

# Ocean Acidification—Carbonate Dissolution Kinetics: Lab vs. Field

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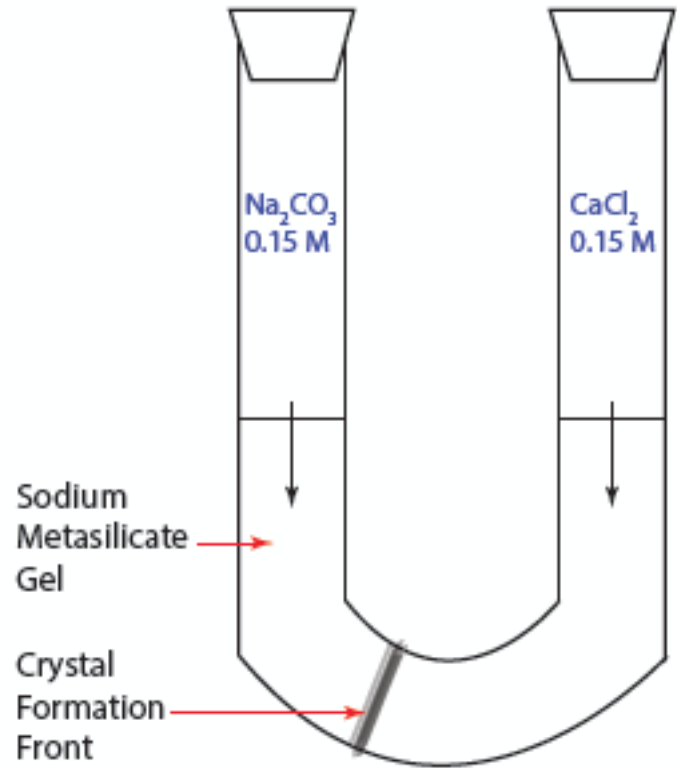
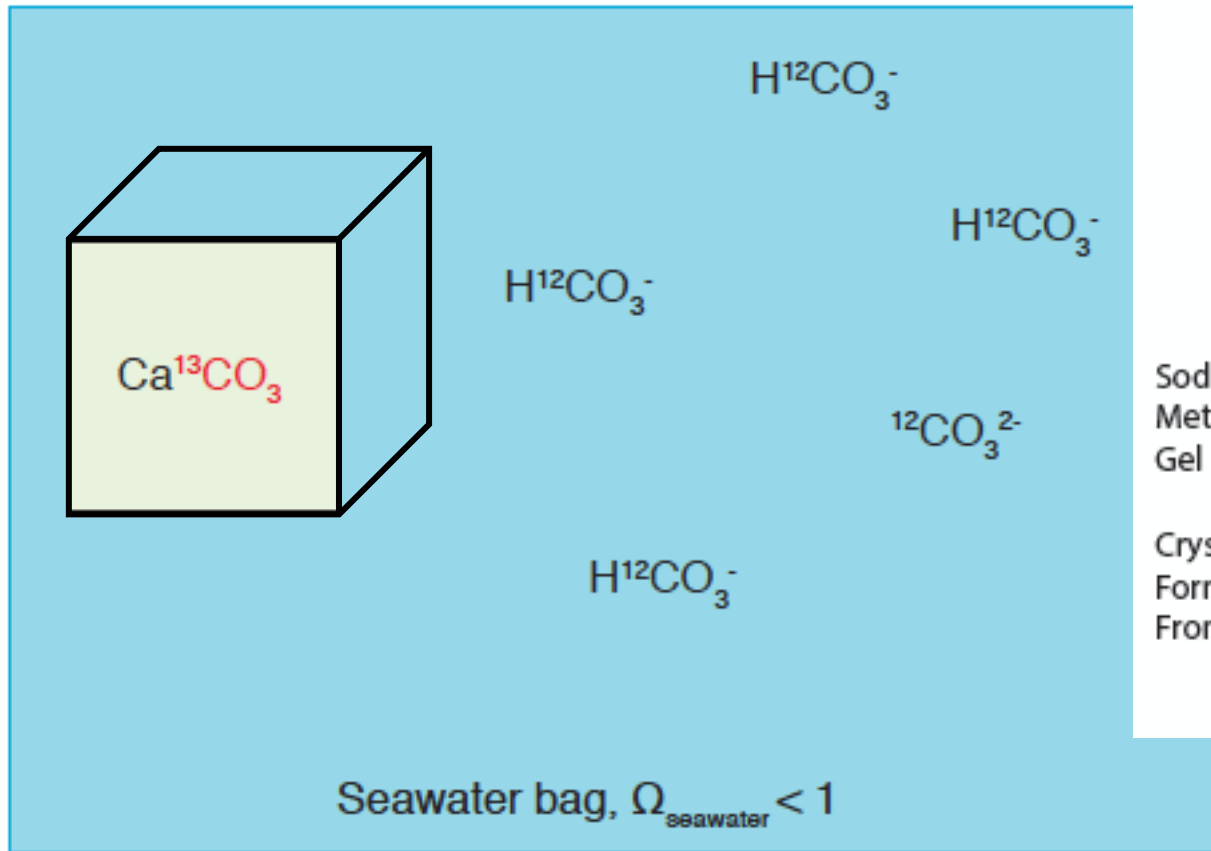
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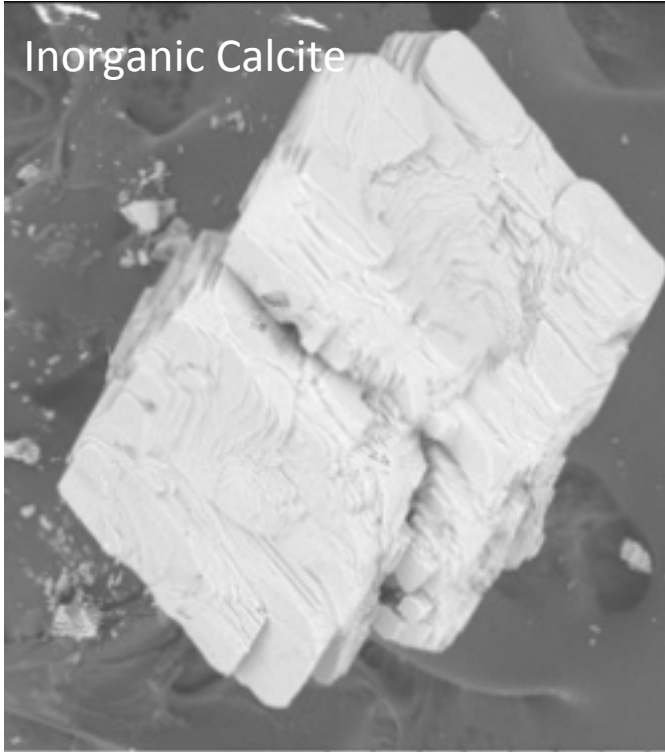
# Methodology



$$\text{DIC} = [\text{H}_2\text{CO}_3] + [\text{HCO}_3^-] + [\text{CO}_3^{2-}]$$
$$\text{Alk} \approx [\text{HCO}_3^-] + 2[\text{CO}_3^{2-}]$$

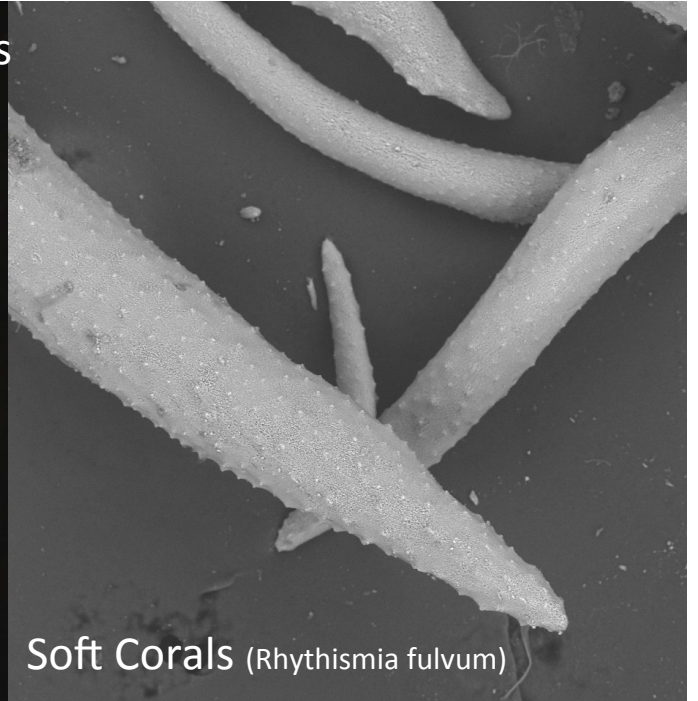
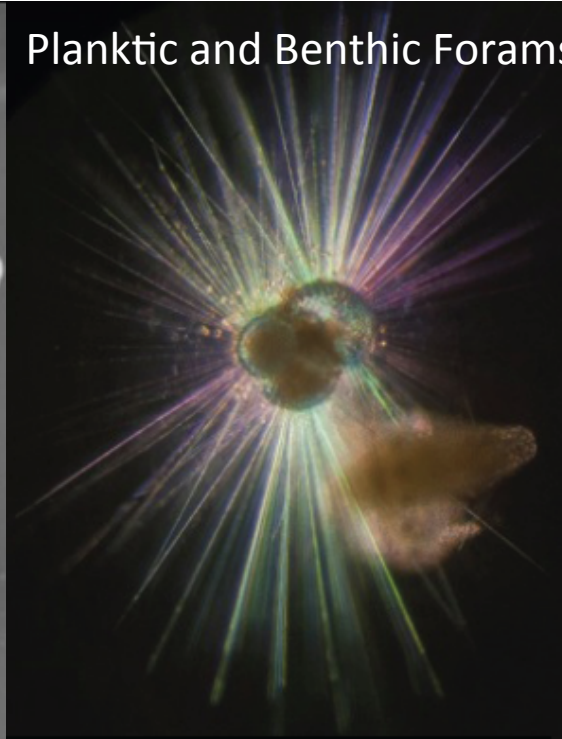
Isotope labeled calcite:  
Store-bought (Aldrich) and home-made calcite

Inorganic Calcite

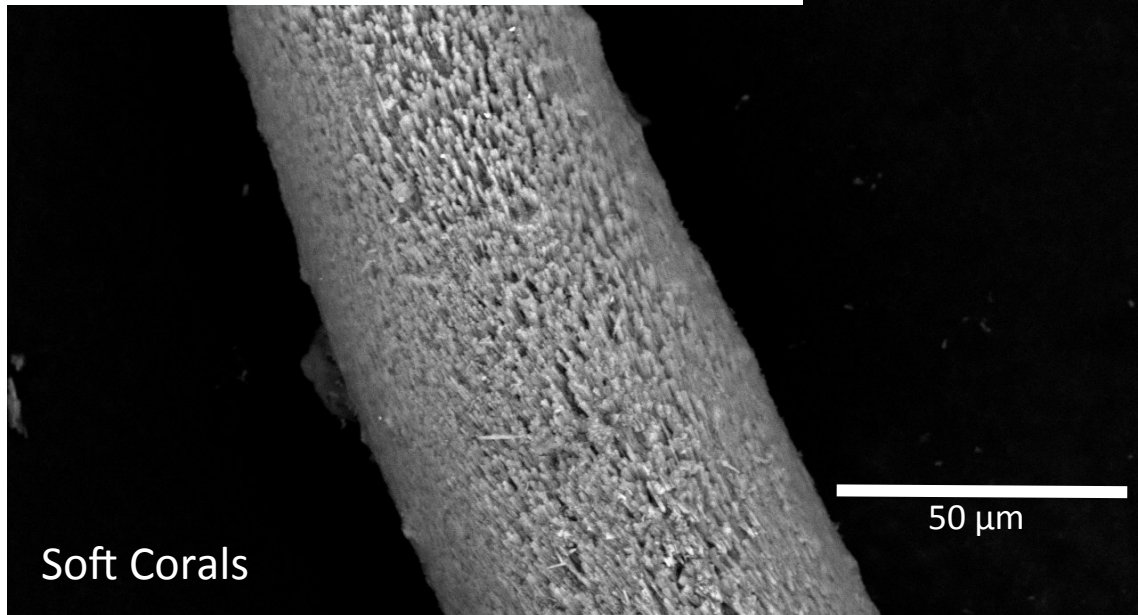


2014/03/03 14:40 D3.1 x180 500 um

Planktic and Benthic Forams



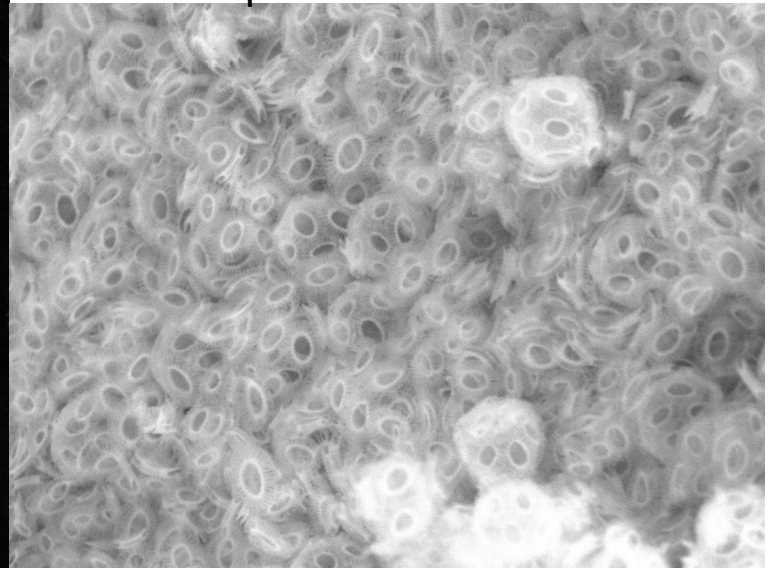
Soft Corals (*Rhythismia fulvum*)



Soft Corals

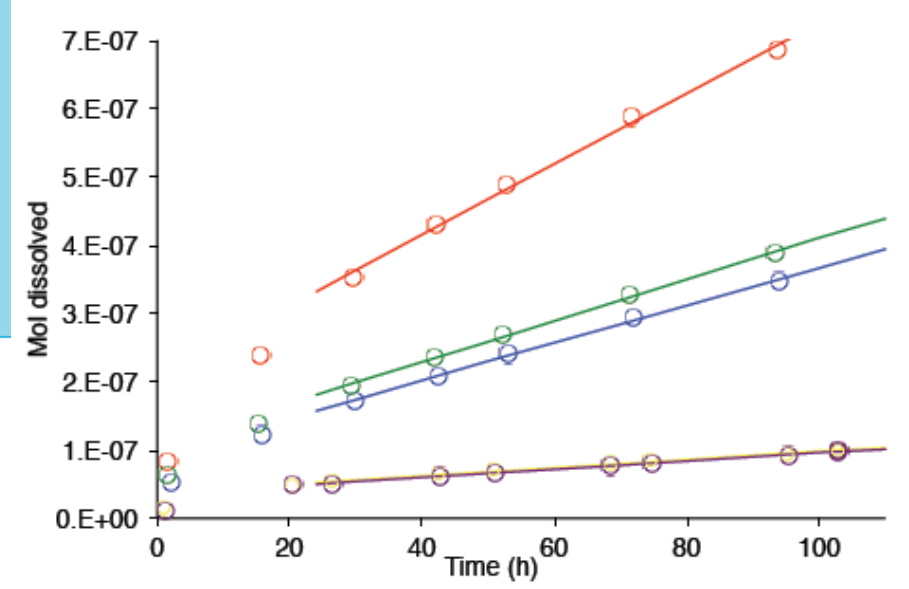
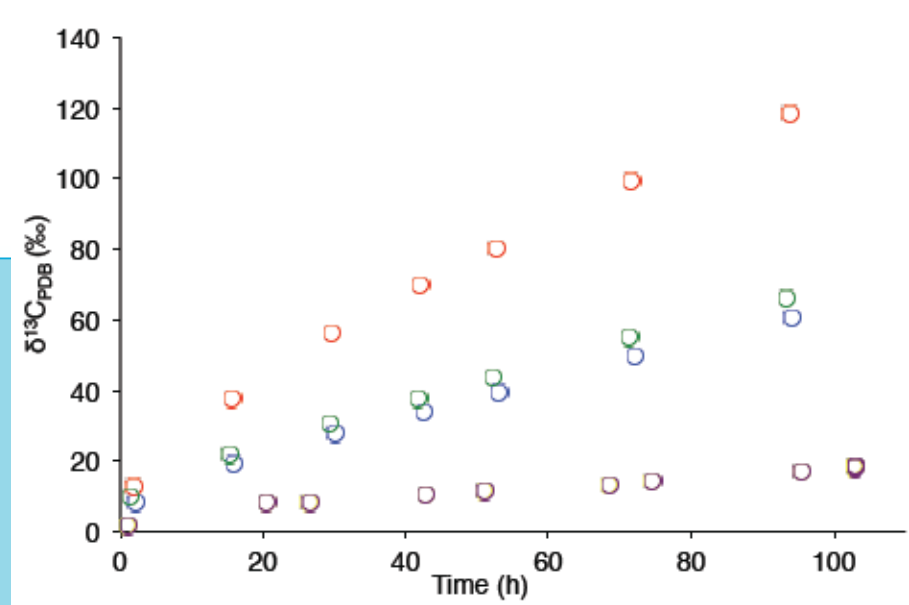
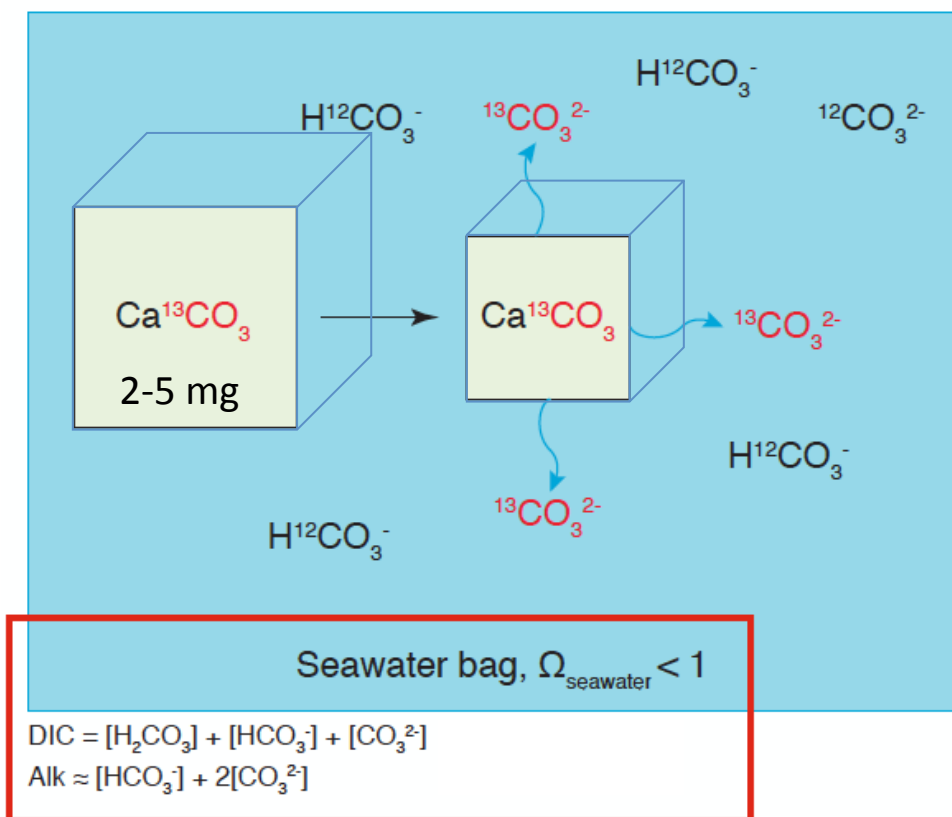
50 μm

Coccolithophores



Coccoliths

2015/03/17 11:11 D4.4 x5.0k 20 um



## Some Analytical Details:

TCO<sub>2</sub> is measured on 3-7 ml of sea water, Alk on 16-20 ml. TCO<sub>2</sub> prep per sample, 2 minutes, run per sample 10 minutes (TCO<sub>2</sub> analyzed on Picarro CRDS).

Alk prep per sample, 2 minutes, run per sample 15 minutes.

Del <sup>13</sup>C measured on TCO<sub>2</sub> is analyzed at same time as TCO<sub>2</sub> (Picarro CRDS).

To date, have run >2000 TCO<sub>2</sub>'s, del<sup>13</sup>C's and >500 Alks.

Dickson water is our TCO<sub>2</sub> and Alk Standard. We also make sub-standards from ocean water.

**Continued:**

Solid  $\text{CaCO}_3$  calibrated to VPDB is used as a  $\delta^{13}\text{C}$  standard.

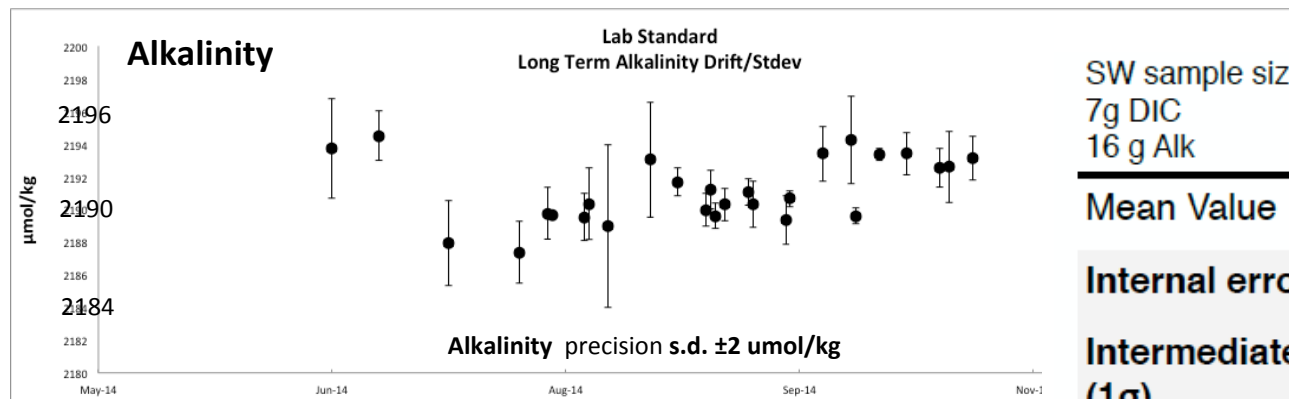
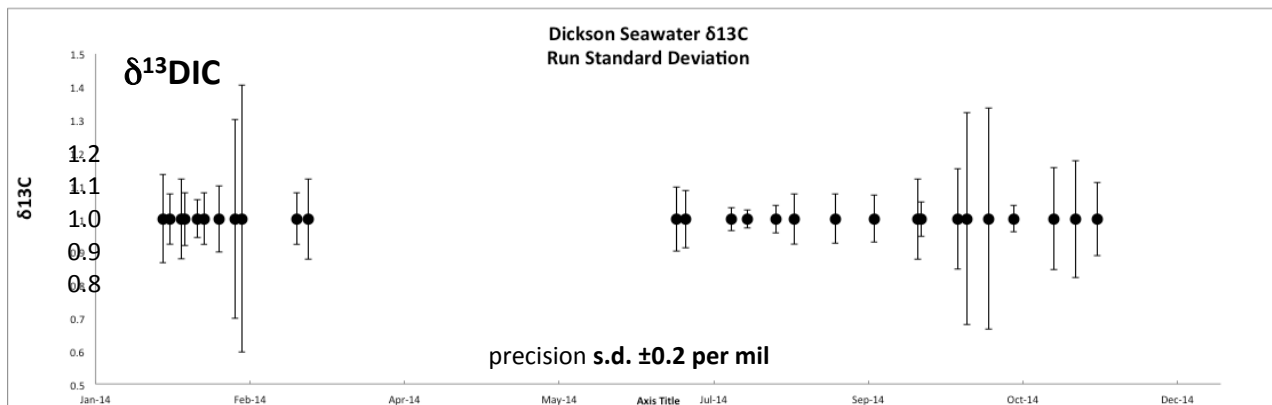
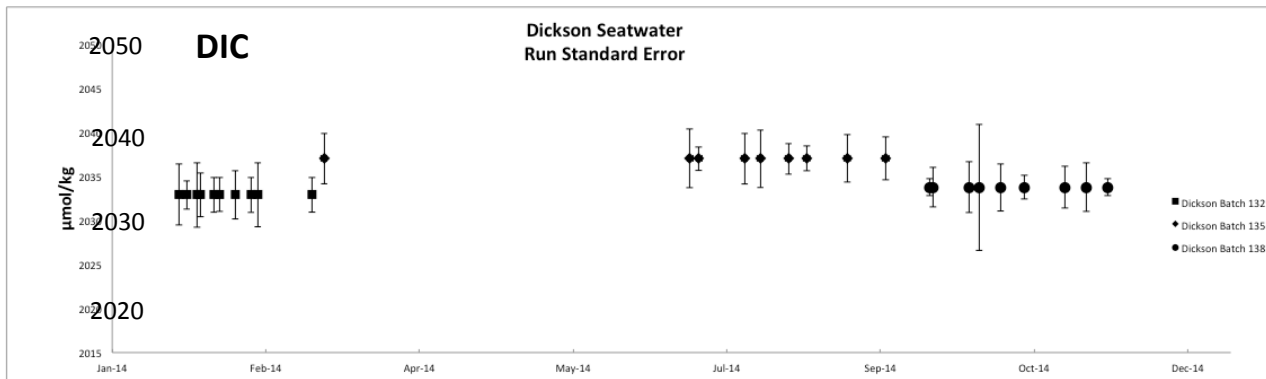
Picarro has been cross-calibrated with Mass Spec at Caltech.

Picarro has been adapted with mass flow control valve, water vapor trap and other personalized data management to improve accuracy and precision

Experiments generally involve 300 ml water in Supelco bags, on shaker table, effect of shaking has been calibrated

We use bags made from interlayered Al foil and polyethylene (Supelco) for experiments.

We have found  $\text{TCO}_2$  and Alk blanks ABOUND, gas loss, gas gain, acid leaching, base leaching!!



Average errors in  $\Omega$  are  $\sim 0.01-0.02$  units

SW sample size:  
 7g DIC  
 16 g Alk

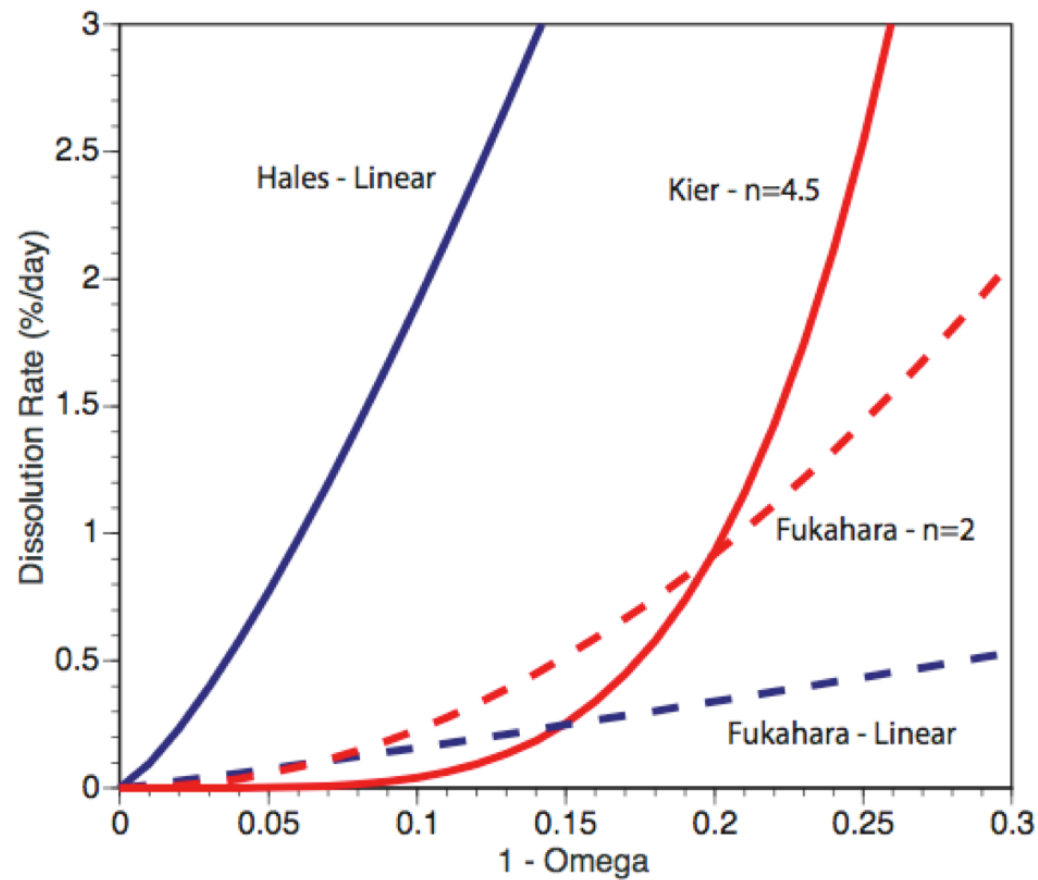
	<b>Alkalinity</b> ( $\mu\text{eq/kg}$ )	<b>DIC</b> ( $\mu\text{mol/kg}$ )
Mean Value	$\sim 2100$	$\sim 2000$
Internal error	2.5	5.1
Intermediate error (1 $\sigma$ )	2.0	5.6
External error (1 $\sigma$ )	2.5	5.2
MSWD	0.7	1.0

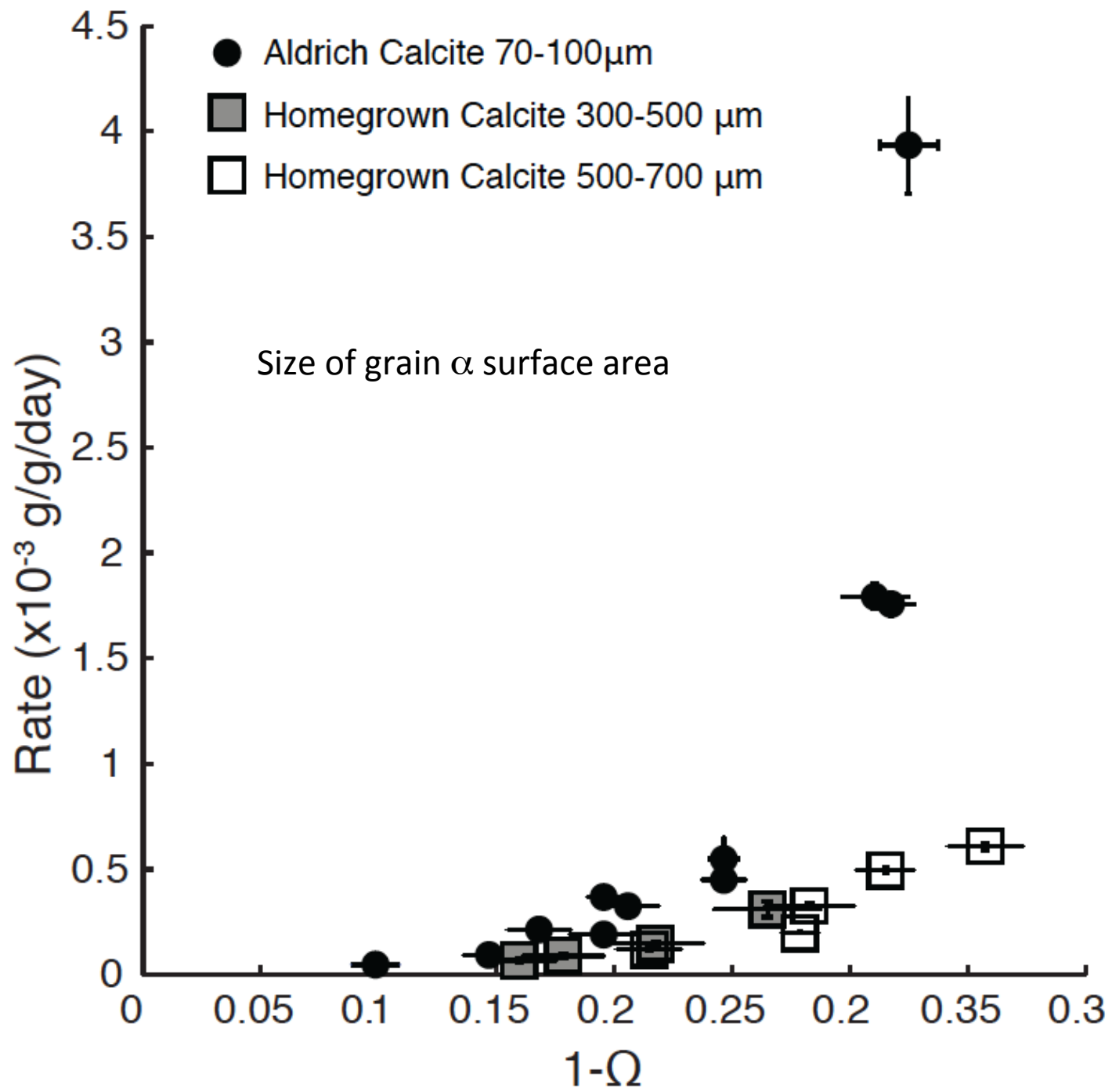
Using 100% labeled materials,  
**A 20‰ signal in  $\delta^{13}\text{C}$  changes alkalinity by 1  $\mu\text{eq/kg}$ !**

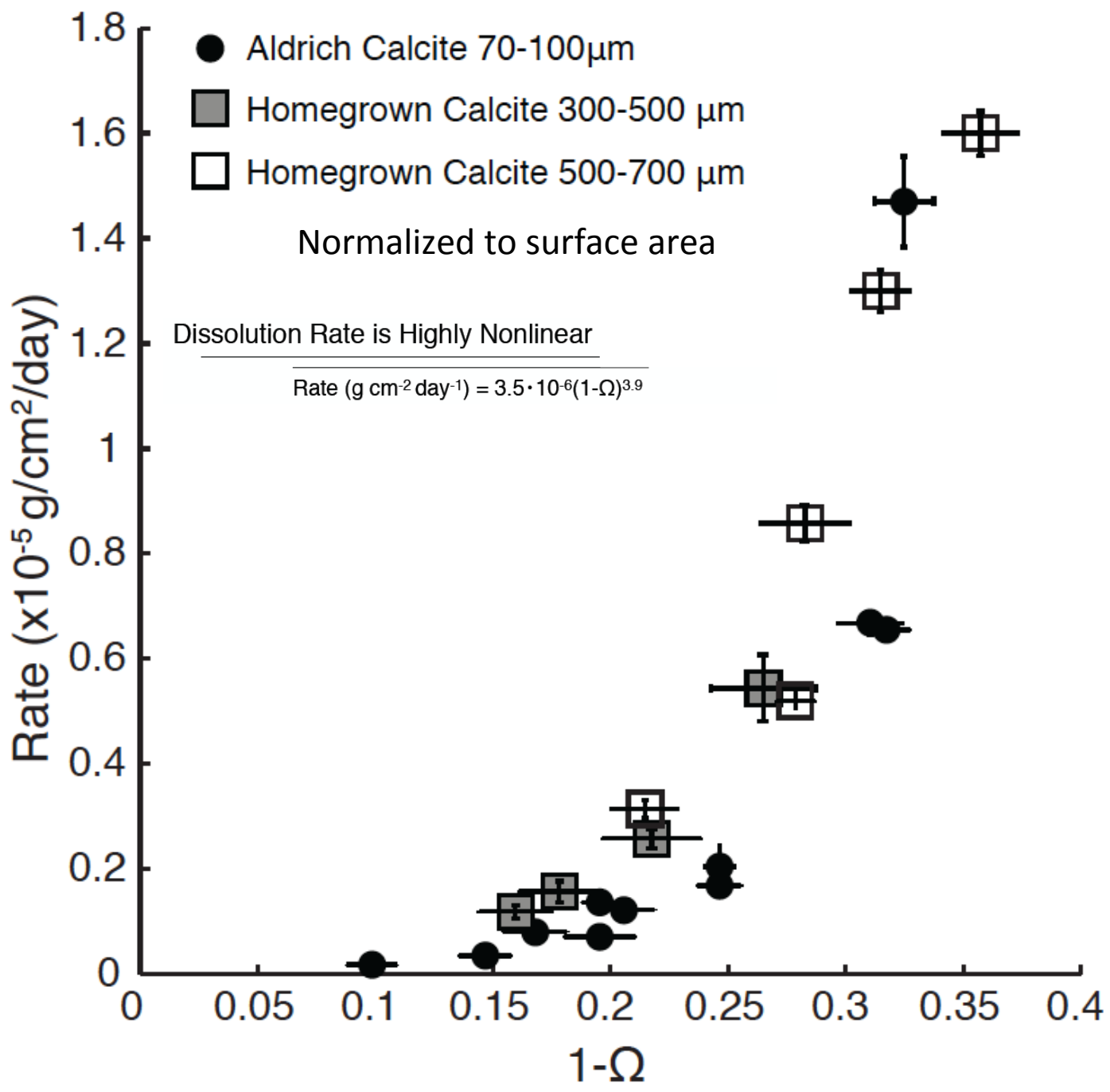
**Average errors in  $\Omega$  are ~0.01-0.02 units**

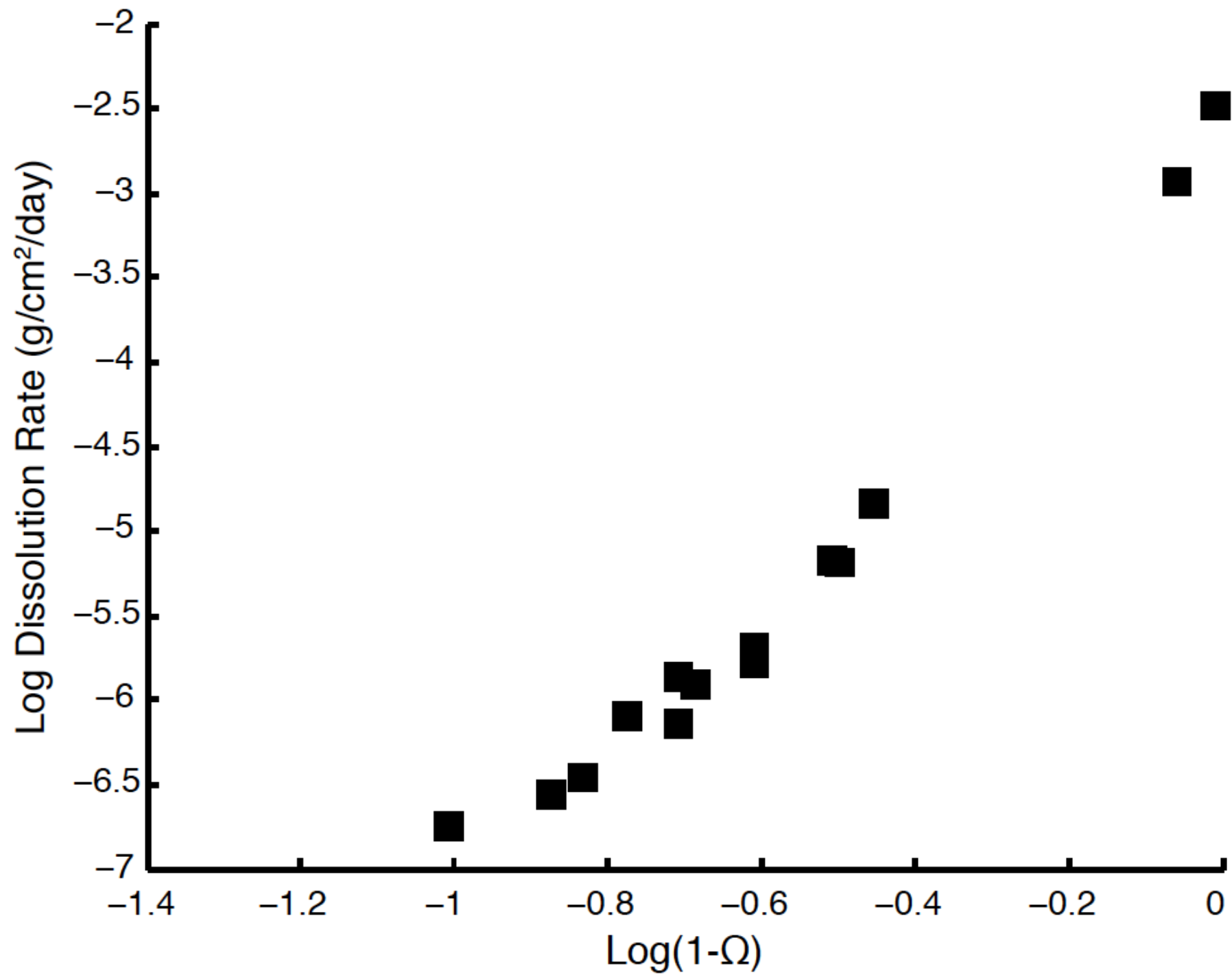


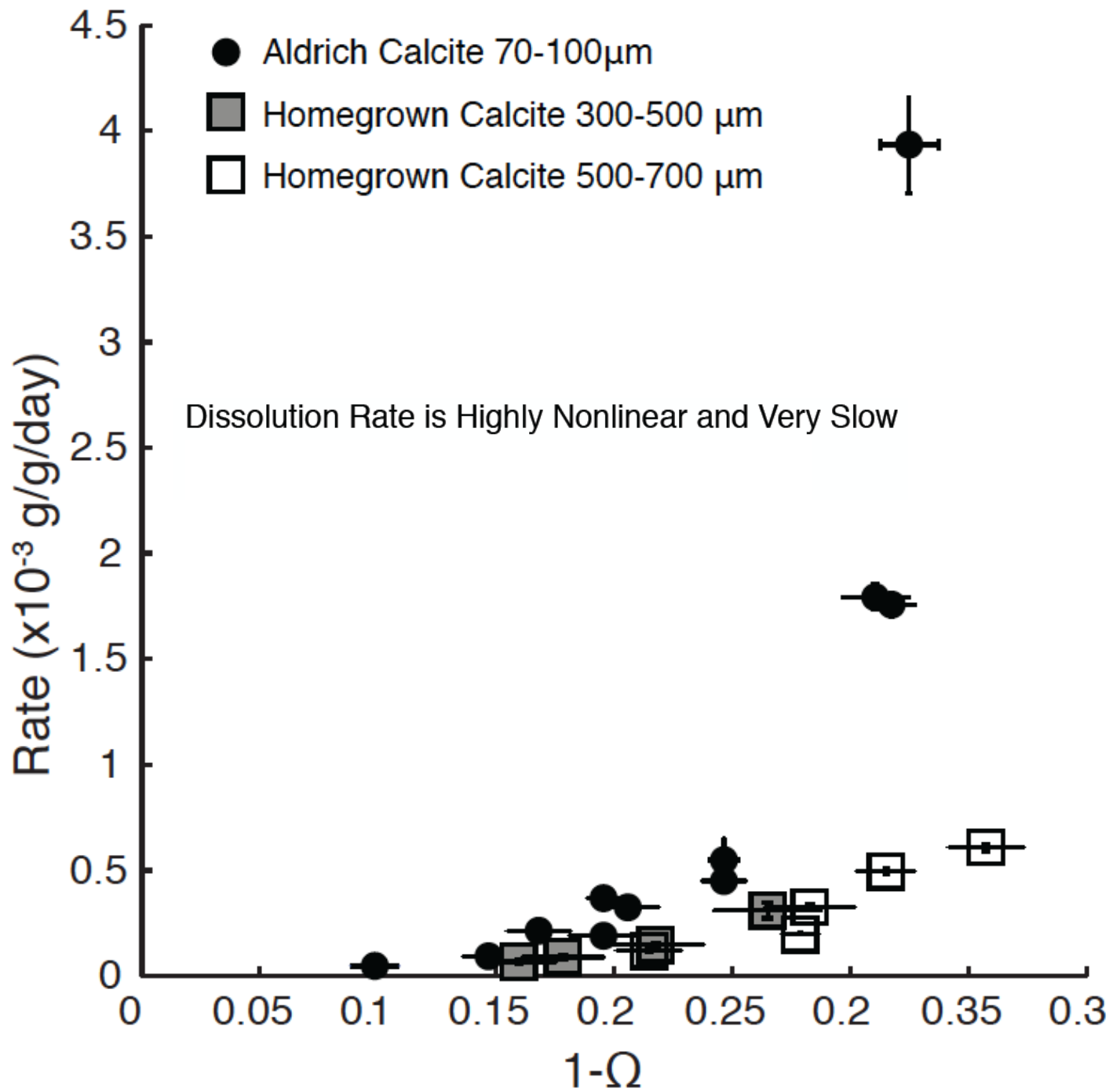
What is the correct formulation of dissolution as function of Omega???

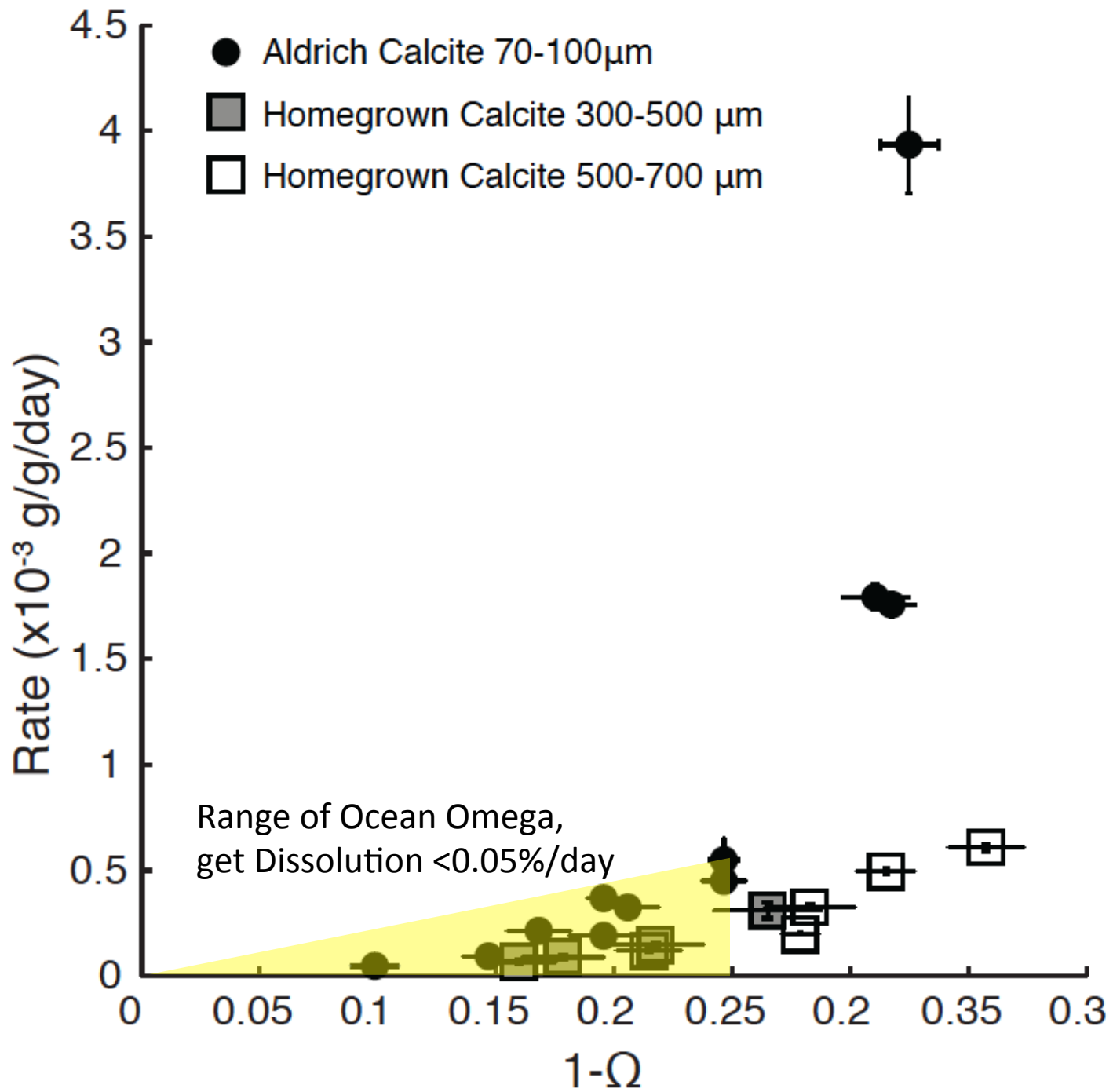




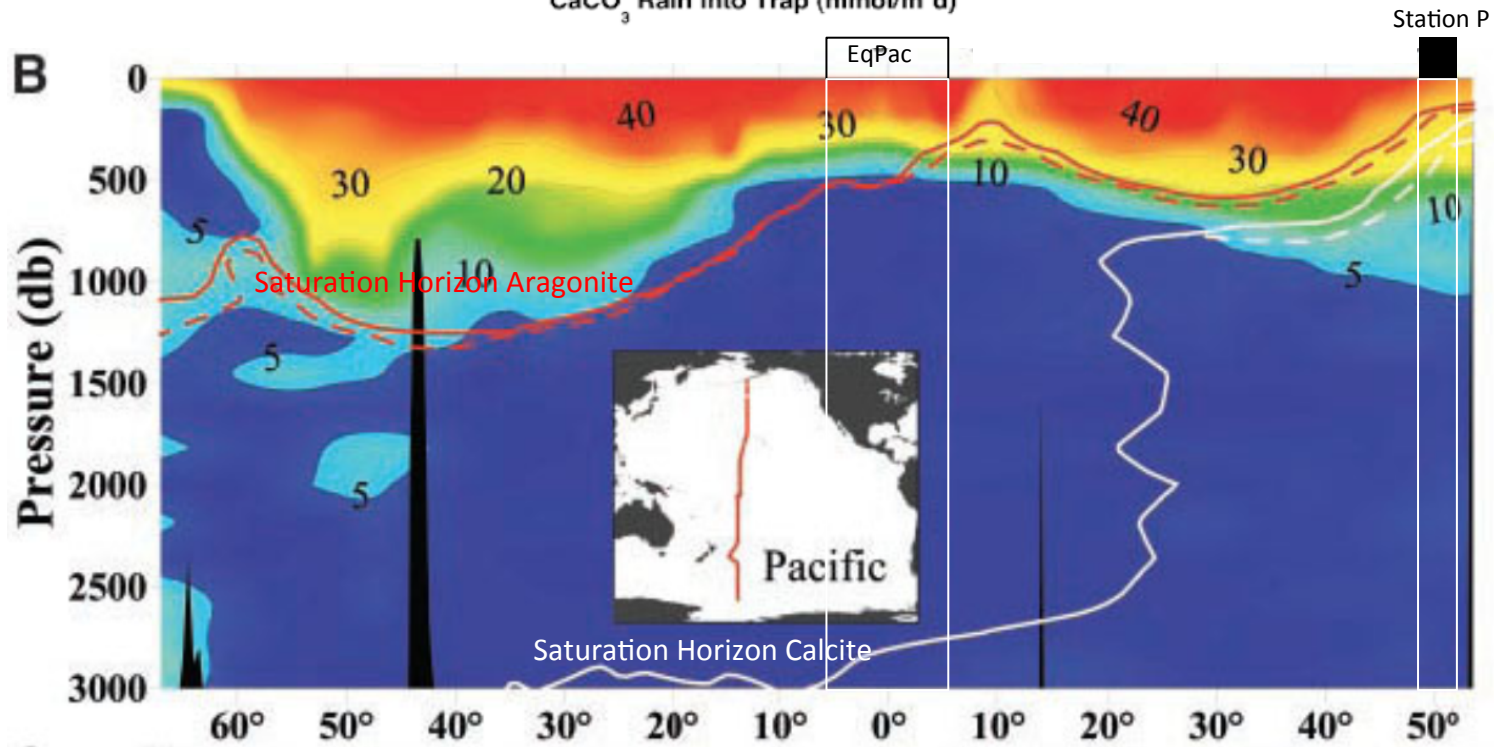
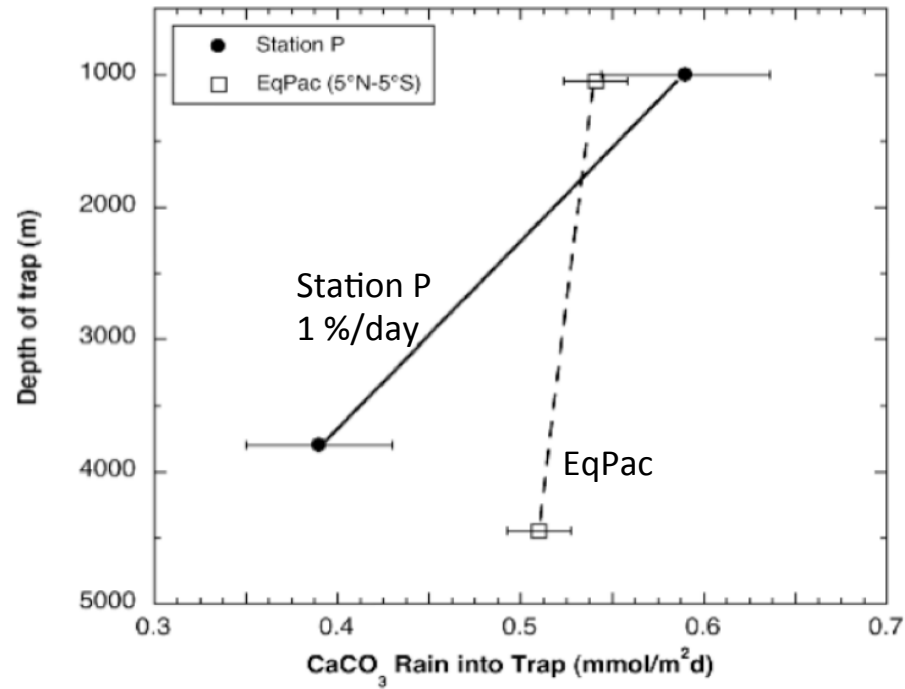


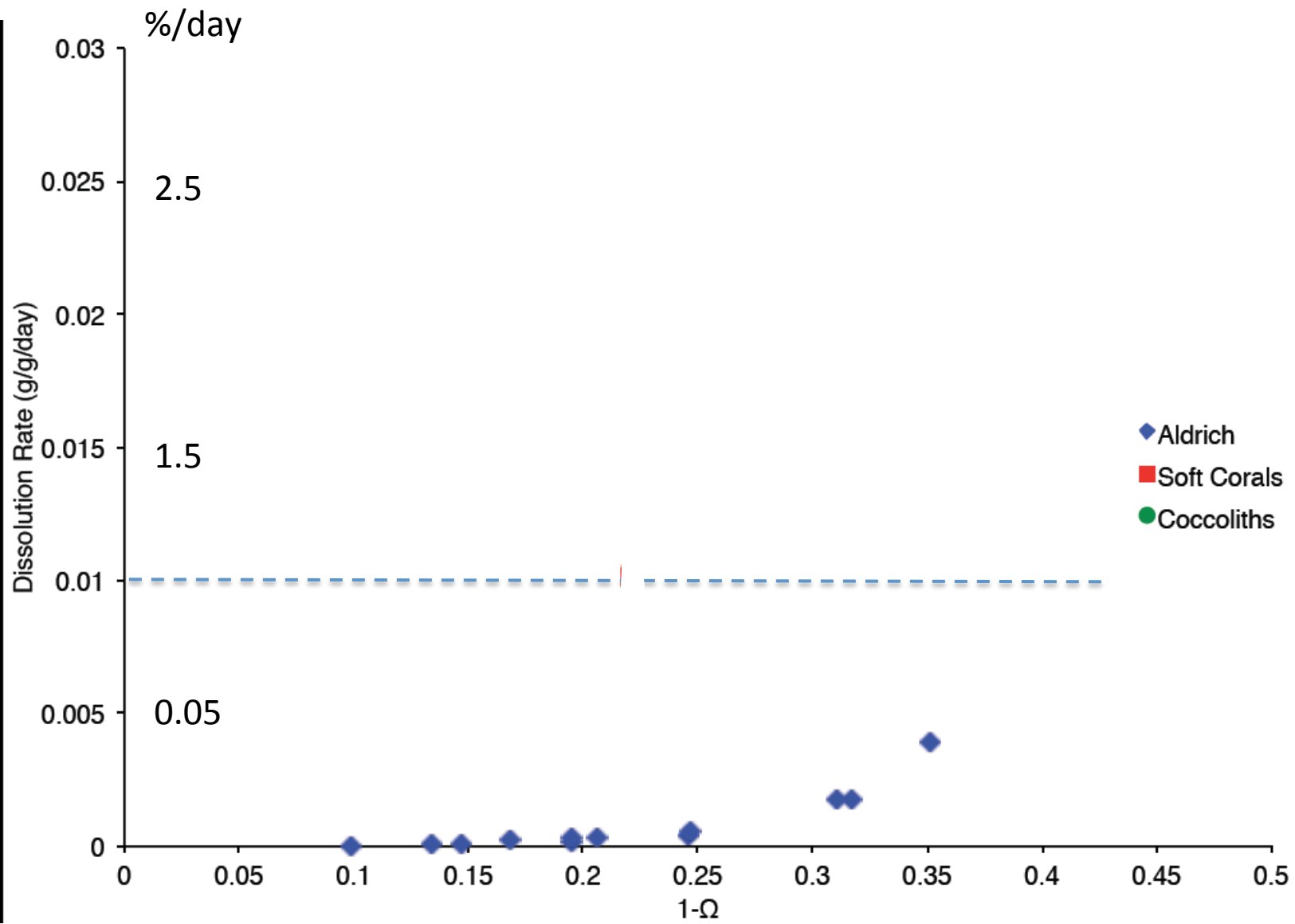




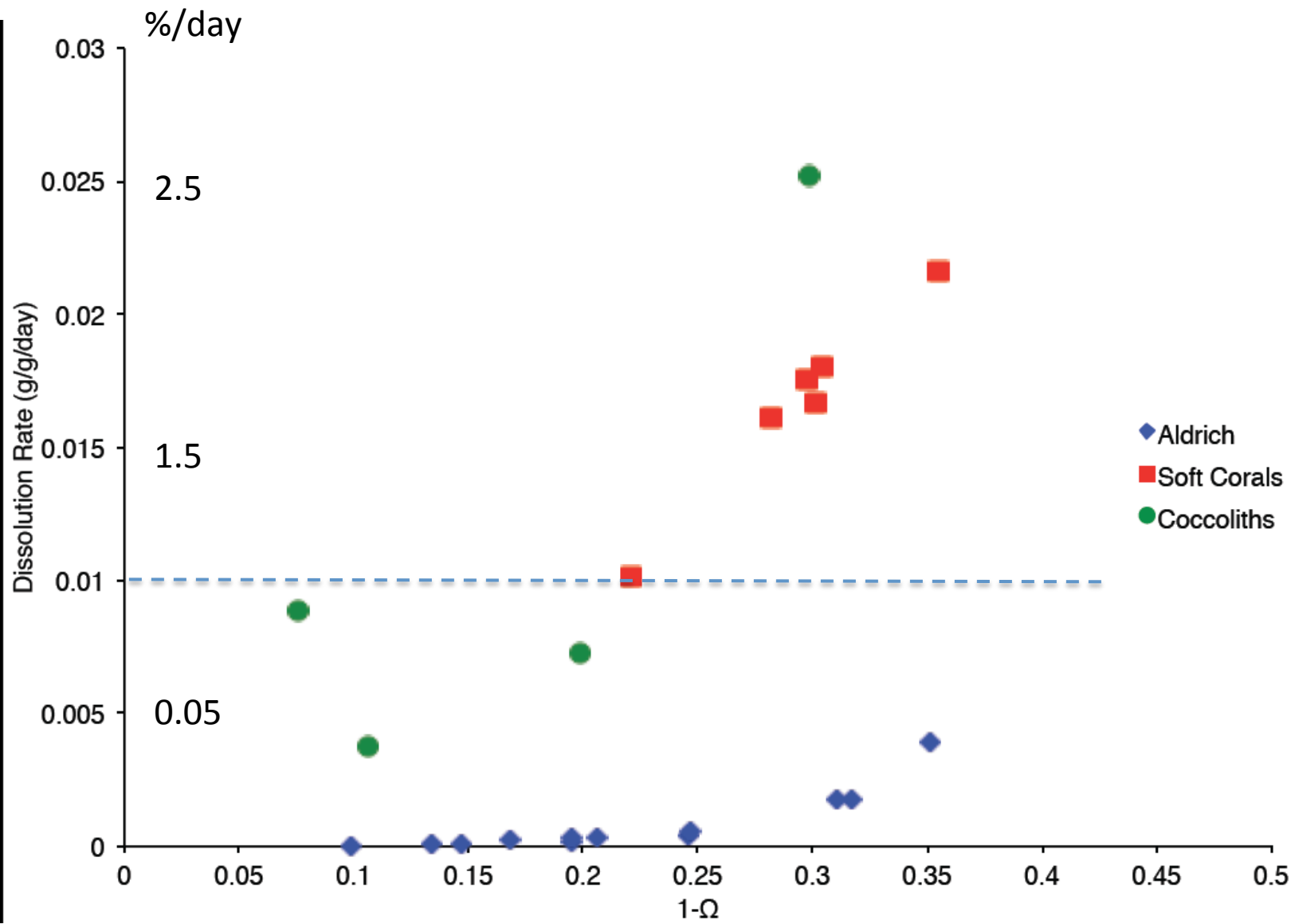


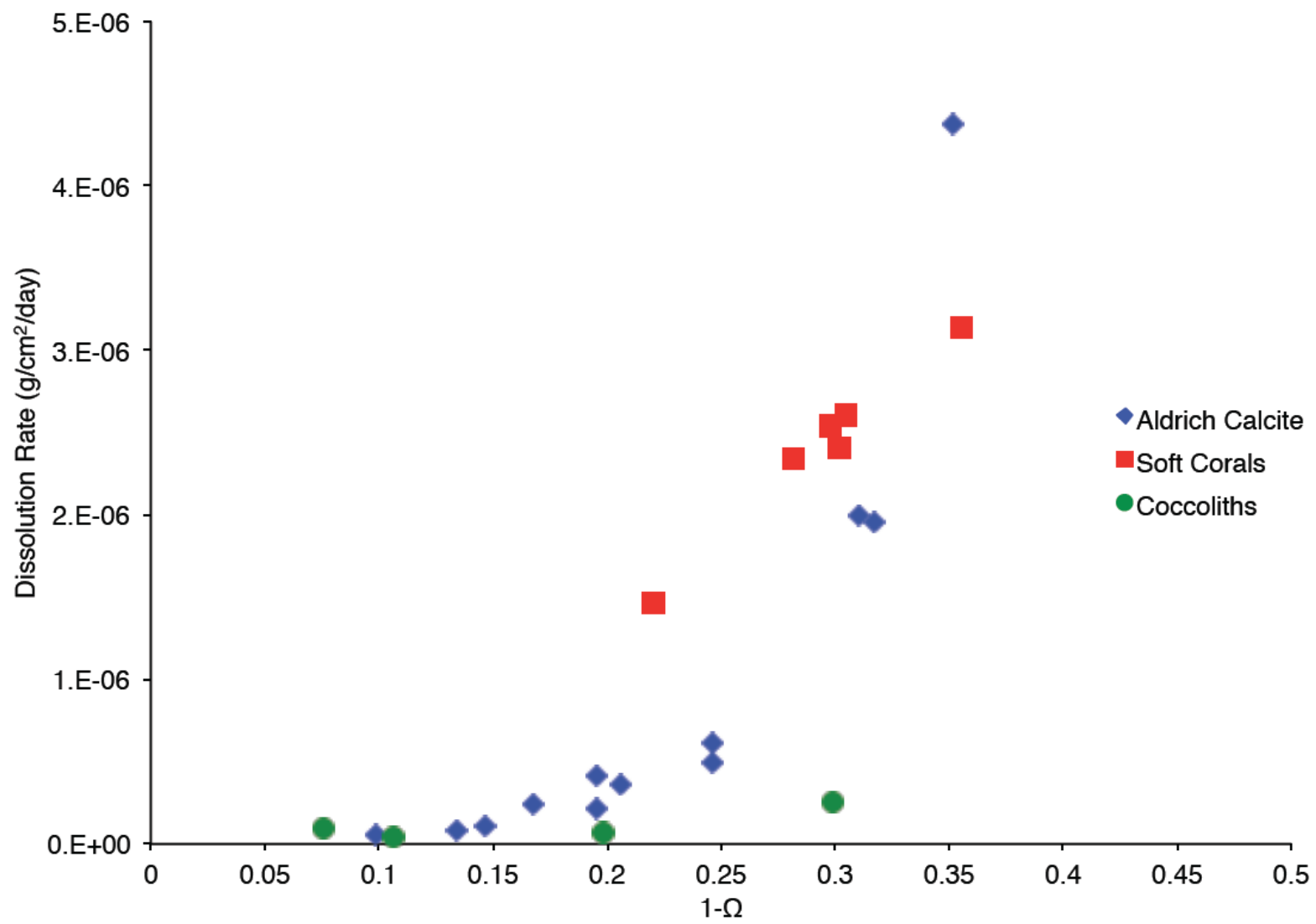
# Field Data







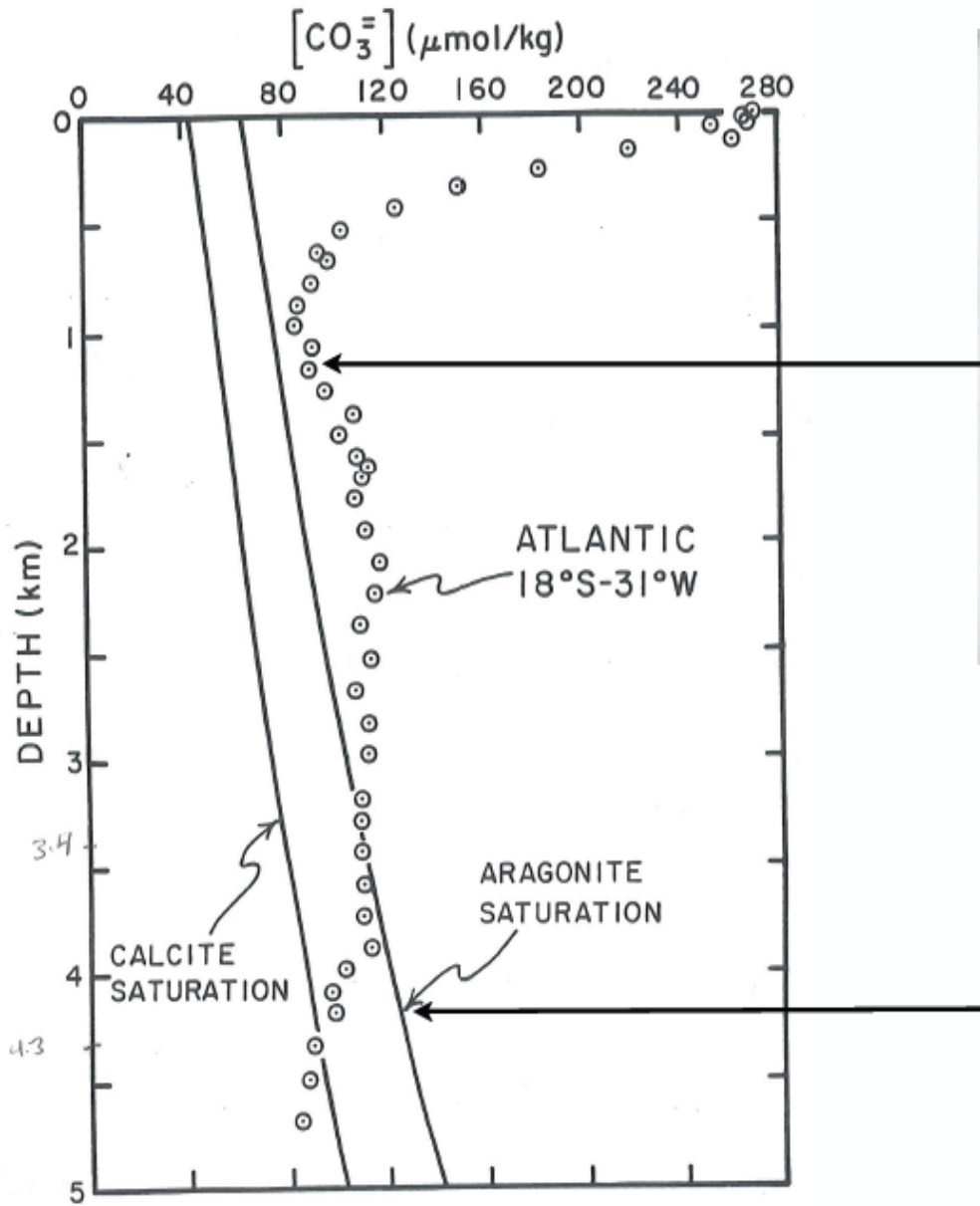




What other Relationships between Lab and Field that are critical to understanding carbonate dissolution??

- phosphate inhibition
- organic coatings

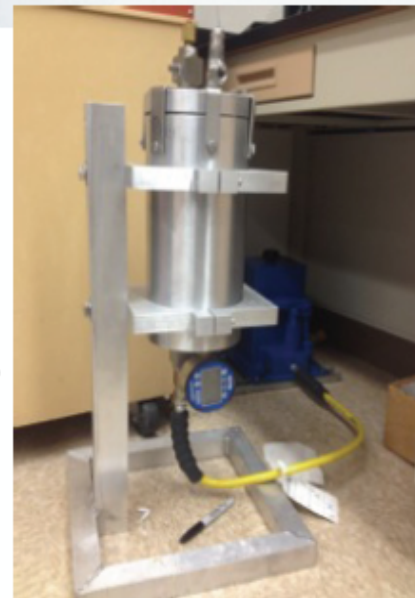
What role does pressure play in dissolution?



Broecker and Peng, 1982

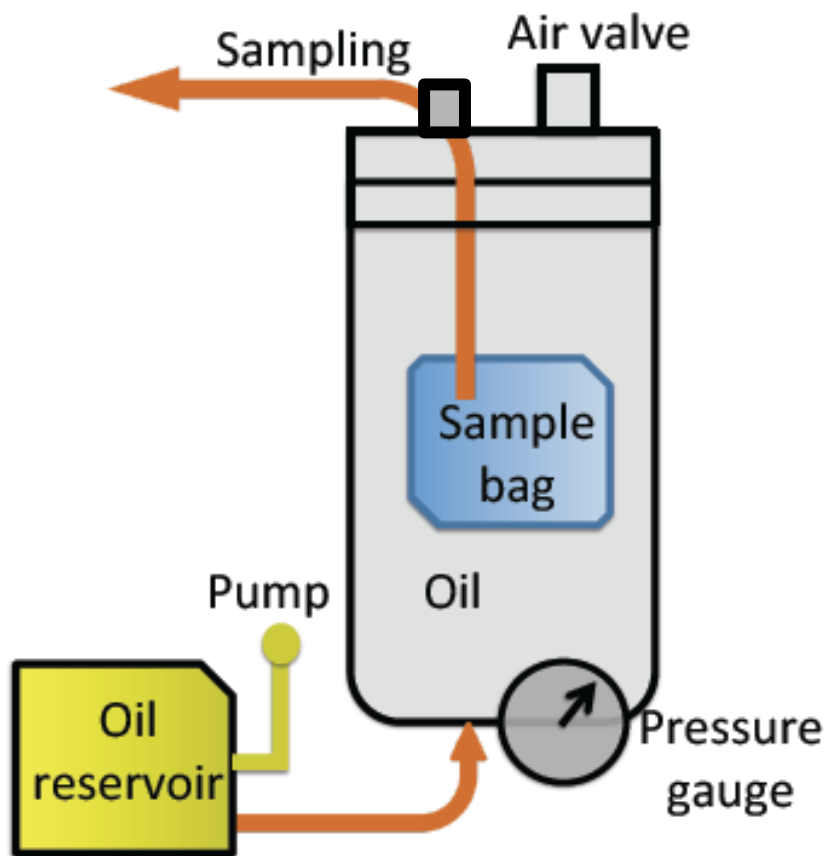


Professor Glickman, the lab practical joker, deftly places a single drop of hydrochloric acid on the back of Professor Bingham's neck.

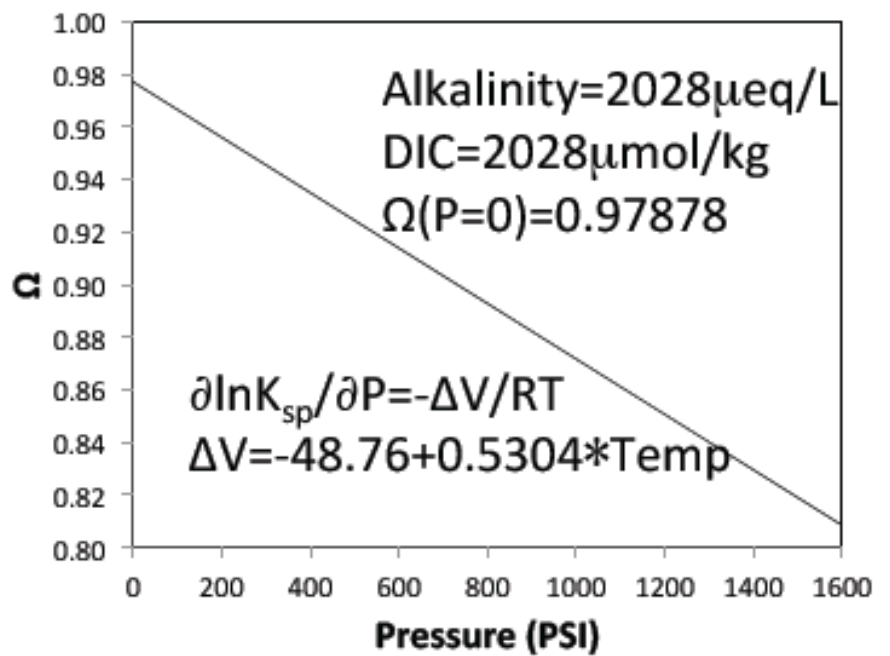


$$\Omega = \frac{[\text{Ca}^{2+}][\text{CO}_3^{2-}]}{K'_{\text{sp}}}$$

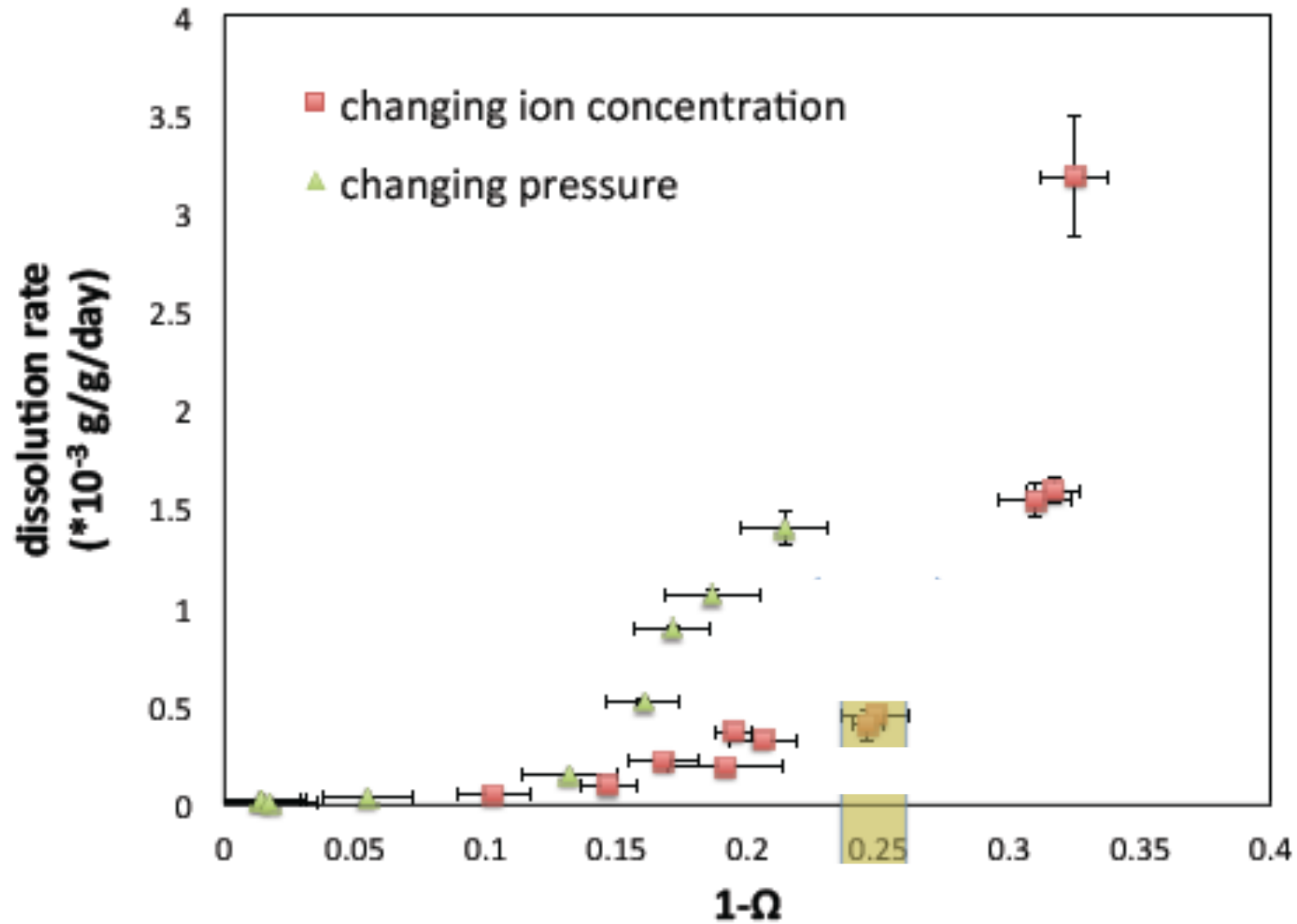
$$\Omega = \frac{[\text{Ca}^{2+}][\text{CO}_3^{2-}]}{K'_{\text{sp}}}$$



**Fig.1** Pressure chamber diagram



**Fig.2**  $\Omega$ -pressure correlation (calculated by CO2SYS)



**Fig.4** Dissolution rate vs. undersaturation

Variations in  $K_{sp}$ -pressure correlation expressions (empirical equations):

- $\partial \ln K_{sp} / \partial P = -\Delta V / RT$  (Millero and Berner, 1972; Duedall, 1972)
- $\text{Log}(K'_{sp}{}^P / K'_{sp}{}^1) = \Delta V(P-1) / 2.303RT$  (Edmond and Gieskes, 1970)
- $RT \ln(K'_{sp}{}^P / K'_{sp}{}^1) = -\Delta V(P-1) + 1/2 \Delta K(P-1)^2$  (Ingle, 1975)
- $\text{Ln}(K^P / K^0) = -(\Delta VP + 0.5 \Delta KP^2) / RT$  (Millero, 1979)

**Table 2** Apparent partial molal volume change for calcite dissolution

Substance	T (°C)	$\Delta V$ (cm <sup>3</sup> )	Author
Calcite	25	-34.4	Ingle (1975)
Calcite	25	-35.5	Ingle (1975)
Calcite	2	-42.3	Ingle (1975)
Calcite	2	-47.7	Ingle (1975)
Iceland spar	25	-35.5	Ingle (1975)
Oolites	2	-31.8	Ingle (1975)
Oolites	2	-33.1	Hawley and Pytkowicz (1969)
Oolites	22	-27.8	Hawley and Pytkowicz (1969)
Foraminifera	22	-30.7	Pytkowicz and Fowler (1967)
Calcite	25	-39.4	Millero and Berner (1969)
Calcite	20	-43.4	Duedall (1972)

A revised  $\Delta V$  from  $-37.6 \text{ cm}^3$  to  $-41.1 \text{ cm}^3$  would make the dissolution formulation equation agree.

-Pressure could contribute to enhanced dissolution in field

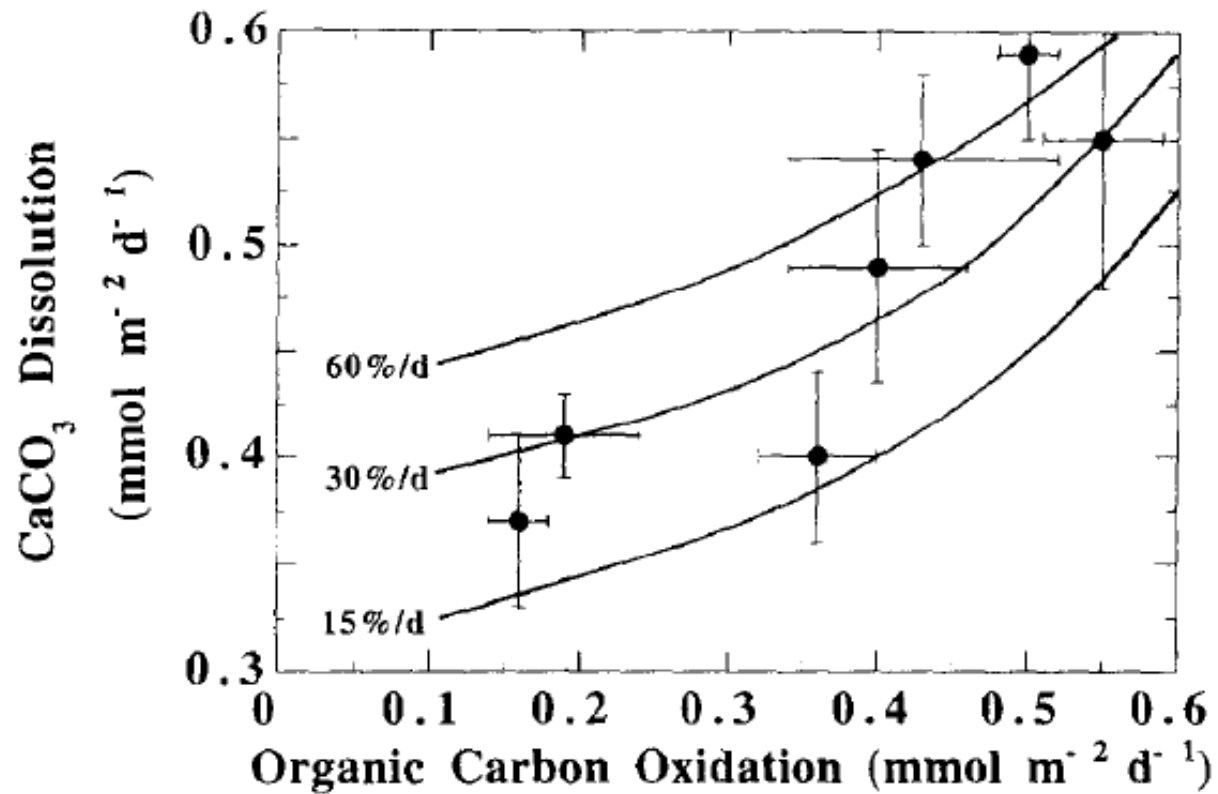
-Surface Area

-Different types of  $\text{CaCO}_3$ , mole fraction Mg

Take Home: no simple function will describe carbonate dissolution as function of omega

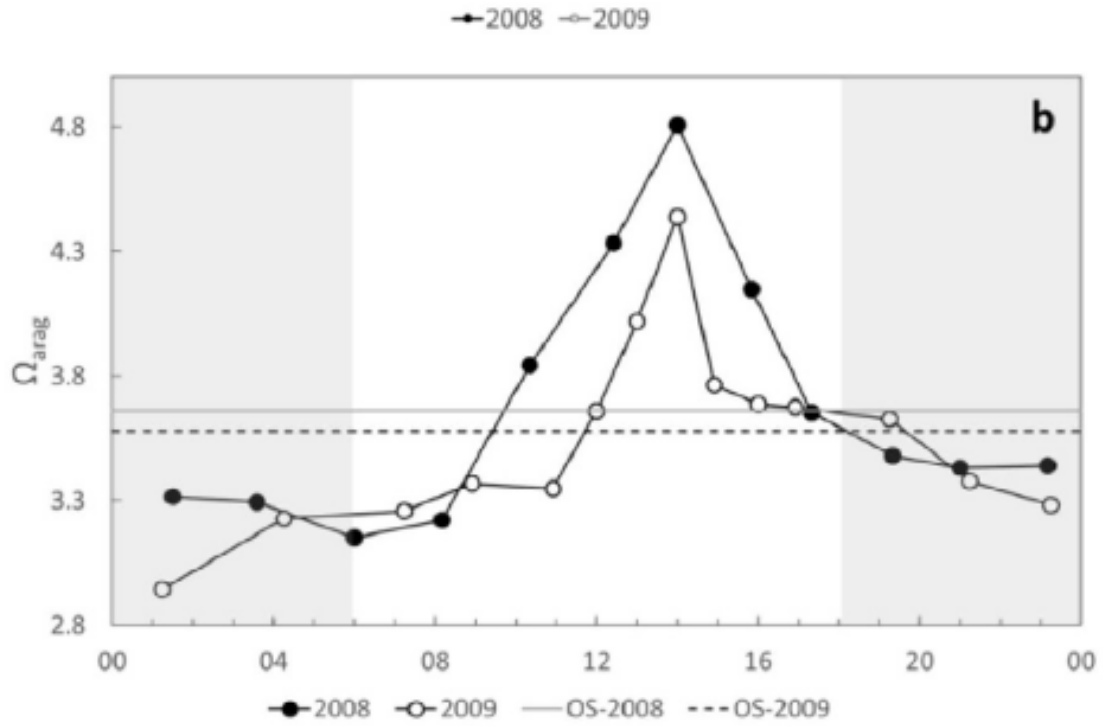
What else can field dissolution rates be attributed to?



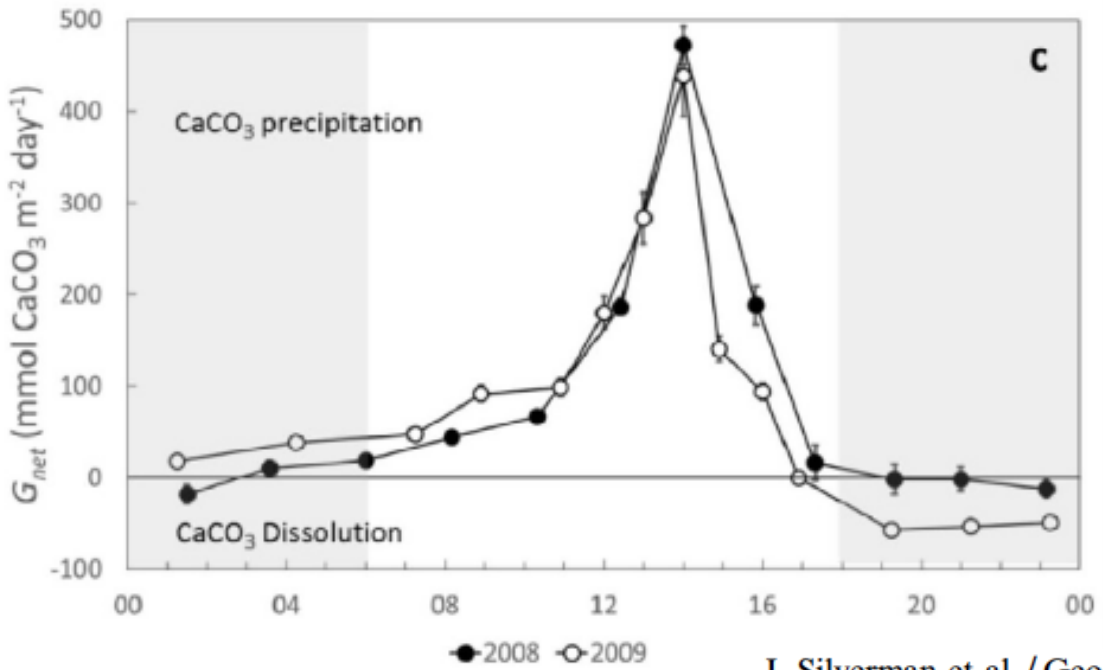


Equatorial Pacific, sea floor, all sites same Omega ~ 0.8

Carbonate dissolution 'driven' by C<sub>org</sub> oxidation



24 hours on a reef in Australia



Dissolution occurs at  $\Omega > 3$

Our Future Directions:

Field measurements of dissolution rate (novel Niskin apparatus)

Dissolution rate of more biogenics

Pressure effect

One of the big challenges we are facing is how to bridge our current understanding of  $\text{CaCO}_3$  dissolution based on controlled laboratory experiments to natural systems. This will require a holistic approach combining laboratory, field and numerical modelling studies.

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nature  
climate change

REVIEW ARTICLE

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Thanks