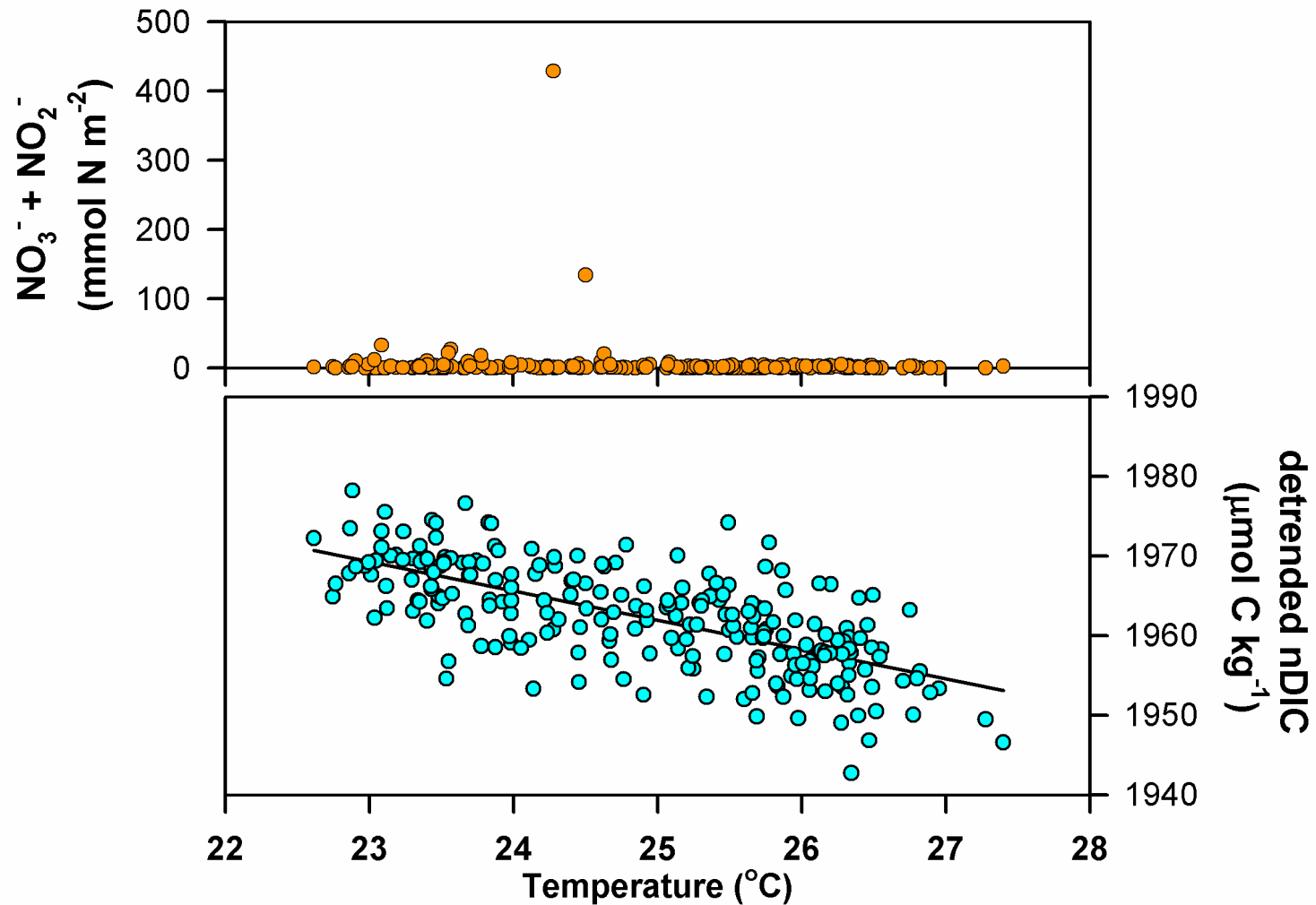
Variability in mixed layer inorganic carbon



- Seasonal variations in $p\text{CO}_2$ largely a function of temperature.
- Most of the year mixed layer CO_2 is undersaturated ($\sim 13 \mu\text{atm}$), becoming a weak source of CO_2 in the late summer.
- Salinity normalized DIC decreases from winter-spring into summer-fall.

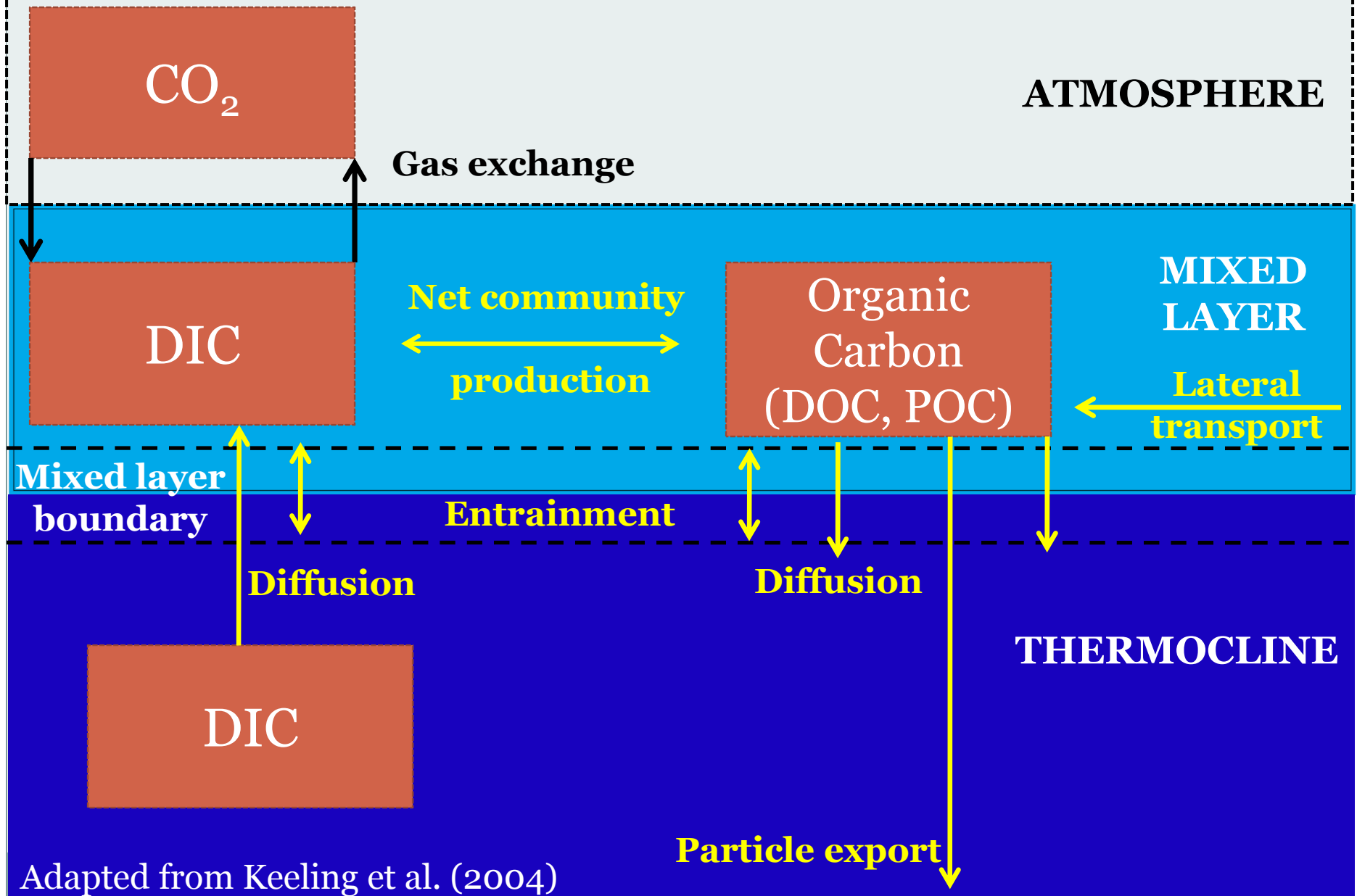
What processes control the summertime drawdown of DIC at Station ALOHA?

Spring-fall drawdown of DIC in the absence of nitrate is a common feature of the subtropical gyres



(Michaels et al. 1994, Bates et al. 1996, 1998, Gruber et al. 1998, Lee et al. 2000, Karl et al. 2003, etc.)

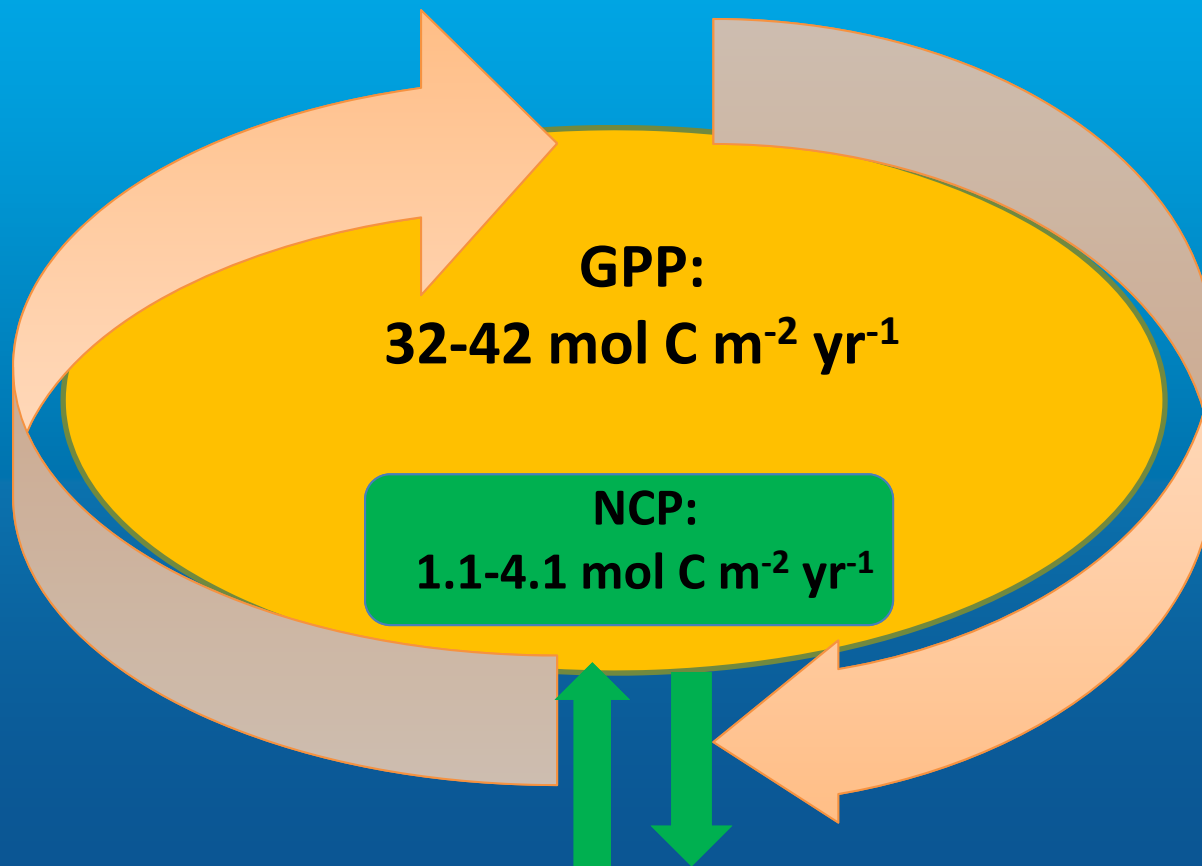
Biological and physical processes controlling variability in carbon inventories and fluxes





Quantifying net community production in the open sea is difficult

1. Small net changes superimposed on large background pools and fluxes
2. Need to accurately quantify impacts of biology (production and respiration) and physics (vertical and lateral transport, air-sea exchange)
3. Ecology matters – nutrient fluxes and growth efficiencies



GPP:
32-42 mol C m⁻² yr⁻¹

NCP:
1.1-4.1 mol C m⁻² yr⁻¹

$$\text{GPP} - \text{R} = \text{NCP}$$

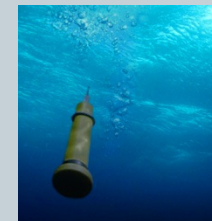
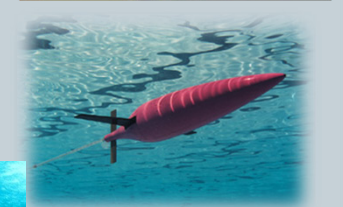
Station ALOHA

$$\text{NCP} : \text{GPP} = \\ 0.03-0.13$$

Measurements of net community production at Station ALOHA



- **Mixed layer DIC and $^{13}\text{C}/^{12}\text{C}$**
- **Seasonal evolution of dissolved oxygen, O_2 :inert gas ratios, and oxygen isotopes**
- **Nitrogen-based determinations of new production; over annual scales \approx NCP**
- **Vertical transport of organic matter export (POC, DOC, and migrant flux)**



Estimates of NCP at ALOHA abound...

SCIENCE VOL 316 27 APRIL 2007
**Revisiting Carbon Flux Through
the Ocean's Twilight Zone**

Ken O. Buesseler,^{1*} Carl H. Lamborg,¹ Philip W. Boyd,² Phoebe J. Lam,¹ Thomas W. Trull,⁴
Robert R. Bidigare,⁴ James K. B. Bishop,^{5,6} Karen L. Casciotti,¹ Frank Dehairs,⁷ Marc Elskens,⁷
Makio Honda,⁸ David M. Karl,⁴ David A. Siegel,⁹ Mary W. Silver,¹⁰ Deborah K. Steinberg,¹¹
Jim Valdes,^{1,2} Benjamin Van Mooy,¹ Stephanie Wilson¹¹

Limnol. Oceanogr., 50(2), 2005, 646-657
© 2005, by the American Society of Limnology and Oceanography, Inc.

Biogeochemical cycling in the oligotrophic ocean: Redfield and non-Redfield models

James Robert Christian

Deep-Sea Research II, Vol. 43, No. 2-3, pp. 539-568, 1996

**Seasonal and interannual variability in primary production and
particle flux at Station ALOHA**

D. M. KARL,^{*} J. R. CHRISTIAN,^{*} J. E. DORE,^{*} D. V. HEBEL,^{*}
R. M. LETELIER,^{*} L. M. TUPAS^{*} and C. D. WINN^{*}

NATURE | VOL 389 | 30 OCTOBER 1997

**Experimental determination of
the organic carbon flux from
open-ocean surface waters**

S. Emerson^a, P. Quay^a, D. Karl^b, C. Winn^a, L. Tupas^a
& M. Landry^a

*...nutrient mass balances
and carbon export*

NATURE · VOL 373 · 19 JANUARY 1995

**Ecosystem changes in the
North Pacific subtropical
gyre attributed to
the 1991-92 El Niño**

D. M. Karl, R. Letellier, D. Hebel, L. Tupas,
J. Dore, J. Christian & C. Winn

Deep-Sea Research I 48 (2001) 2595-2611

A time-series study of particulate matter export in the North
Pacific Subtropical Gyre based on ²³⁴Th: ²³⁸U disequilibrium

Claudia Benitez-Nelson^{a,*}, Ken O. Buesseler^b, David M. Karl^a, John Andrews^b

Deep-Sea Research II 53 (2006) 698-717

On the relationships between primary, net community, and
export production in subtropical gyres

Holger Brix^{a,*}, Nicolas Gruber^a, David M. Karl^b, Nicholas R. Bates^c

NATURE | Vol 465 | 24 June 2010

**Nitrate supply from deep to near-surface waters of
the North Pacific subtropical gyre**

Kenneth S. Johnson¹, Stephen C. Riser² & David M. Karl³

Estimates of net community production at ALOHA

	Method of Determination	Rate (\pm stdev) $\text{mol C m}^{-2} \text{ yr}^{-1}$	Period	References
<i>O₂ based approaches</i>	Mixed Layer O ₂ : Ar	1.4 - 3.7 (\pm 1.0)	1992–2008	Emerson et al. (1997); Hamme and Emerson (2006); Juranek and Quay (2005); Quay et al. (2010)
	Mooring O ₂	4.1 (\pm 1.8)	2005	Emerson et al. (2008)
	Sub-mixed layer float profiles	1.1 - 1.7 (\pm 0.2)	2003-2010	Riser and Johnson (2008)
	Sub-mixed layer glider surveys	0.9 (\pm 0.1)	2005	Nicholson et al. (2008)
<i>C-based approaches</i>	Mixed layer ¹³ C/ ¹² C and DIC dynamics	1.6 \pm 0.9	1990-1995	Emerson et al. (1997)
		2.7 \pm 1.4	1994-1999	Quay and Stutsman (2003)
		2.8 \pm 0.8	1988-2002	Keeling et al. (2004)
<i>Passive and active OM fluxes</i>	Sediment traps (150 m)	0.9 (\pm 0.3)	1988-2012	HOT core data
	²³⁴ Th deficits	1.5 (\pm 0.8)	1999-2000	Benitez-Nelson et al. (2001)
	DOC export	0.4	2002-2012	HOT core data
	Zooplankton-mediated	0.1 (\pm 0.09)	1994-2005	Al-Mutairi and Landry (2001) Hannides et al. (2009)

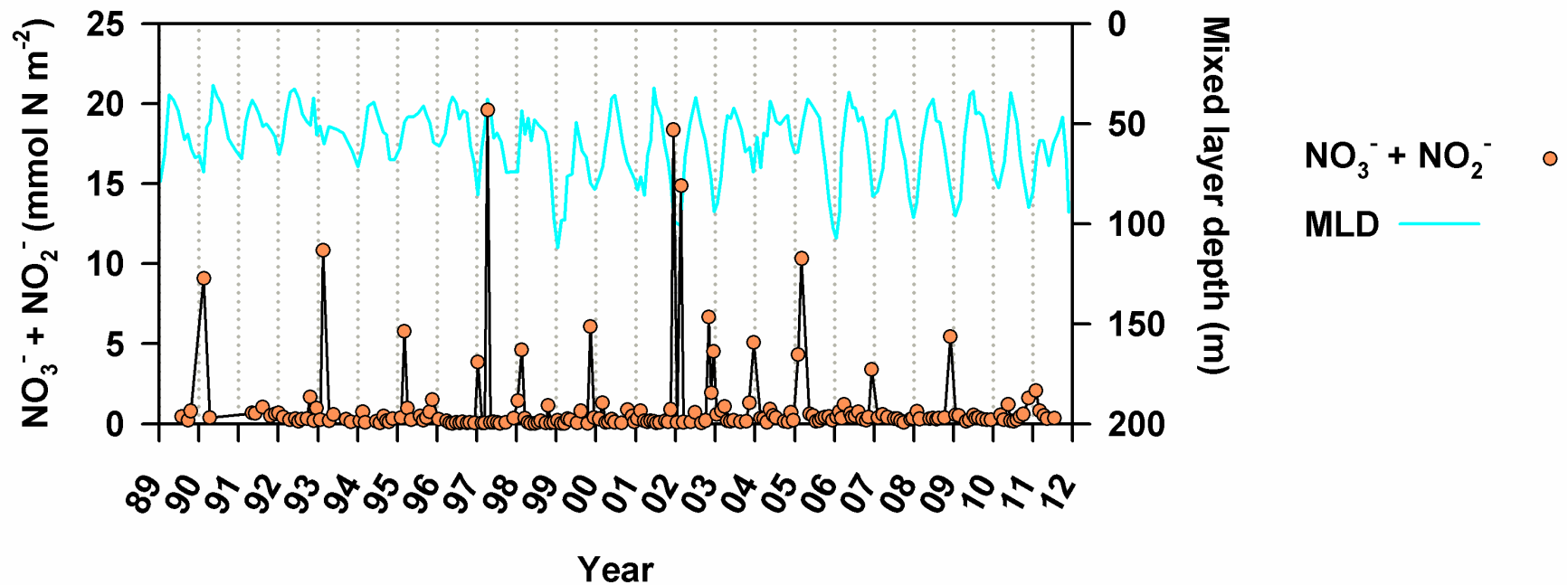
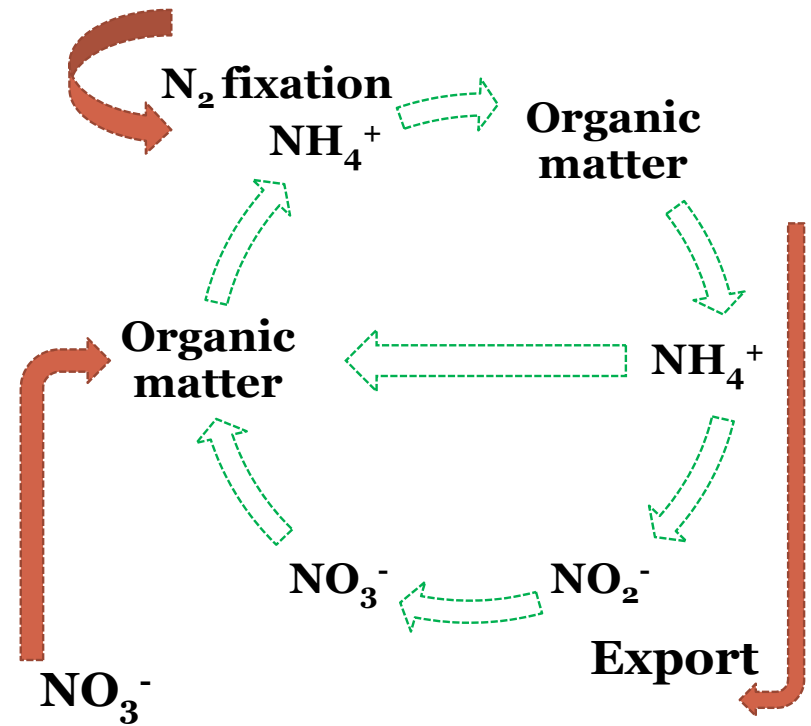
NCP ranges 1.1-4.1 mol C m⁻² yr⁻¹, averaging ~2 mol C m⁻² yr⁻¹
Many of these estimates have uncertainties of ~30-60%

The devil is in the details...

- **Diagnostic models depend on accuracy of :**
 - **Air-sea flux ($\pm 30\%$) - limited by gas exchange parameterization**
 - **Lateral transport ($\pm 50-70\%$) – limited by horizontal velocities**
 - **Vertical entrainment/diffusion ($\pm 50\%$) - poor constraint on K_z**
- **Some estimates based on mixed layer dynamics (i.e. O_2 : Ar or $^{13}C/^{12}C$) while others based on sub-mixed layer (seasonal evolution of O_2)**
- **O_2 based approaches require appropriate PQ and RQ**
- **Sediment trap biases**
- **Particles and DOC**
- **Nutrient sources supporting NCP**

Where do nutrients come from to support NCP?

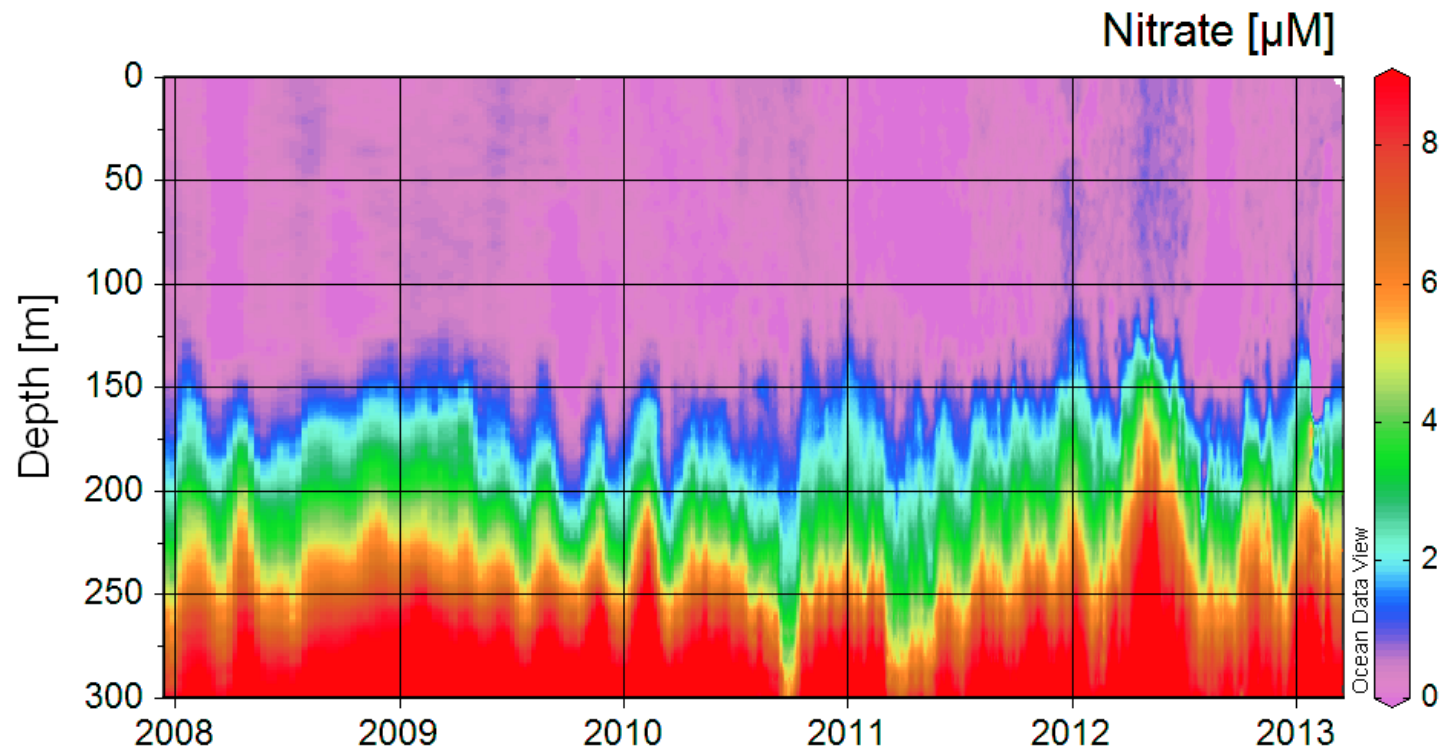
- Physical :
 - ❖ Mixing, upwelling, diffusion, advection, etc.
 - ❖ NO_3^- supported new production
- Biological:
 - ❖ N_2 fixation ($\text{N}_2 \rightarrow \text{NH}_3$)



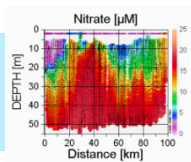
LETTERS

Nitrate supply from deep to near-surface waters of the North Pacific subtropical gyre

Kenneth S. Johnson¹, Stephen C. Riser² & David M. Karl³

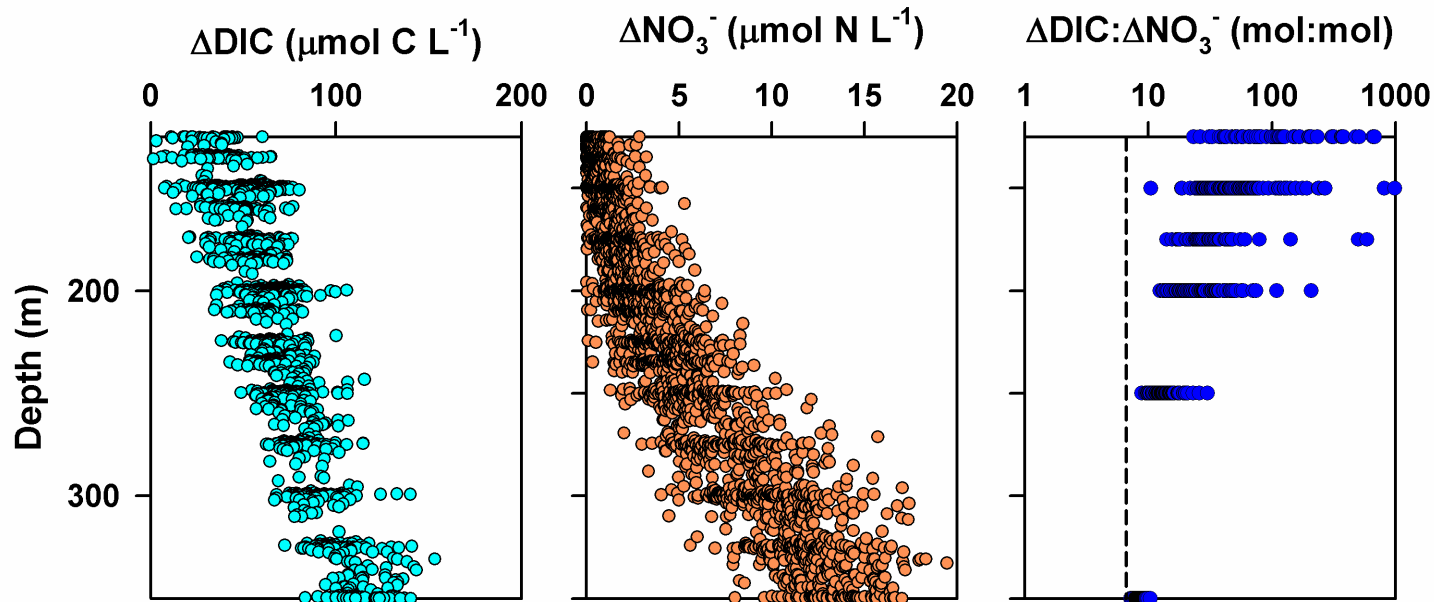


Annual N supply: $>88 \text{ mmol N m}^{-2} \text{ yr}^{-1}$ ($0.6 \text{ mol C m}^{-2} \text{ yr}^{-1}$)

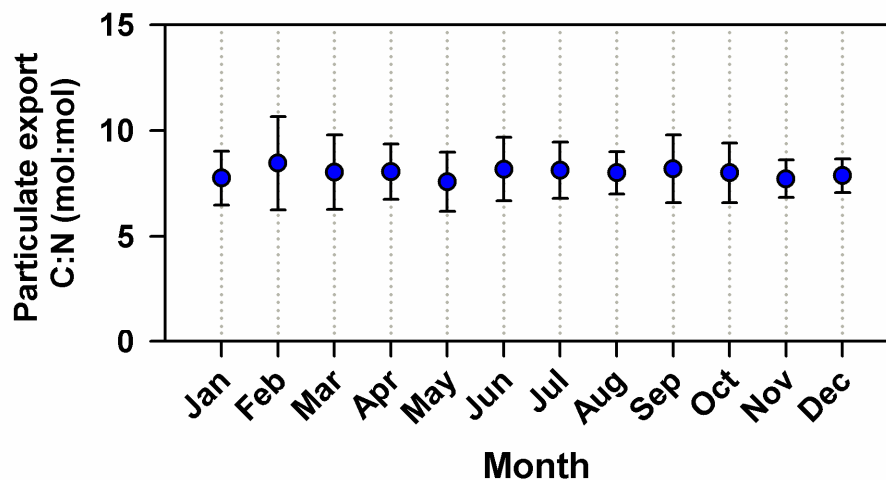


Courtesy of Ken Johnson, MBARI Chemical Sensor Lab

Vertical transport of nutrients by physical processes introduces C-enriched waters



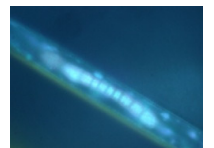
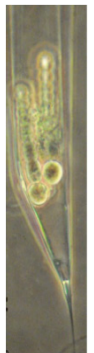
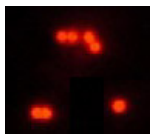
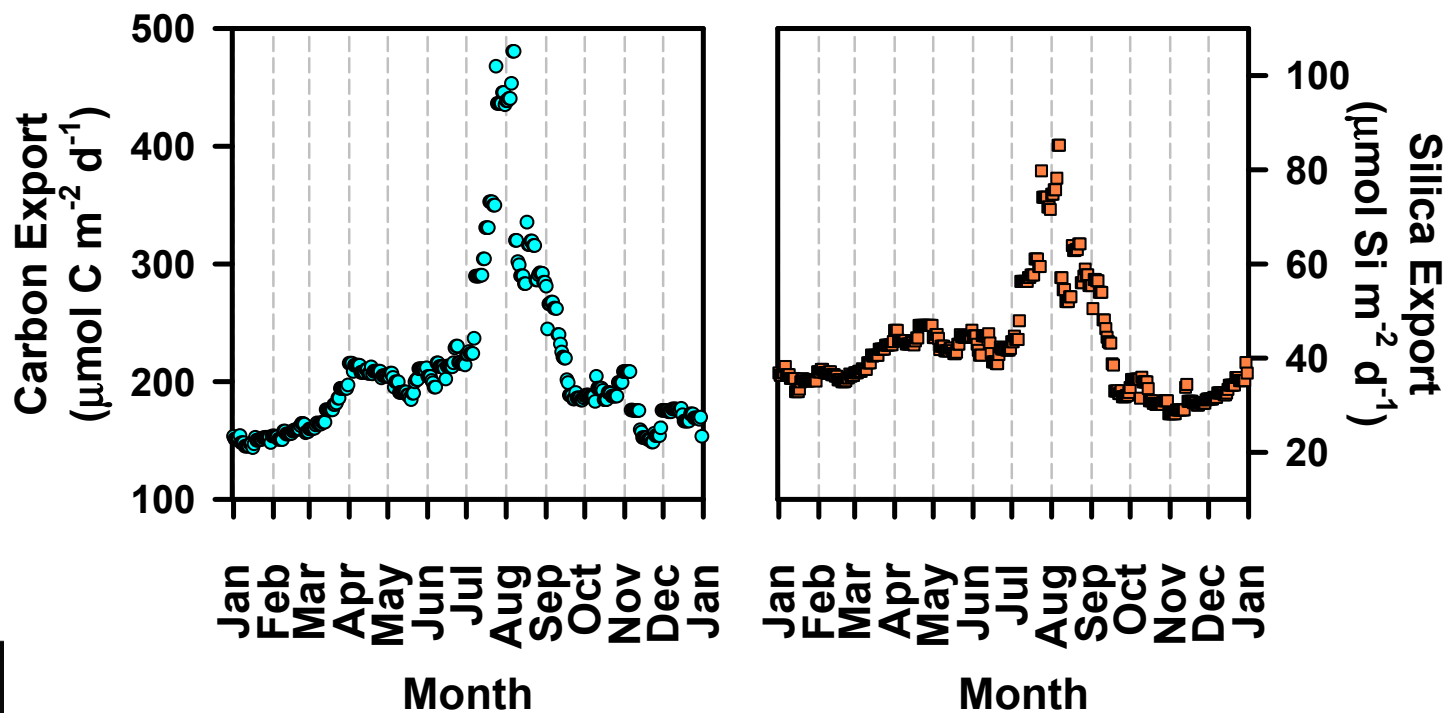
Sub-euphotic zone waters have $\Delta\text{C}:\Delta\text{N}$ ratio of $\sim 25:1$ to $>400:1$



Sinking particles C:N $\sim 7:1$

Biological N supply to the ocean: N₂ fixation

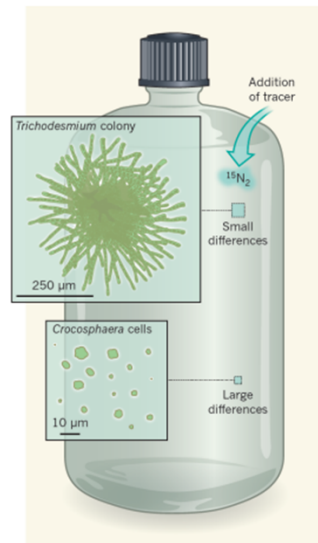
- N₂ fixation estimated to fuel ~50% of particulate export in the subtropical N. Pacific
- Numerous taxa of N₂ fixing microorganisms
- N₂ fixation supported diatom-driven export estimated to contribute ~35% of the annual C-flux to the deep sea



290 | NATURE | VOL 488 | 16 AUGUST 2012

The trouble with the bubble

ANGELIQUE E. WHITE



PLOS one September 2010 | Volume 5 | Issue 9 | e12583

Methodological Underestimation of Oceanic Nitrogen Fixation Rates

Wiebke Mohr*, Tobias Großkopf, Douglas W. R. Wallace, Julie LaRoche

only 40% of the maximum rate measured in the incubations to which $^{15}\text{N}_2$ -enriched seawater had been added. In other words, for the 12-h incubation period under the described experimental conditions, the N_2 fixation rate was underestimated by 60% when the $^{15}\text{N}_2$ was introduced as a gas bubble. In contrast, in both the isotopic equilibration and the culture experiments, the concentration of dissolved $^{15}\text{N}_2$ remained stable at the predicted value throughout the 24 h in incubations to which $^{15}\text{N}_2$ -enriched water was added.

16 AUGUST 2012 | VOL 488 | NATURE | 361

Doubling of marine dinitrogen-fixation rates based on direct measurements

Tobias Großkopf^{1*}, Wiebke Mohr^{1*†}, Tina Baustian¹, Harald Schunck¹, Diana Gill¹, Marcel M. M. Kuypers², Gaute Lavik², Ruth A. Schmitz³, Douglas W. R. Wallace⁴ & Julie LaRoche^{1†}

Our data show that in areas dominated by *Trichodesmium*, the established method underestimates N_2 -fixation rates by an average of 62%. We also find that the newly developed method yields N_2 -fixation rates more than six times higher than those from the established method when unicellular, symbiotic cyanobacteria and γ -proteobacteria dominate the diazotrophic community. On the

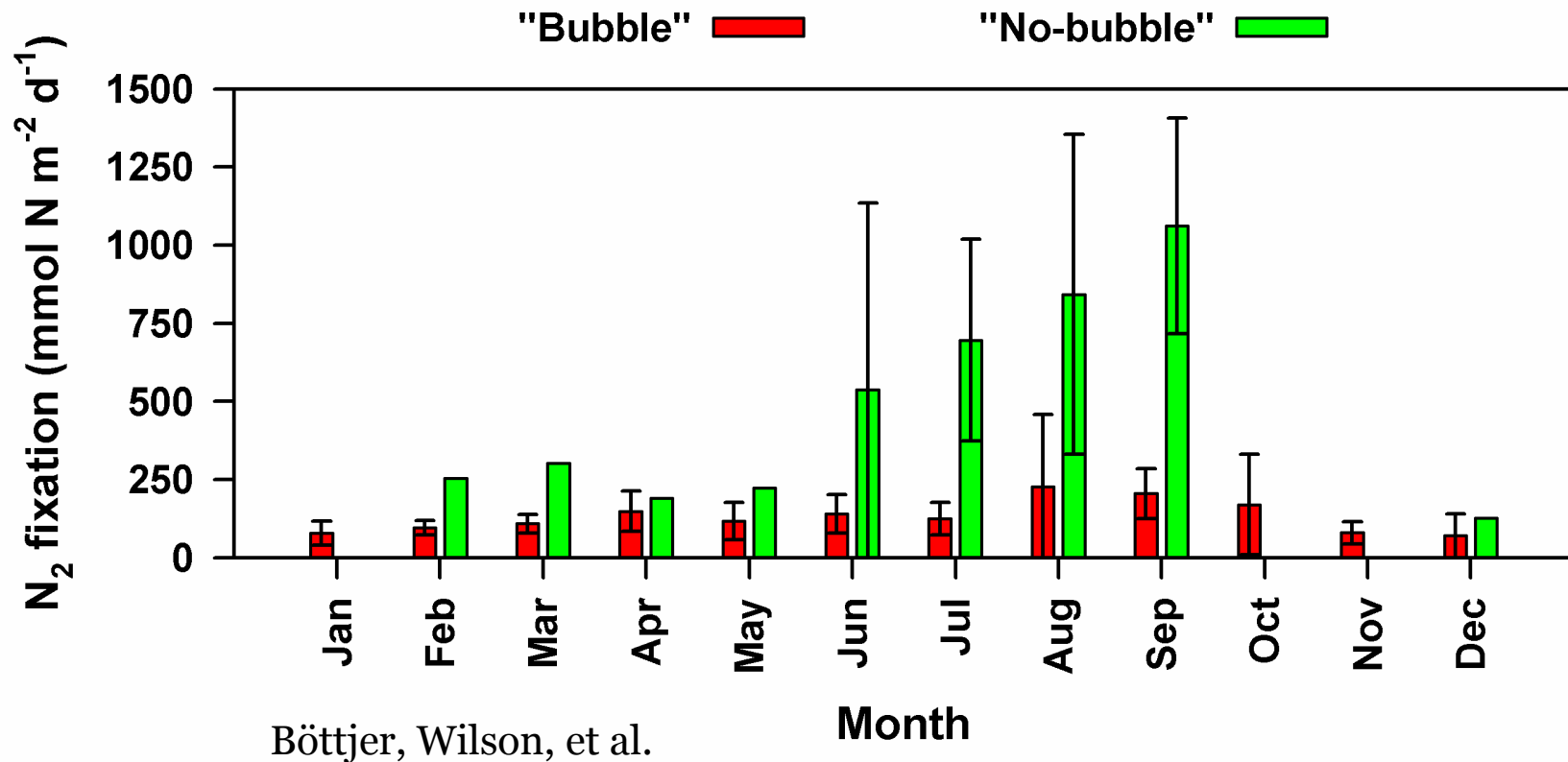
Applied and Environmental Microbiology p. 6516–6523 September 2012 Volume 78 Number 18

Comparative Assessment of Nitrogen Fixation Methodologies, Conducted in the Oligotrophic North Pacific Ocean

Samuel T. Wilson,^{a,b} Daniela Böttjer,^{a,b} Matthew J. Church,^{a,b} and David M. Karl^{a,b}

production were measurable for nonconcentrated seawater samples after an incubation period of 3 to 4 h. The $^{15}\text{N}_2$ tracer measurements compared the addition of $^{15}\text{N}_2$ as a gas bubble and dissolved as $^{15}\text{N}_2$ enriched seawater. On all sampling occasions and at all depths, a 2- to 6-fold increase in the rate of $^{15}\text{N}_2$ assimilation was measured when $^{15}\text{N}_2$ -enriched seawater was added to the seawater sample compared to the addition of $^{15}\text{N}_2$ as a gas bubble. In addition, we show that the $^{15}\text{N}_2$ -enriched seawater can be prepared prior to its use with no detectable loss (<1.7%) of dissolved $^{15}\text{N}_2$ during 4 weeks of storage, facilitating its use in the

Annual climatology of N₂ fixation at Station ALOHA



- Annual (2005-2012) average N₂ fixation (bubble):
48 mmol N m⁻² yr⁻¹ (0.3 mol C m⁻² yr⁻¹)
- Annual (2012-2013) average N₂ fixation (no-bubble):
143 mmol N m⁻² yr⁻¹ (0.9 mol C m⁻² yr⁻¹)

Summary



- NCP at ALOHA averages $\sim 2 \text{ mol C m}^{-2} \text{ yr}^{-1}$, with an uncertainty of $\sim \pm 50\%$.
- Uncertainties in NCP derive from poor constraint on both physical (lateral advection, vertical entrainment, air-sea exchange) and biological processes (DOC flux, sediment traps, vertical migrators).
- The processes supplying nutrients supporting NCP remain unclear.
- Time series programs (augmented by autonomous technologies) continue to improve our ability to constrain the magnitude and variability in carbon fluxes in the open sea.

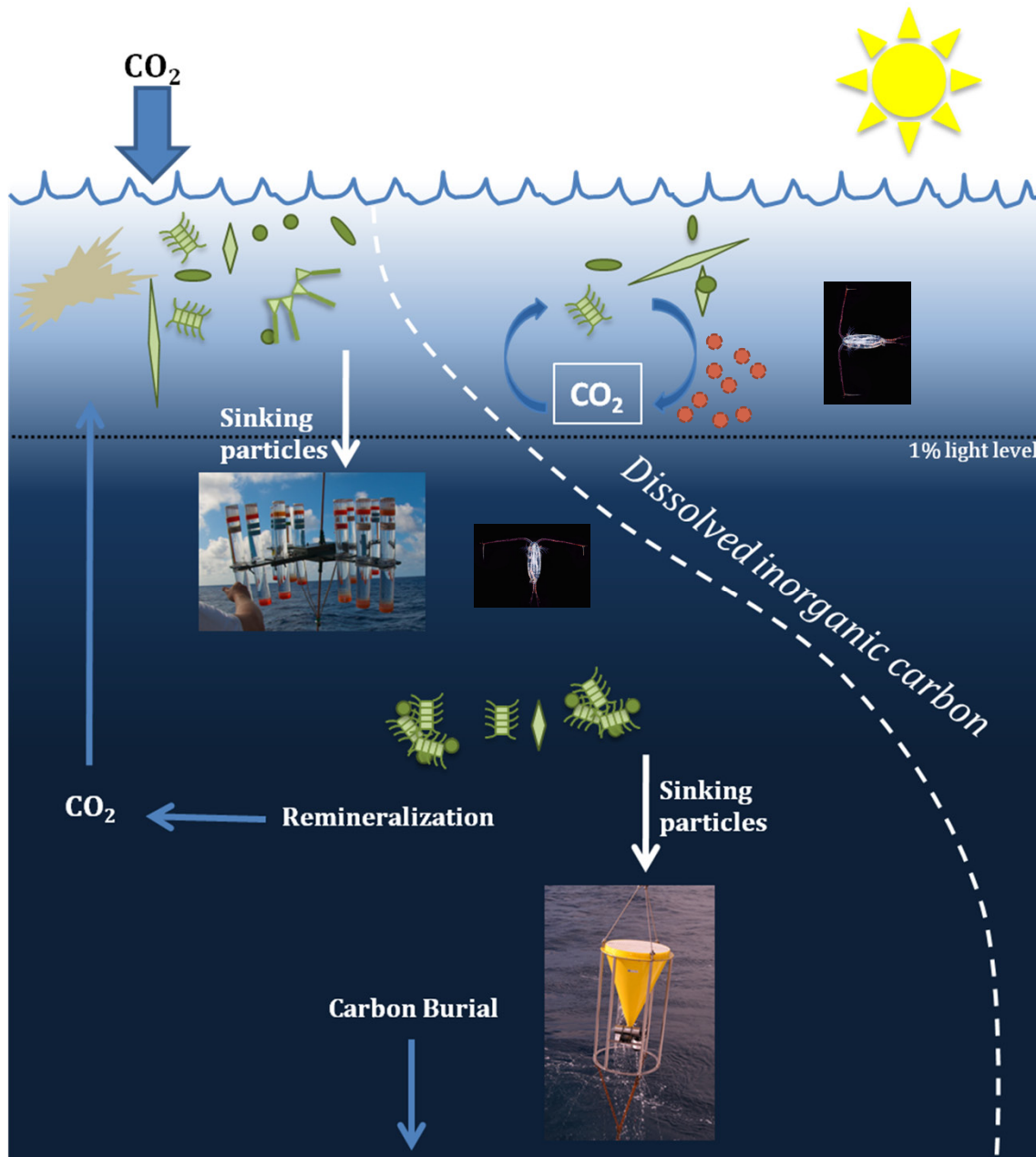


THANK YOU



EXTRA SLIDES

Station ALOHA is one of the few places on Earth where time series measurements enable mass balance constraint on ocean NCP

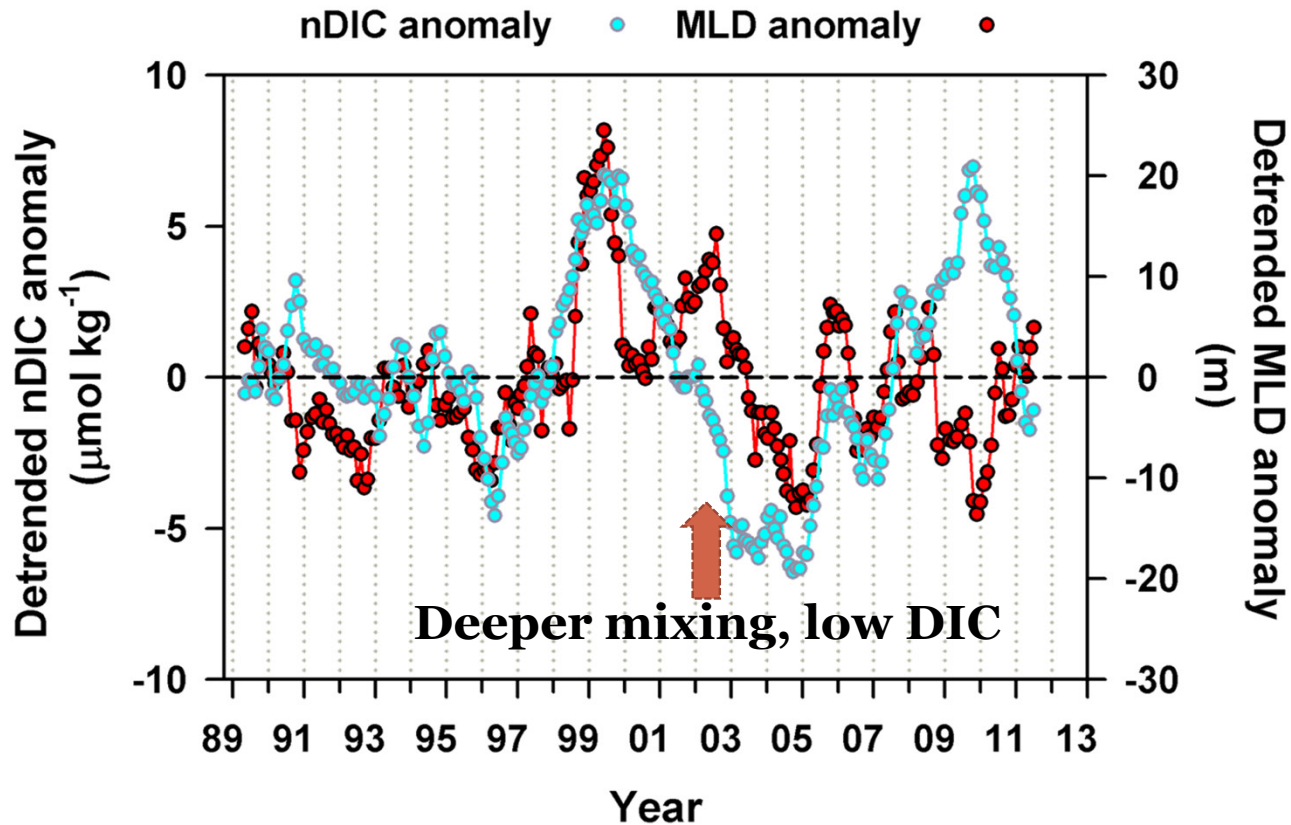


➤ How temporally variable are rates of production and export in the central North Pacific?

➤ What processes control this variability?

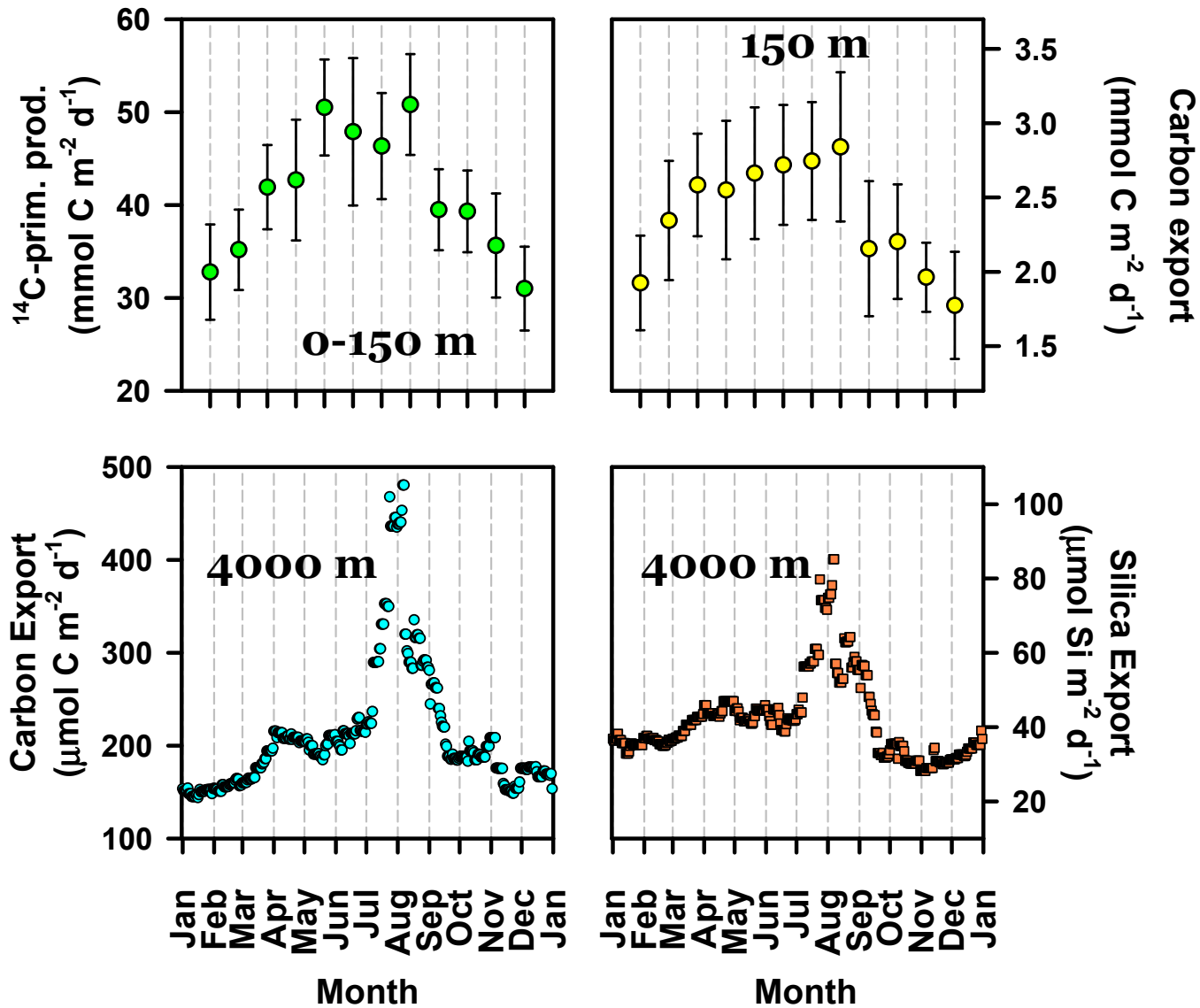
➤ What is the fate of biologically fixed carbon?

Interannual to subdecadal scale variability in mixing and mixed layer DIC



What role do biological and physical processes play in controlling the magnitude and variability in seasonal- to decadal-scale ocean carbon fluxes?

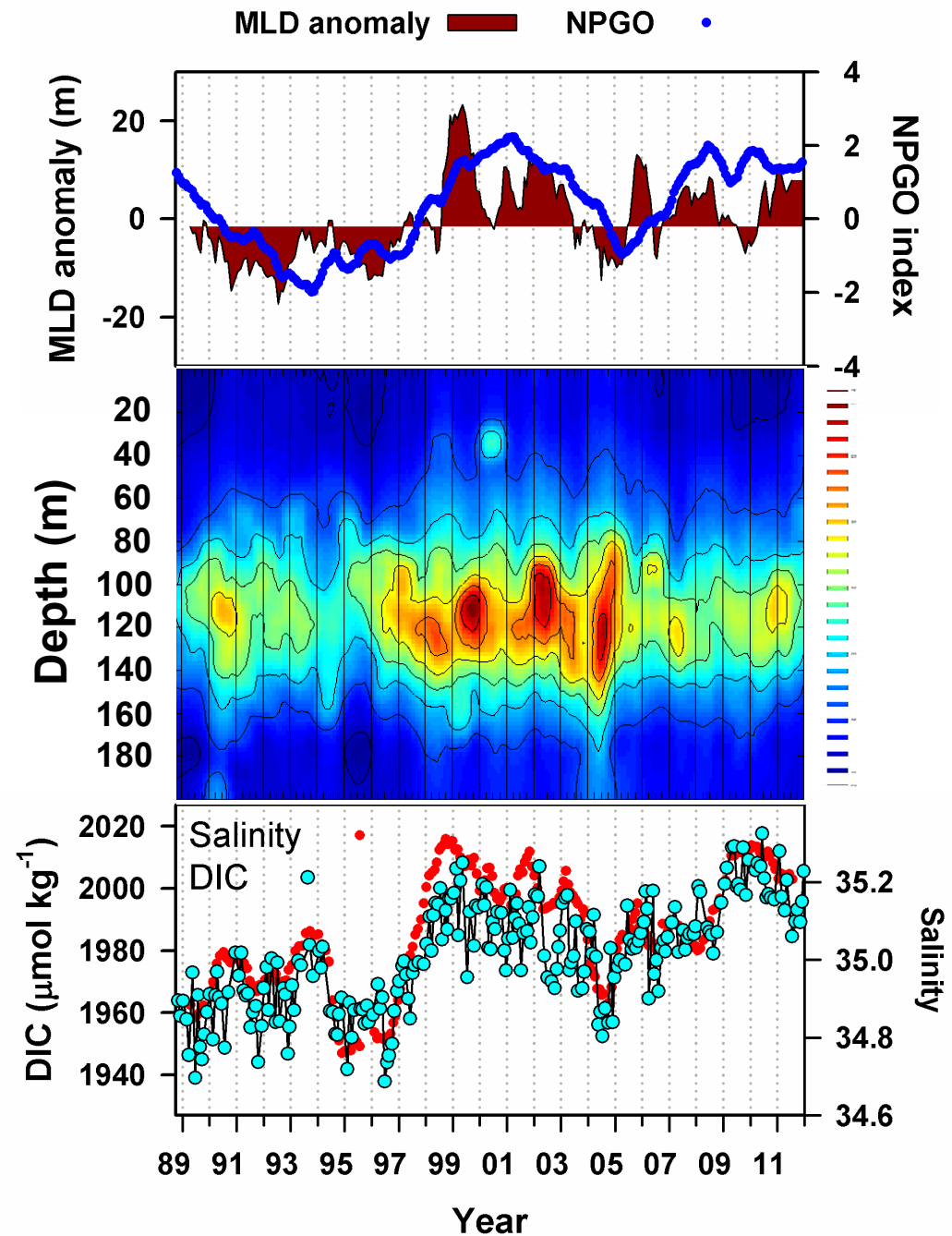
Annual cycle of productivity and export



Plankton community structure plays a key role in controlling carbon export

Climate modulated changes in the NPSG ecosystem

- Interannual to subdecadal variability in upper ocean mixing appears linked to basin-scale climate fluctuations.
- Changes in nutrient and light availability alters biological productivity and biomass.
- Variability in inorganic carbon inventories, ocean pH, and air-sea CO₂ flux appear correlated to variations in salinity.





- ^{14}C -primary production provides a highly sensitive means of quantifying daily rates of carbon fixation.
 - Requires confinement of samples
 - DO^{14}C ? At ALOHA ~20-30% of particulate carbon fixation
 - Daytime only or 24 hours? ~20-30% loss overnight
- ^{14}C -primary production \neq gross primary production
- ^{14}C -primary production \neq net community production
- Depending on how the method is employed, ^{14}C -primary production \approx net primary production?