Ocean Acidification and Biogeochemical Cycling



Uta Passow

Introduction

Facts About Ocean Acidification (OA)





- Burning of fossil-fuel lead to increase in *p*CO₂
- CO₂ dissolves in ocean water
- Decrease in pH, increase in DIC & pCO₂, TA remains constant

Introduction

In situ variability of carbonate system parameters

N Sea annual pH range Annual S. Ocean 56 (b) (a) 0.9 8.4 55 0.8 2200 0.7 54 8.2 DIC, µmol kg⁻¹ 2100 0.6 53 pH_{sws} 8.0 0.5 2000 52 0.4 7.8 1800 0.3 1600 51 1400 0.2 7.6 1200 50 0.1 JFMAMJJASONDJF DJFMAMJJASONDJF D Month, 1993-95 -2 8 0 2 (a) Blackford & Gilbert 2007 25 Chlorophyll a, mg m⁻³ W Coast USA 20 0.0 С 15 8.5 8.1 8.3 Depth (m) 10 8.1 7.9 7.8 5 7.8 7.9 7.7 0 J F 200 D J F М S 0 Ν D 7.5 pH 7.3 Month, 1993-95 D 0 2350 Gibson & Trull 1999 20.01 2250 Depth (m) 100 2150 Dissolved 2050 Inorganic 2220 1950 2240 200 Carbon

Feely et al. 2008

Ocean acidification work shop Woods Hole 3/22- 3/24 2011

(umol kg-1)

1850



(1) Primary Production & growth rate



Pelagic Ecosystem CO₂ Enrichment study (PeECE III 2005), but no difference PeECE I & II (2001 & 2003)

(1) Primary Production & growth rate



(2a) Nitrogen Fixation by Cyanobacteria

(2a) Nitrogen Fixation by Cyanobacteria

(2b) Nitrification

% Change in NO₃: DIN ratio (Δ 385 & 1000 ppm CO₂) based on Husemann et al. 2002.

(2b) Nitrification

(2c) Denitrification

Δ dissolved of dissolved O₂ between 2080-2100 and 1980-2000

(2c) Denitrification

Stratification increases & shallower remineralization

- Δ di Suboxic zones in the ocean increase •
 - **Denitrification quadruples in next 2000 years**
 - Less nitrate available, decrease in biological production
 - Positive feed-back

28.

27.

Synthesis

(2) Nitrogen Cycle

Hutchins

Decrease in fraction of inorganic N that is nitrate

(3a) CaCO₃: Saturation Horizon

 $Ca^{2+} + 2HCO_3^{-} \leftrightarrow CaCO_3 + (CO_2) + H_2O$ Simulated aragonite saturation at the surface

- Aragonite and Calcite saturation is decreasing rapidly.
- Saturation is spatially variable with low values in the high latitudes

(3b) (Bio-)Calcification: Coccolithophores & Foraminifera

Irie et al 2010: Optimality-model predicts that natural selection will favors heavy calcification if calcification is a defensive strategy and photosynthesis is enhanced.

(3b) (Bio-)Calcification: Coccolithophores & Foraminifera

(3c) CaCO₃ Dissolution

Gehlen & Bopp 2008

(3c) CaCO₃ Dissolution

- Decrease in calcification & increase in production of POM potentially results in an increase in rain ratio (POC: PIC) in exported material
- Increased rain ratio → promotes dissolution, timescales of 1000-10,000 years).
- However, less ballasting might reduce sedimentation of POC not leading to an increased rain ratio
- Carbonate compensation: Dissolution of CaCO₃ in sediments have a buffer capacity for the deep ocean on geological time scales. Negative feed back!

40°E

60*W

Gehlen & Bopp 2008

100°E

(4) DMS, DMSP Production

(4) DMS, DMSP Production

(5a) Phytoplankton Species Composition

(5a) Phytoplankton Species Composition

(5b) Genetic Acclimation & Adaptation

Adaptation: (genetic change) 200 - 1000 generations

(5b) Genetic Acclimation & Adaptation

Stress response: immediate

- Green algae Chlamydomonas 1000 generations. Some lines did not grow in lower pCO₂ conditions - neutral mutation of CCM making it less efficient - Collins & Bell (2004):
- Similar experiments underway
- Indications from paleo-record of coccolithophores suggest some adaptation
- Evolutionary models
- Impact unknown

Synthesis

Bottom - up Controls of Primary Producers

Synergistic & Antagonistic Interactions between CO₂, temperature, light, nutrients, trace metals

Wyatt et al 2010: surface layer $NH_4 \leftrightarrow NH_3$

Synthesis

Bottom - up Controls of Primary Producers

Synergistic & Antagonistic Interactions between CO₂, temperature, light, nutrients, trace metals

Wyatt et al 2010: surface layer $NH_4 \leftrightarrow NH_3$

(6) Organic Matter & Trace element cycling

(6) Organic Matter & Trace element cycling

umol quanta m ² s ¹

(7) Stoichiometry of organic matter

(7) Stoichiometry of organic matter

(8a) Microbial Loop: DOM Production

(8a) Microbial Loop: DOM Production

(8b) Microbial Loop: Bacterial Concentration, Turnover and Diversity

Ocean acidification work shop Woods Hole 3/22- 3/24 2011

(8b) Microbial Loop: Bacterial Concentration, Turnover and Diversity

(8c) Microbial Loop: Bacterial Degradation of DOM

(8c) Microbial Loop: Bacterial Degradation of DOM

Synthesis

(8) Concept: Microbial Carbon Pump

Jiao et al. Nature Reviews Microbiology 2010

Synthesis

(9) Carbon Flux: Partitioning

(9) Aggregation and Transparent Exopolymer Particles, TEP

(9) Aggregation and Transparent Exopolymer Particles, TEP

- No change in abiotic TEP formation expected
- Possibly enhanced TEP formation due to increase in DOC exudation but not observed in PeECE II & III.
- Impact for aggregation or microbial dynamics are unclear

(9) Sinking velocity of aggregates

Biermann & Engel 2010

(9) Sinking velocity of aggregates

 Lack of ballasting due reduction in calcification reduces sinking velocity

 Increase in the volume fraction of TEP in aggregates decreases sinking velocity

• Decrease of flux to deep sea in the future ocean \rightarrow positive feed back

CO2 treatment

Biermann & Engel 2010

(9) Efficiency of Biological pump

•Decrease in diatom abundance

(10) Zooplankton

Protozooplankton day 14

Effects on grazing rate & vertical migration

(10) Zooplankton

T14 Ambient

0.6 -

ns

Protozooplankton day 14

T14 High pCO₂

Growth rate of amphipod

- Efffects on microzooplankton appear mostly indirect (shift in food source)
- Effects on mesozooplankton sublte, in physiology (gene expression) but not in mortality or growth
- Consequences for biogeochemical cycling impossible to predict at this time

Effects on grazing rate & vertical migration

Synthesis

Effects on Biogeochemistry on Biological Timescales

Summary

- Number of available studies very limited considering the complexity of the problem
- Ocean acidification seems to significantly alter the environment of microbes
- Changes in phytoplankton and bacterial community structure and physiology are likely
- Changes in microbial processes and community structure impact biogeochemical cycles
- Magnitude and direction of impact presently unpredictable, because of its complexity
- ➤ Associated changes of climate add to the complexity → need for multifactorial experiments

Challenges

- 1. The carbonate system structures the environment.
 - a) Spatial and temporal fluctuations are large and *will become larger as OA progresses*
 - b) Feed-backs make changes in carbonate system partially unpredictable
 - c) These effect organisms, populations, ecosystems and biogeochemical cycling
 - d) Till recently ignored: Re-evaluate many "pre-OA" results
- 2. OA does not happen isolated synergistic and antagonistic interactions with other parameters of major importance.
- 3. Ability of organisms to acclimatize and adapt unknown.
- 4. Empirical relationships (Redfield ratio, Martin curve) may not hold.

Conclusions

Approaches

1. Autonomous CO₂ buoys – understand rate of change an variability

2. Process studies (laboratory or shipboard experiments) isolating individual processes and field surveys evaluating temporal or spatial gradients will give us information on the breadth of possible reactions

3. Explore combined influence of global change relevant parameters (light, nutrients, trace elements, temperature) in experiments. Test for parameter range in expected changes of specific regions.

4. Models could help predict regional specific expected changes of these parameter combinations (tighter collaboration between modelers and experimentalists).

5. Proxy evaluation to utilize past data for analysis of long-term trends.

6. "Evolutionary" Experiments

7. Mesocosm studies look at community level net effects (black box for individual processes)

8. Model processes and feed-backs, e.g. sedimentation – ballasting,

