

Mechanisms of Mesoscale Physical/Biological Interaction

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Overview

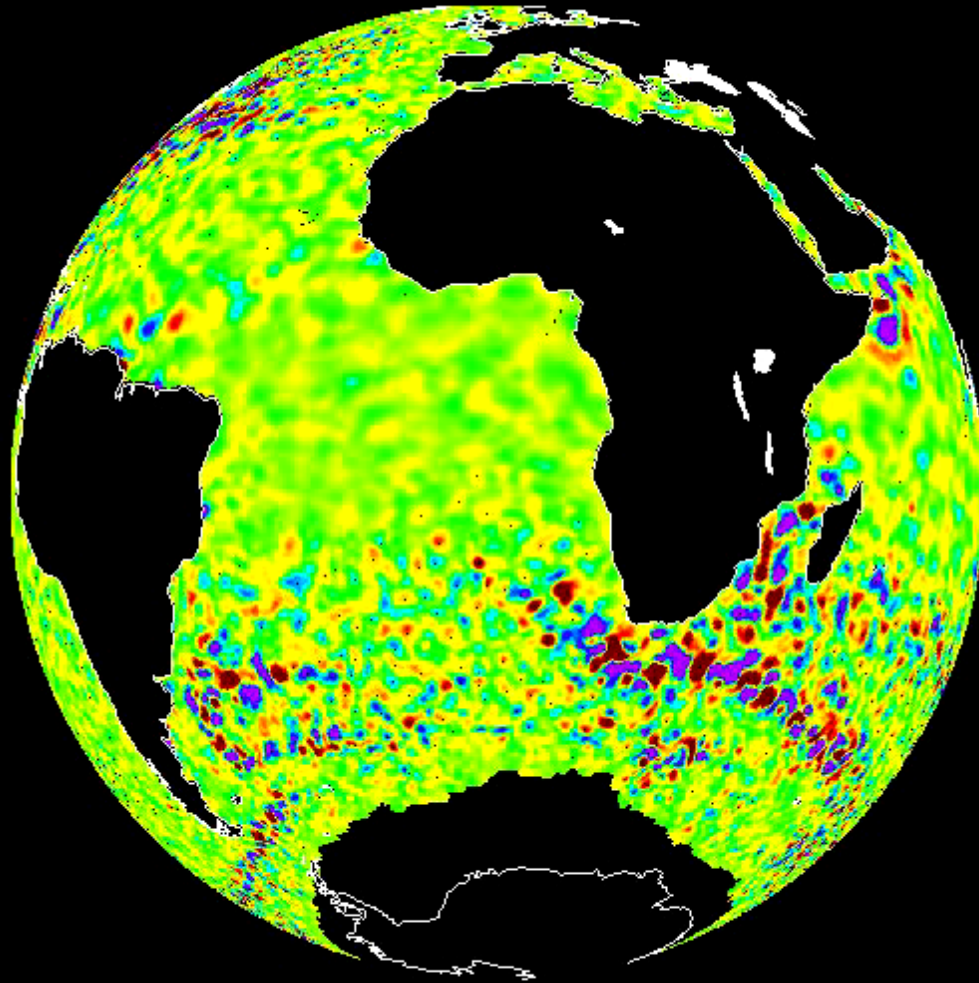
1. Horizontal advection of ecosystems by eddies
2. Eddy-induced vertical fluxes
3. Modulation of surface mixing by eddies



MODIS-Aqua May 9th, 2014
Gulf of Alaska

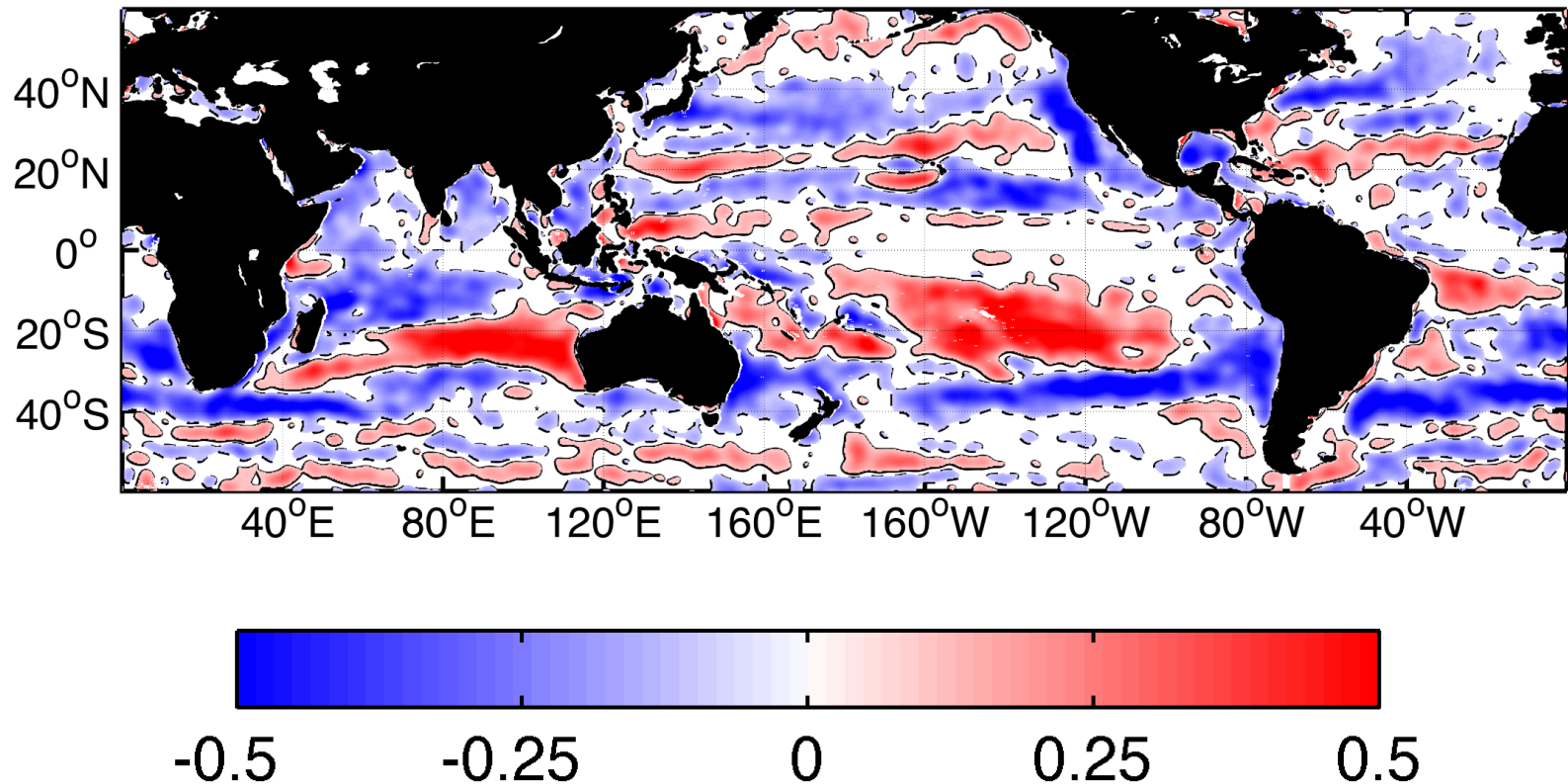
Observations of Nonlinear Mesoscale Eddies

SSH from the merged TOPEX and ERS-1/2 Data (Ducet *et al.*, 2000)

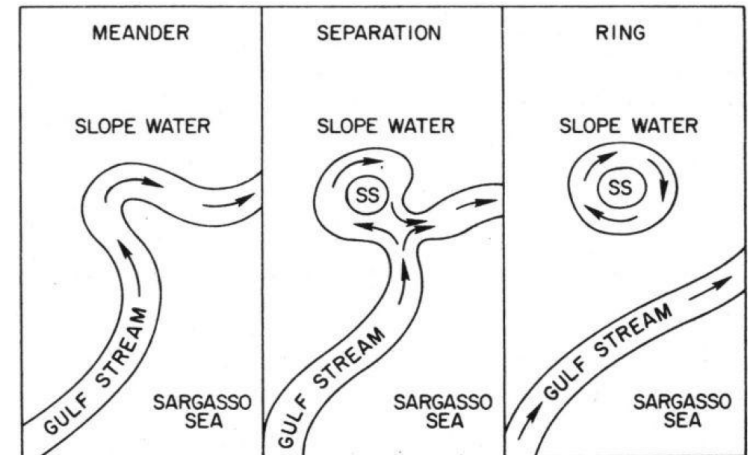
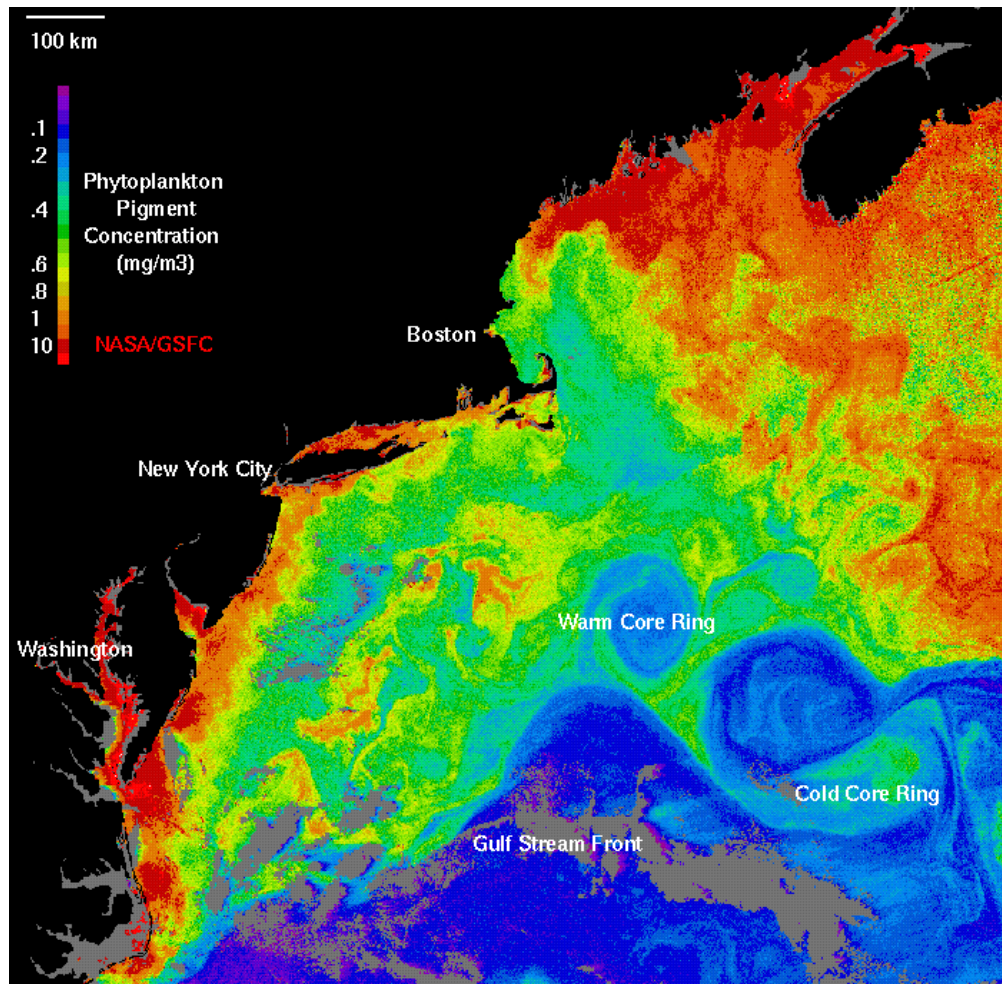


Exploring the Mechanisms Generating Observed Eddy Influence on Chlorophyll

Cross Correlation of CHL' and SLA



Mechanisms of mesoscale physical-biological interaction: Trapped ecosystems



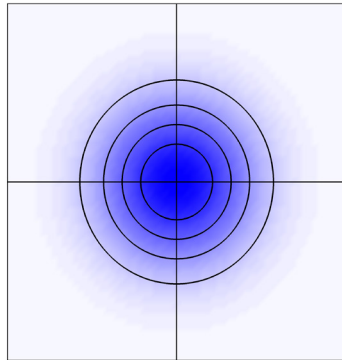
Warm Core Ring Executive Committee, 1982

Mechanisms of Mesoscale Phys/Bio Interaction

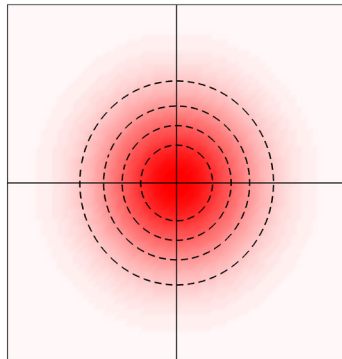
Eddy Trapping

Idealized Composite Average

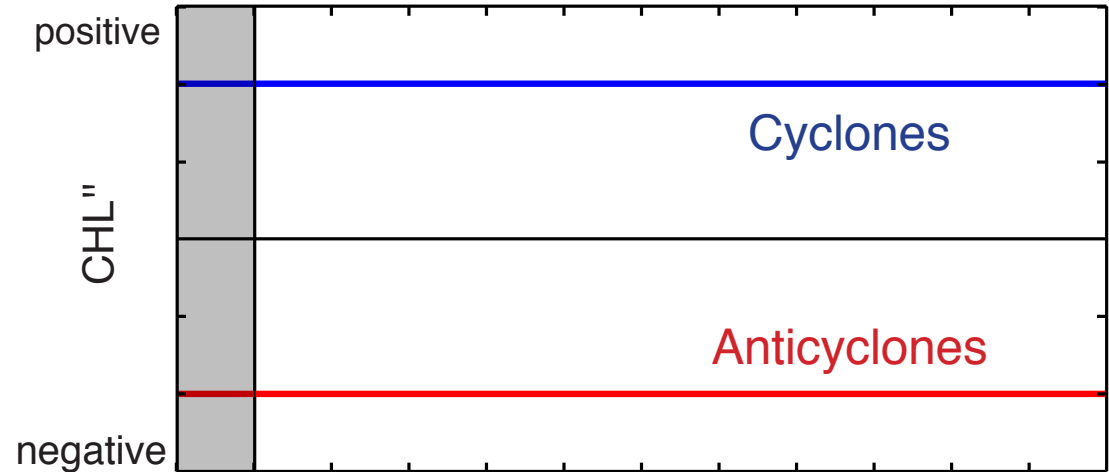
Anticyclones



Cyclones



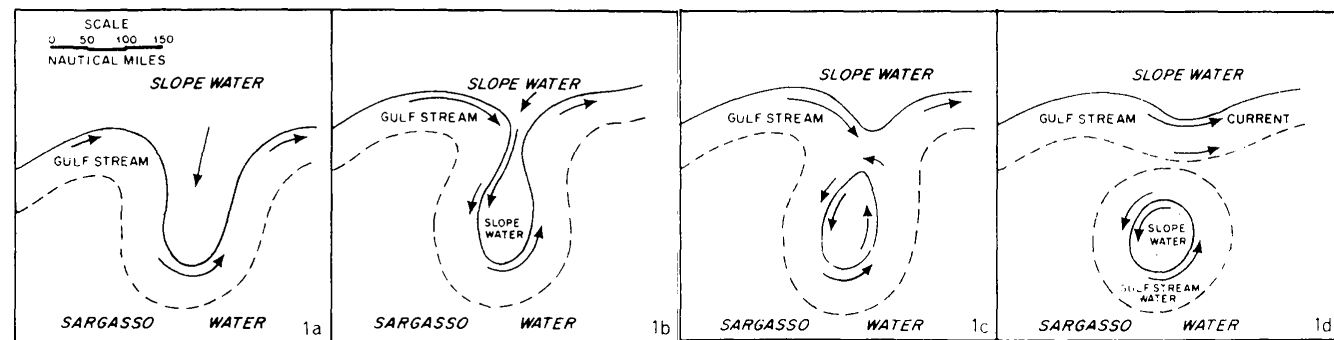
Idealized Composite Average
Time Series



intensification

maturation

Gaube et al., 2014

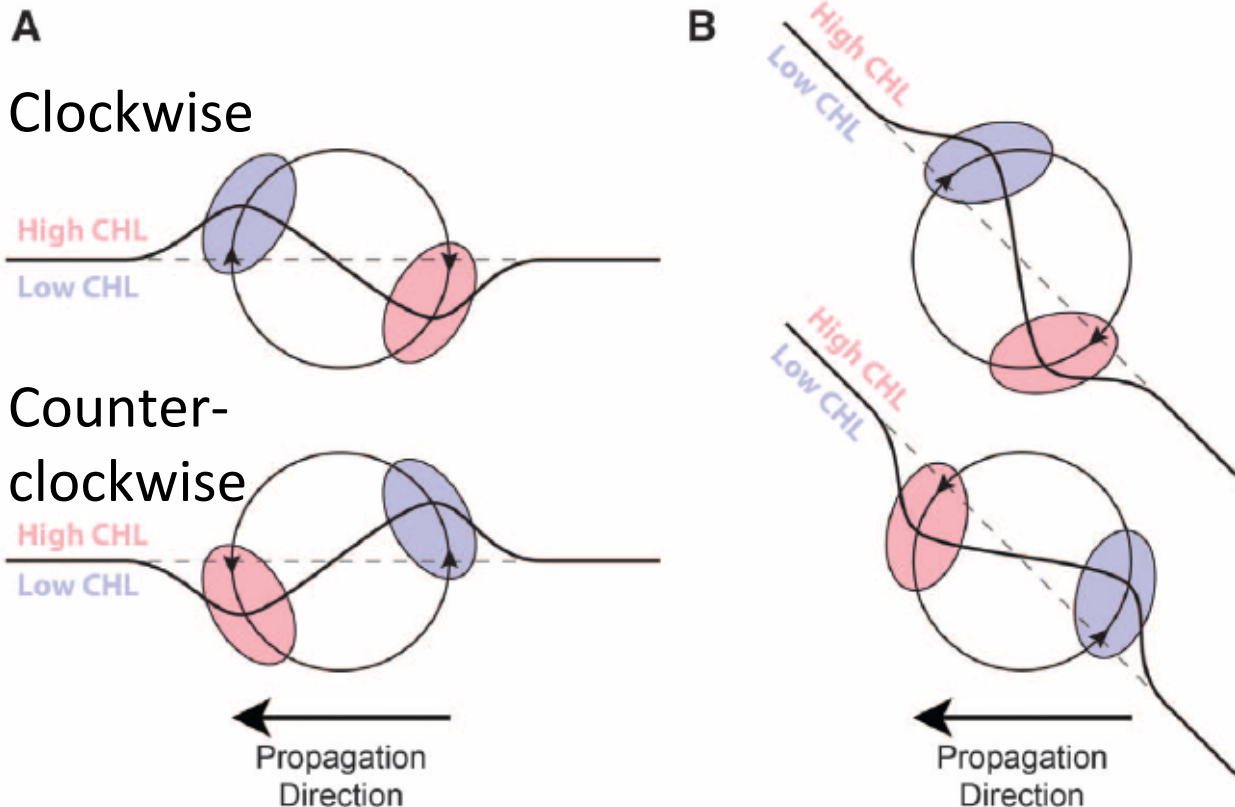


Parker 1971

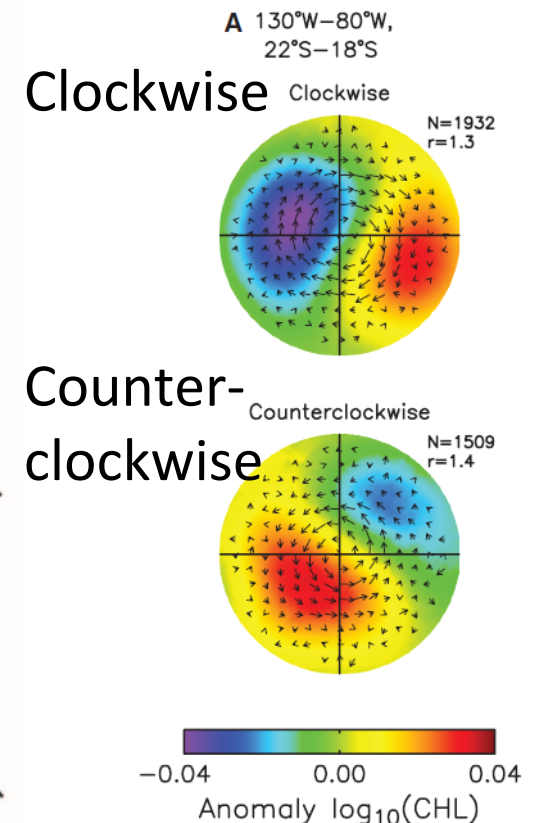
The sign of the CHL" resulting from the **trapping of CHL** is a function of the cross current CHL gradient.

Mechanisms of mesoscale physical-biological interaction: Eddy stirring

Theory

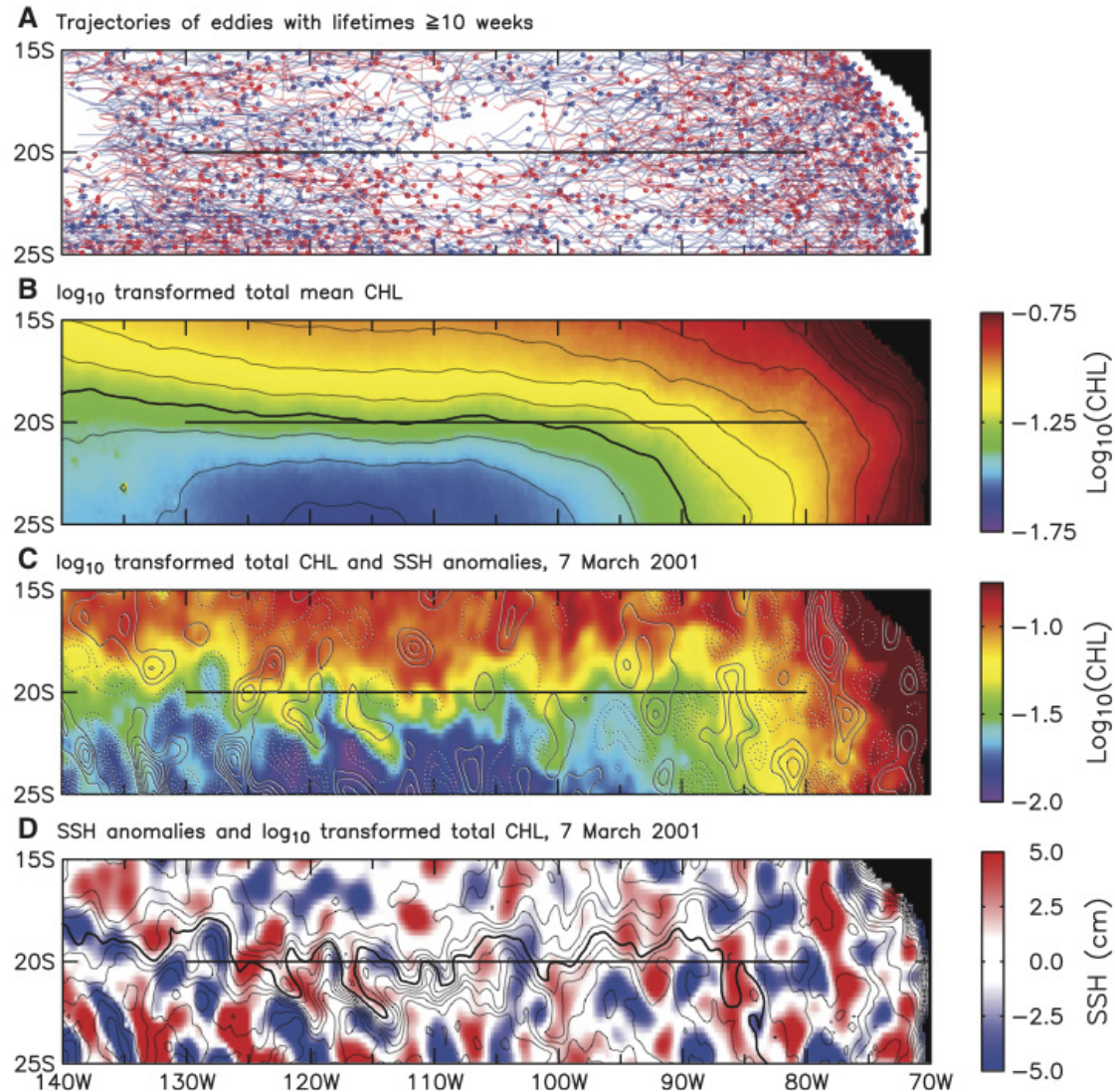


Observations



Eddies of the South Pacific

region investigated by Dandonneau et al., 2003; Killworth, 2004; Dandonneau et al., 2004 and Killworth et al., 2004

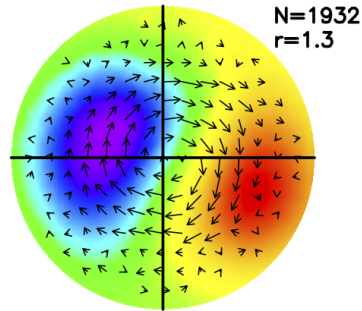


Eddy Composite Averages of Chlorophyll Anomaly

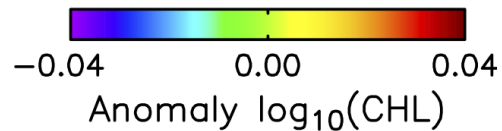
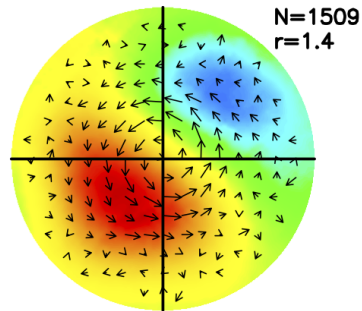
eddies along 20°S section

130°W–80°W,
22°S–18°S

Clockwise



Counterclockwise



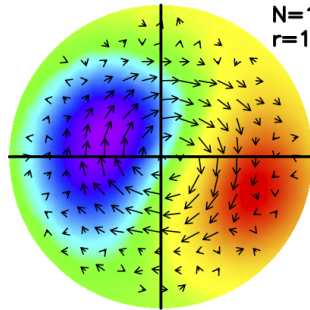
Eddy Composite Averages of Chlorophyll Anomaly

eddies along 20°S section

130°W–80°W,
22°S–18°S

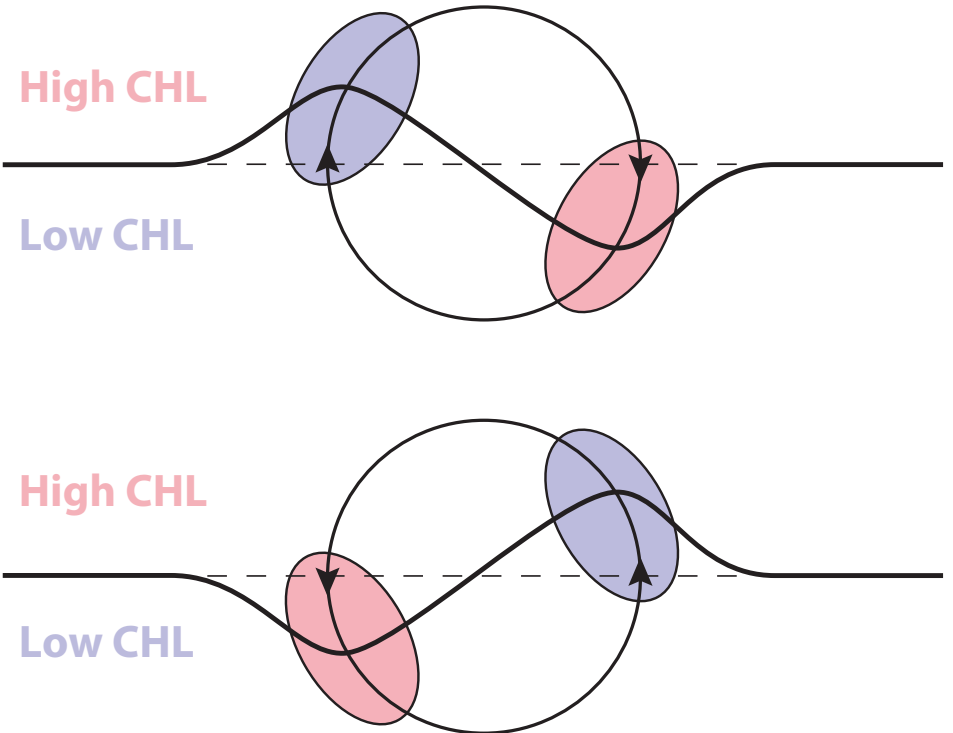
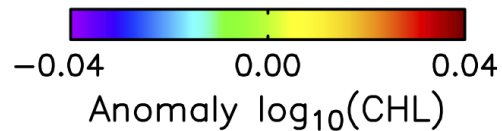
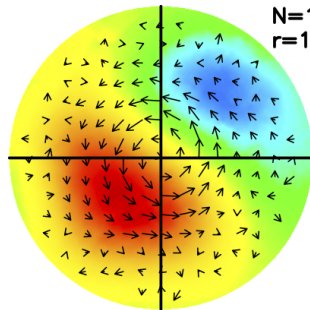
Clockwise

N=1932
r=1.3



Counterclockwise

N=1509
r=1.4



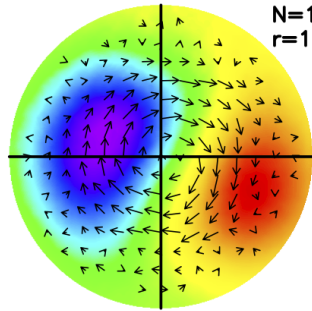
Global Composites of Chlorophyll Anomaly

northward CHL gradient

130°W–80°W,
22°S–18°S

Clockwise

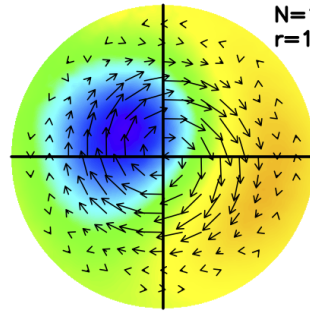
N=1932
r=1.3



Global, 15°–45°
Northward ∇ CHL

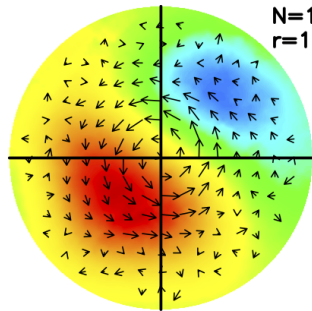
Clockwise

N=113,889
r=1.7



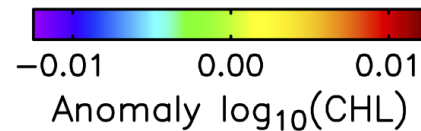
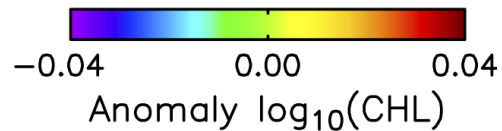
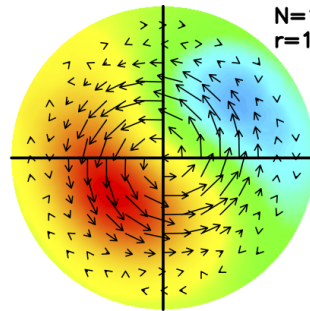
Counterclockwise

N=1509
r=1.4



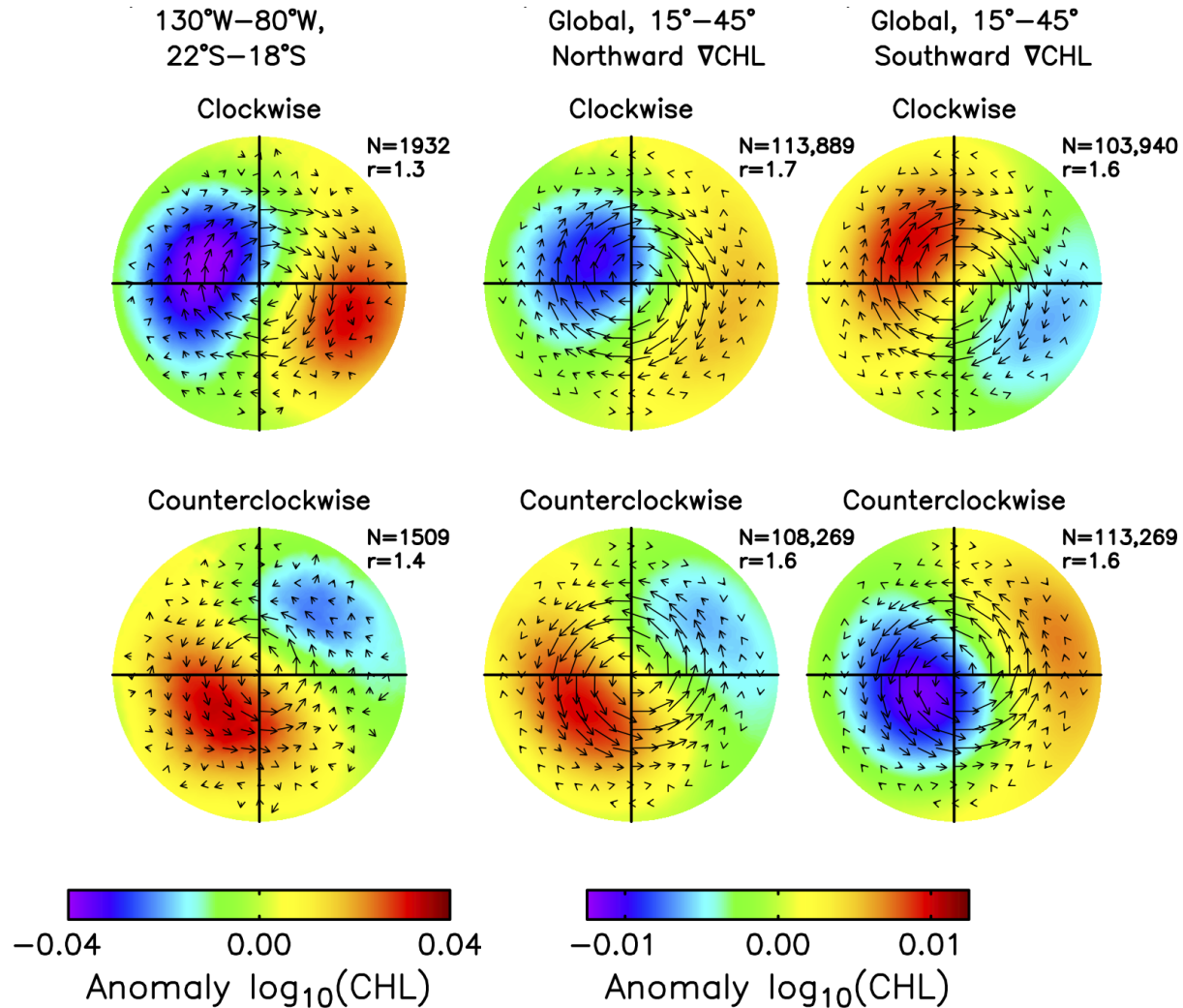
Counterclockwise

N=108,269
r=1.6

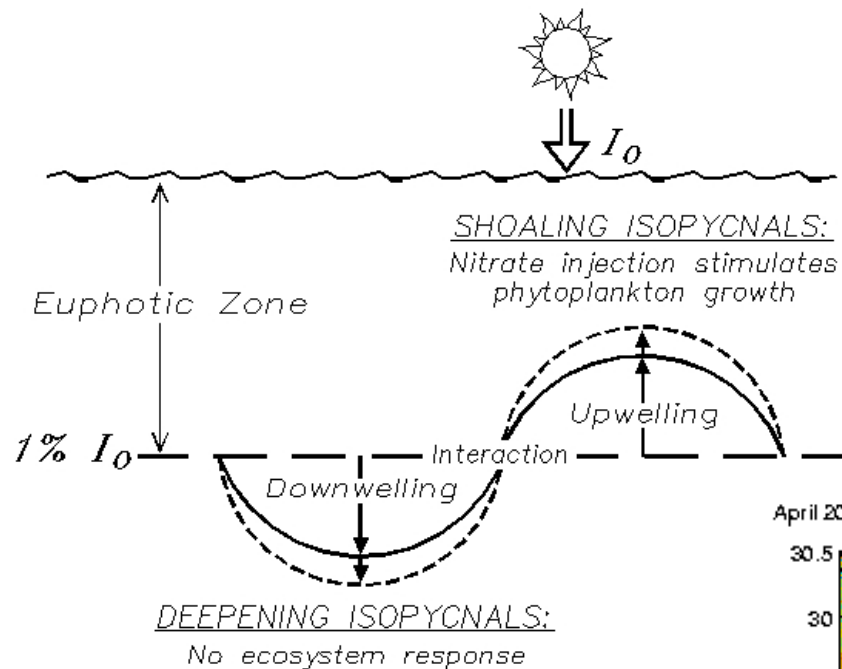


Global Composite of Chlorophyll Anomaly

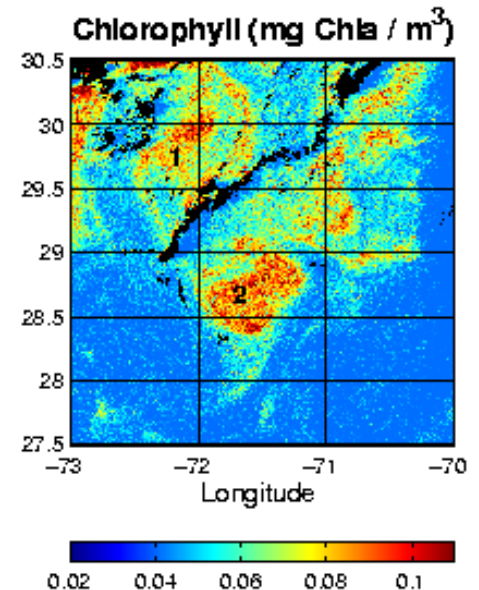
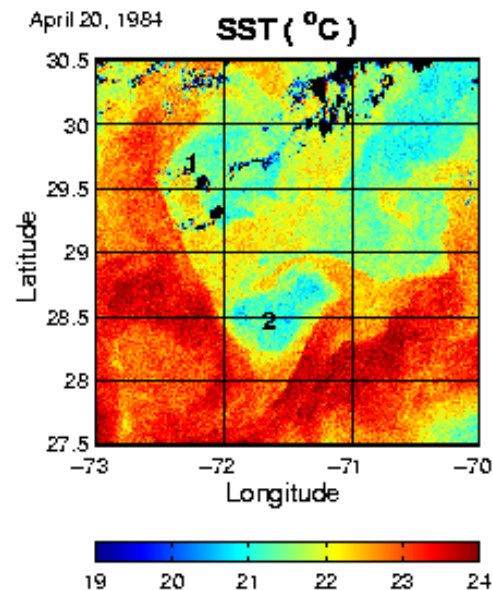
northward and southward CHL gradient



Mechanisms of mesoscale physical-biological interaction: Eddy formation / intensification

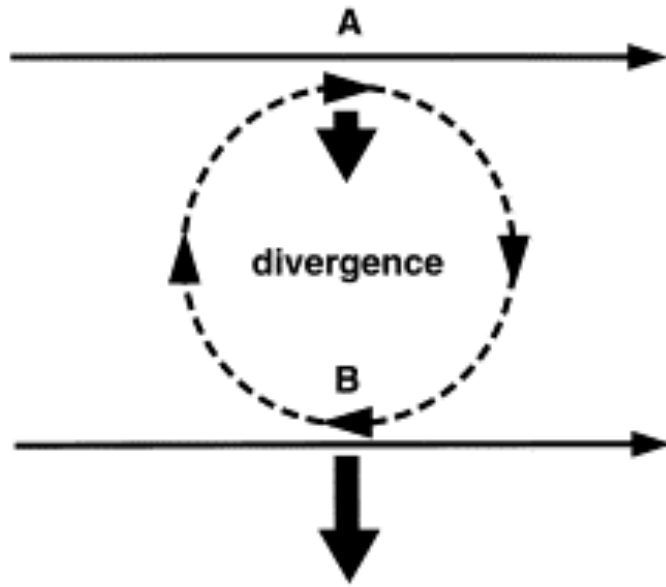


McGillicuddy and Robinson, 1997



McGillicuddy et al., 2001

Mechanisms of mesoscale physical-biological interaction: Eddy-induced Ekman pumping



Key:

wind



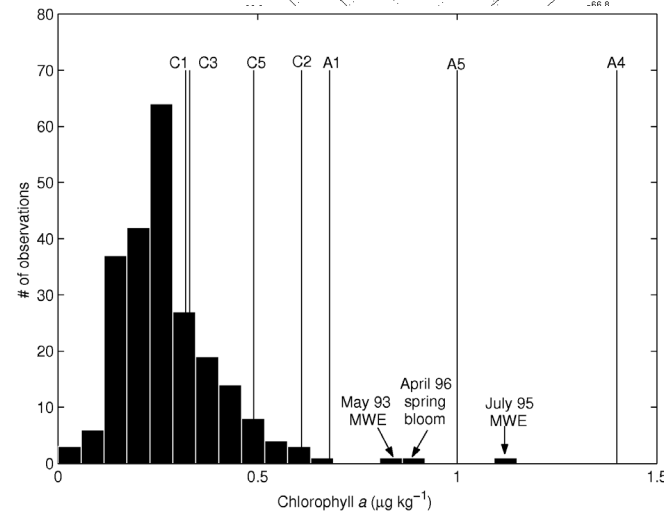
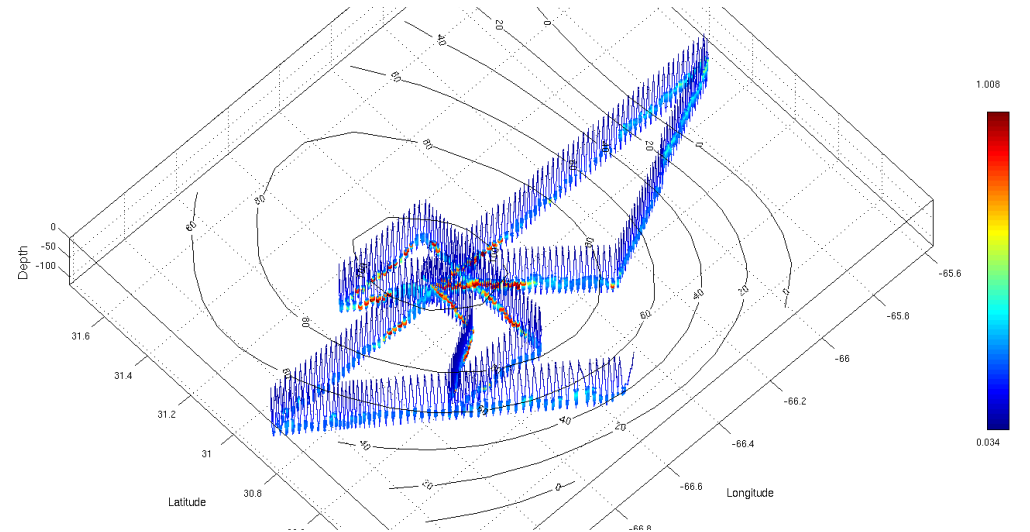
eddy current



Ekman transport



Dewar and Flierl, 1987
Martin and Richards, 2001



Fluorescence
 $\mu\text{g Chl } a \text{ L}^{-1}$

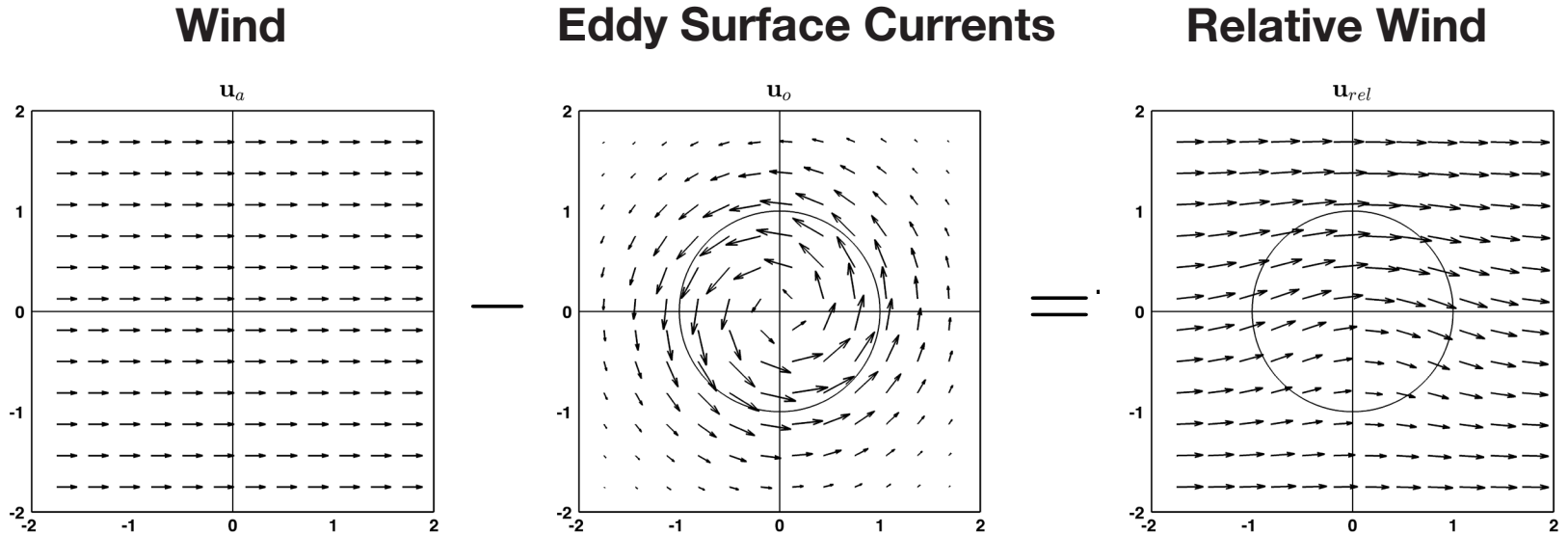
BATS Chl *a* 1988-2003

McGillicuddy et al., 2007

Surface Current-Induced Ekman Pumping

How it works:

- The surface currents of eddies impart a curl in the relative wind.
- Dewar and Flier, 1987; Martin and Richards, 2001



$$\mathbf{u}_{rel} = \mathbf{u}_a - \mathbf{u}_o$$

$$\boldsymbol{\tau} = \rho_a C_D \mathbf{u}_{rel} |\mathbf{u}_{rel}|$$

$$W_{tot} = \frac{1}{\rho_o} \nabla \times \left(\frac{\boldsymbol{\tau}}{(f + \zeta)} \right)$$

Surface Current-Induced Ekman Pumping

$$\begin{aligned}
 W_{tot} &= \frac{1}{\rho_o} \nabla \times \left(\frac{\boldsymbol{\tau}}{(f + \zeta)} \right) \\
 &= \frac{1}{\rho_o} \left[\frac{\nabla \times \boldsymbol{\tau}}{f + \zeta} - \frac{\boldsymbol{\tau} \times \nabla \zeta}{(f + \zeta)^2} - \frac{\boldsymbol{\tau} \times \nabla f}{(f + \zeta)^2} \right] \\
 &= \underbrace{\frac{\nabla \times \boldsymbol{\tau}}{\rho_o (f + \zeta)}}_{W_c} - \underbrace{\frac{1}{\rho_o (f + \zeta)^2} \left(\tau^y \frac{\partial \zeta}{\partial x} - \tau^x \frac{\partial \zeta}{\partial y} \right)}_{W_\zeta} + \underbrace{\frac{\beta \tau^x}{\rho_o (f + \zeta)^2}}_{W_\beta}
 \end{aligned}$$

- Surface current-induced Ekman pumping
- Dominant term at scale $\geq O(100 \text{ km})$

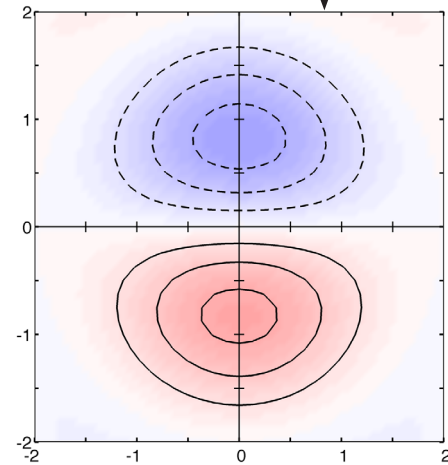
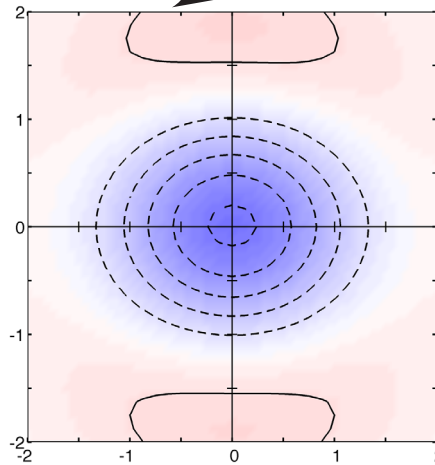
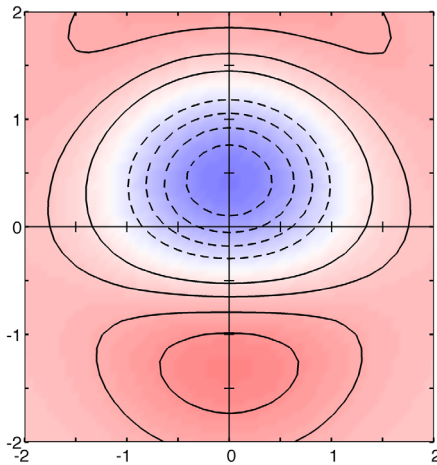
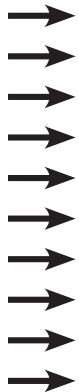
- Vorticity gradient-induced Ekman pumping
- Dominant term at scale $\sim O(10 \text{ km})$

- "Beta" Ekman pumping
- Negligible at scale $< O(100 \text{ km})$

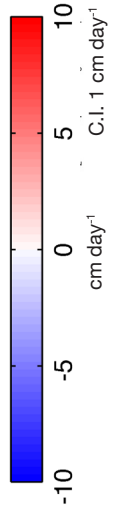
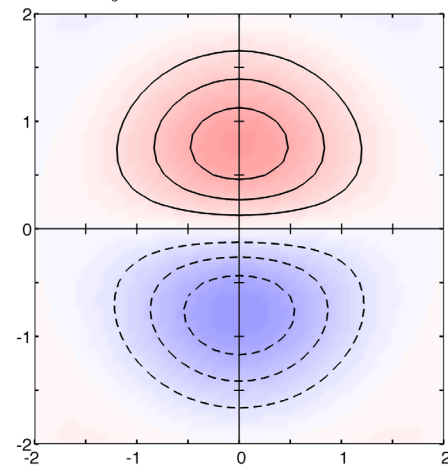
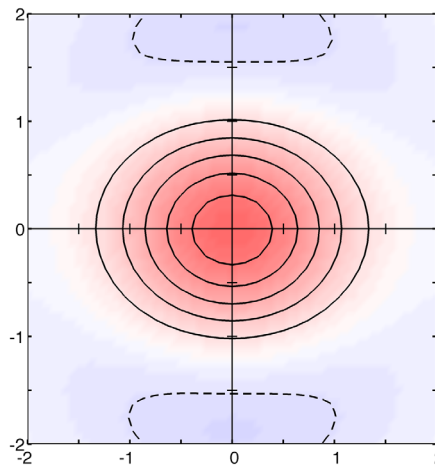
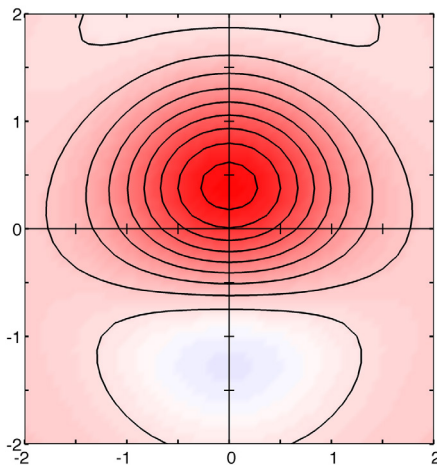
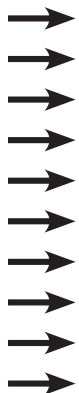
Surface Current-Induced Ekman Pumping

$$W_{tot} = \underbrace{\frac{\nabla \times \boldsymbol{\tau}}{\rho_o (f + \zeta)}}_{W_c} - \underbrace{\frac{1}{\rho_o (f + \zeta)^2} \left(\tau^y \frac{\partial \zeta}{\partial x} - \tau^x \frac{\partial \zeta}{\partial y} \right)}_{W_\zeta}$$

N.H. Cyclones



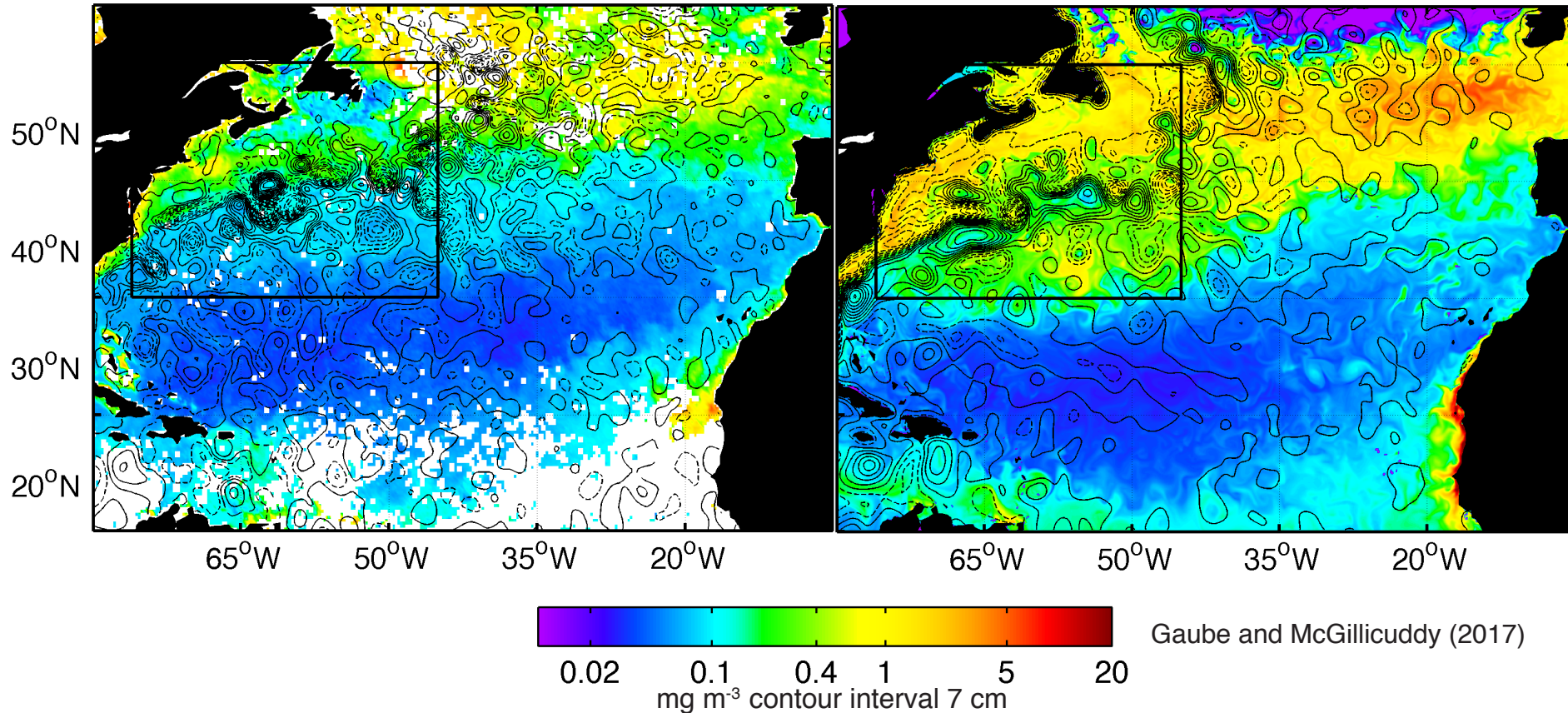
N.H. Anticyclones



Surface Chlorophyll and Sea Level Anomalies in the North Atlantic Ocean

SeaWiFS $\log_{10}(\text{CHL})$ and AVISO SLA
June 6th, 1998

BEC $\log_{10}(\text{CHL})$ at surface and POP SLA
June 14th



BEC=Biogeochemical Ecosystem Model: 3-phytoplankton types and 1-zooplankton.
Limiting nutrients nitrate (NO_3), ammonium (NH_4), phosphate (PO_4), iron (Fe), and silicate (SiO_3).

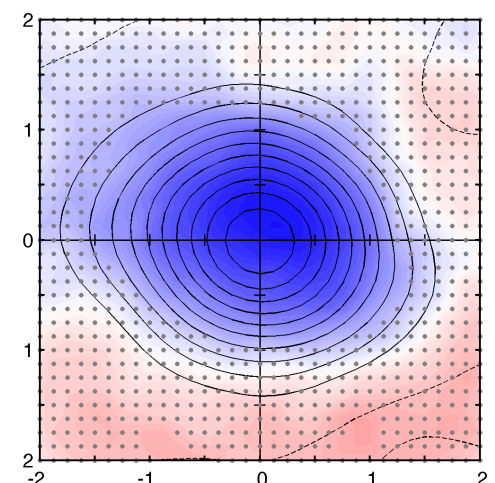
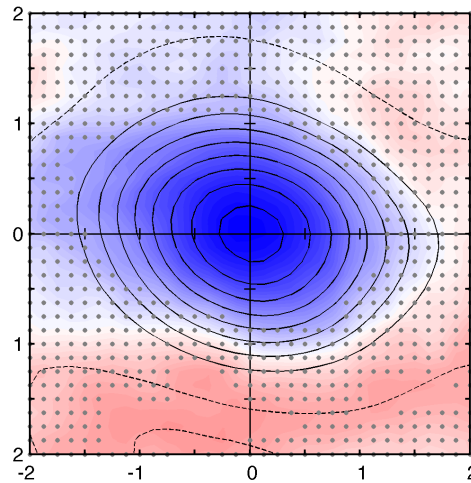
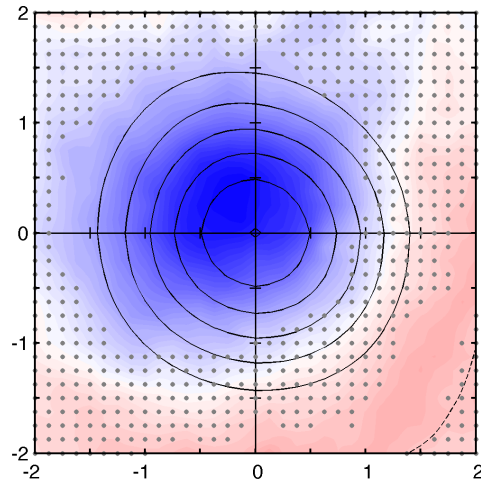
Composite Averages of CHL Anomalies in Westward Propagating Eddies

Observations

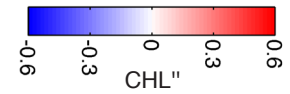
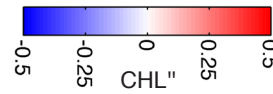
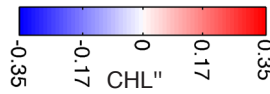
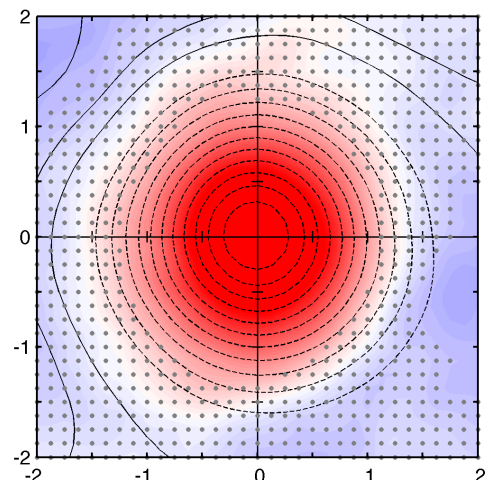
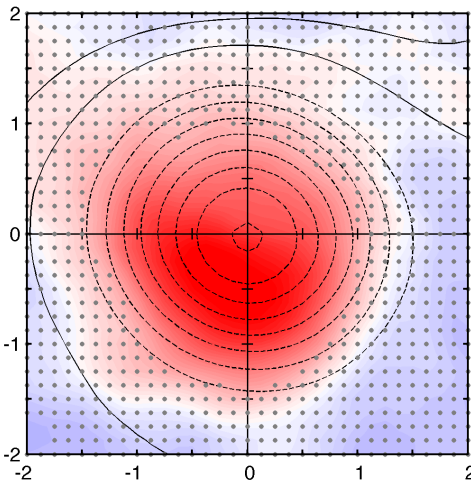
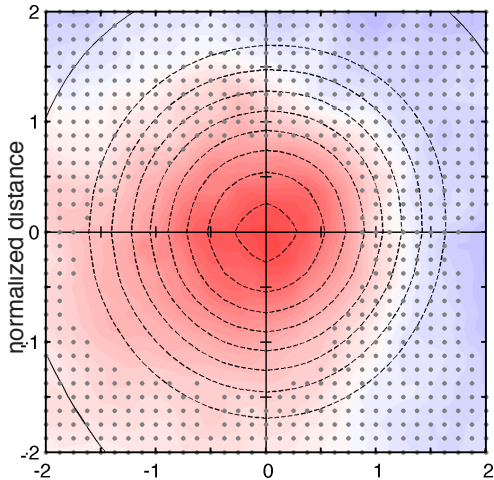
POP-BEC with
Eddy/wind

POP-BEC without
Eddy/wind

Anticyclones

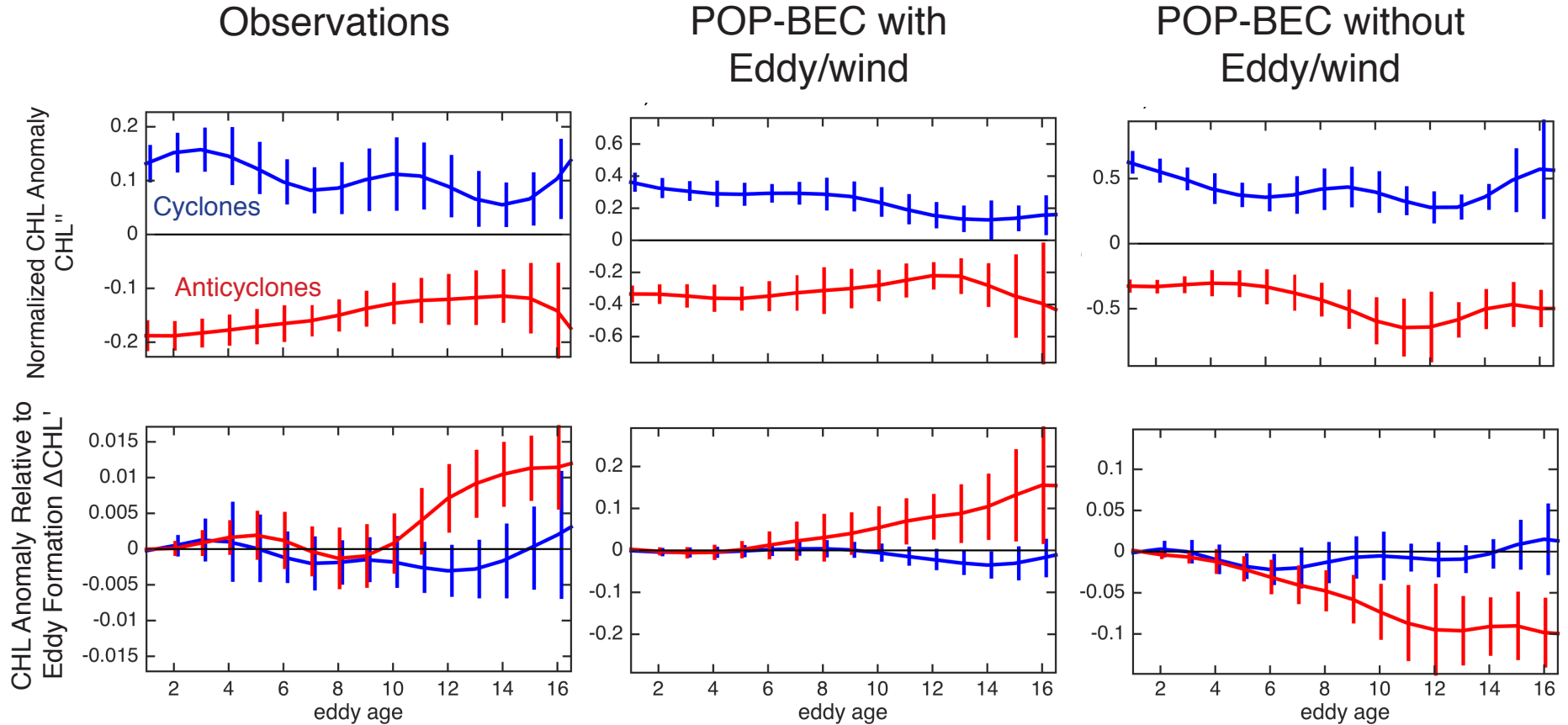


Cyclones



The Evolution of CHL in Gulf Stream Eddies

CHL anomalies averaged within the interiors of eddies as a function of time since eddy formation

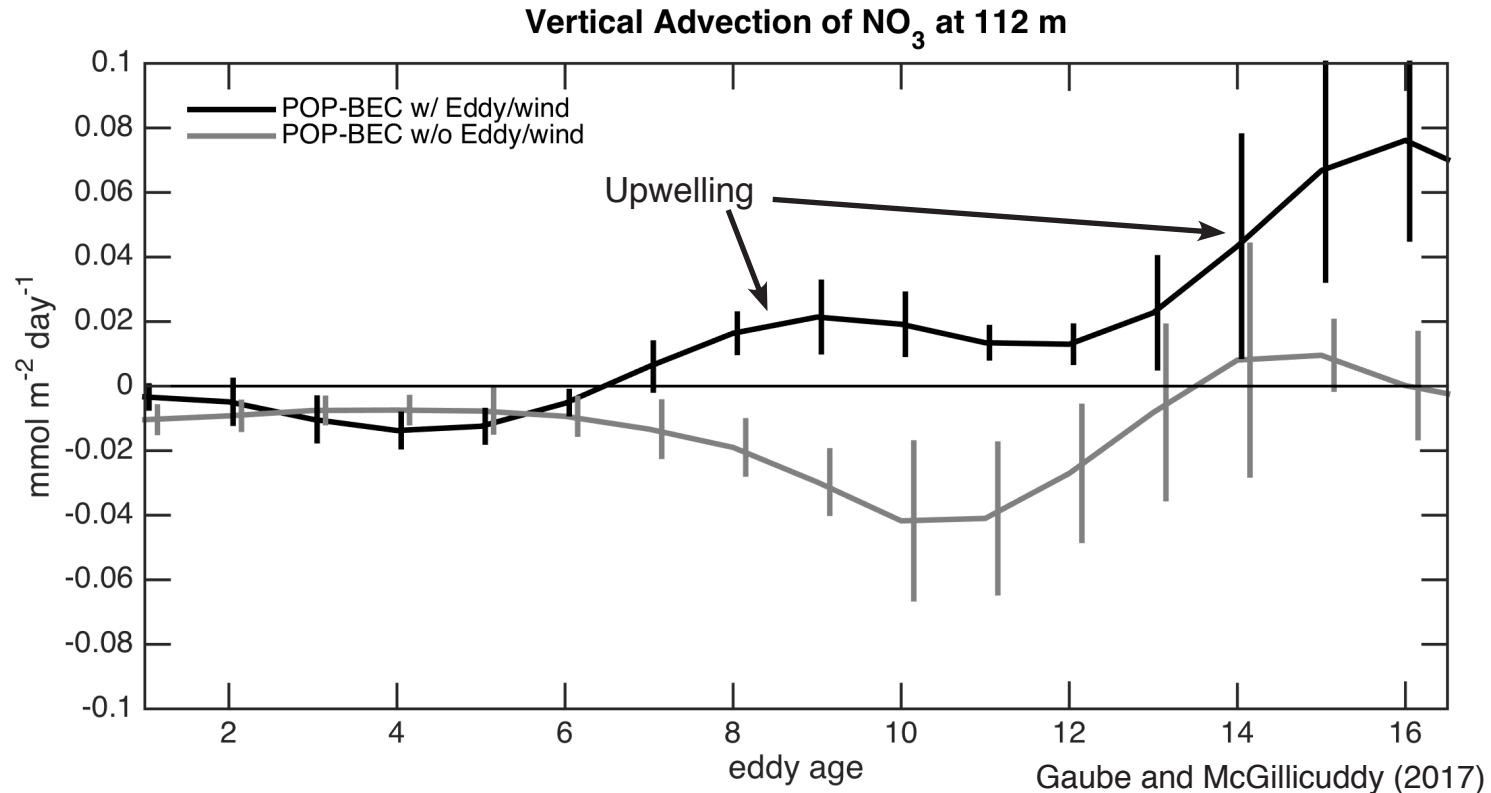


$$\begin{aligned}\Delta CHL'(x, y, t) &= CHL'(x, y, t) - CHL'(x, y, t_0) \\ &= [CHL(x, y, t) - \langle CHL(x, y, t) \rangle] - [CHL(x, y, t_0) - \langle CHL(x, y, t_0) \rangle]\end{aligned}$$

Gaube and McGillicuddy (2017)

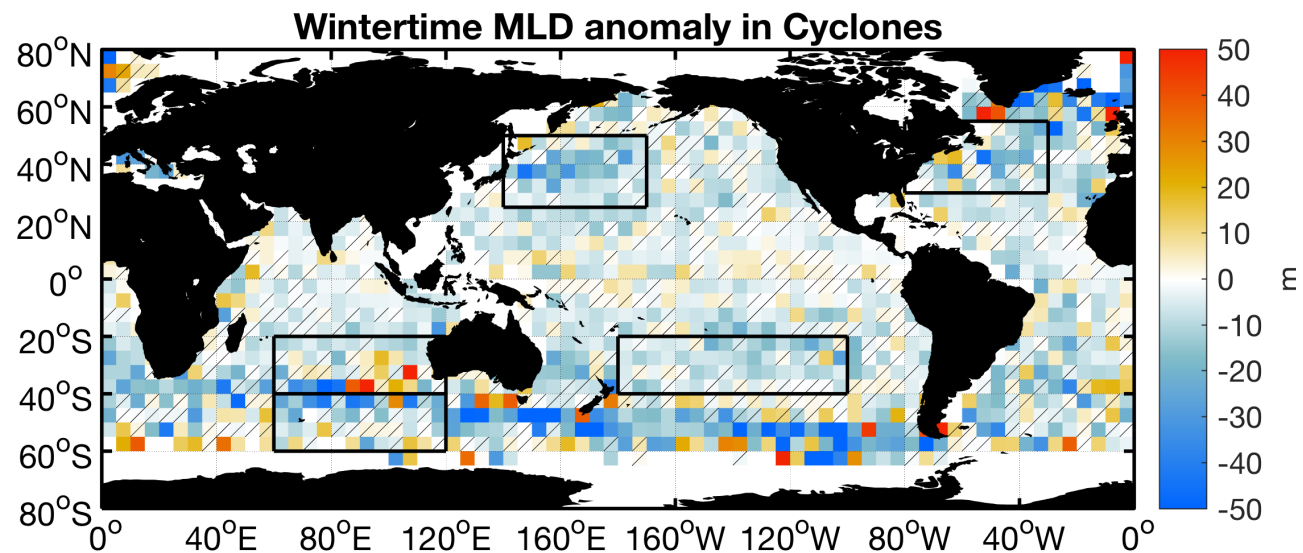
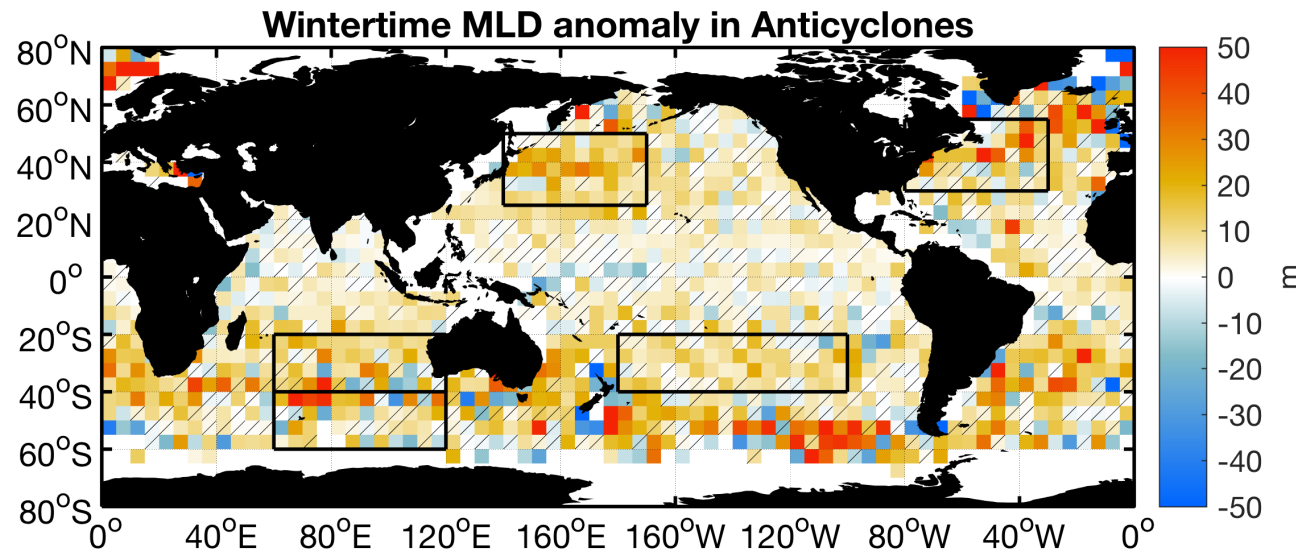
Vertical Nitrate Advection in Gulf Stream Eddies

From the POP-BEC solution



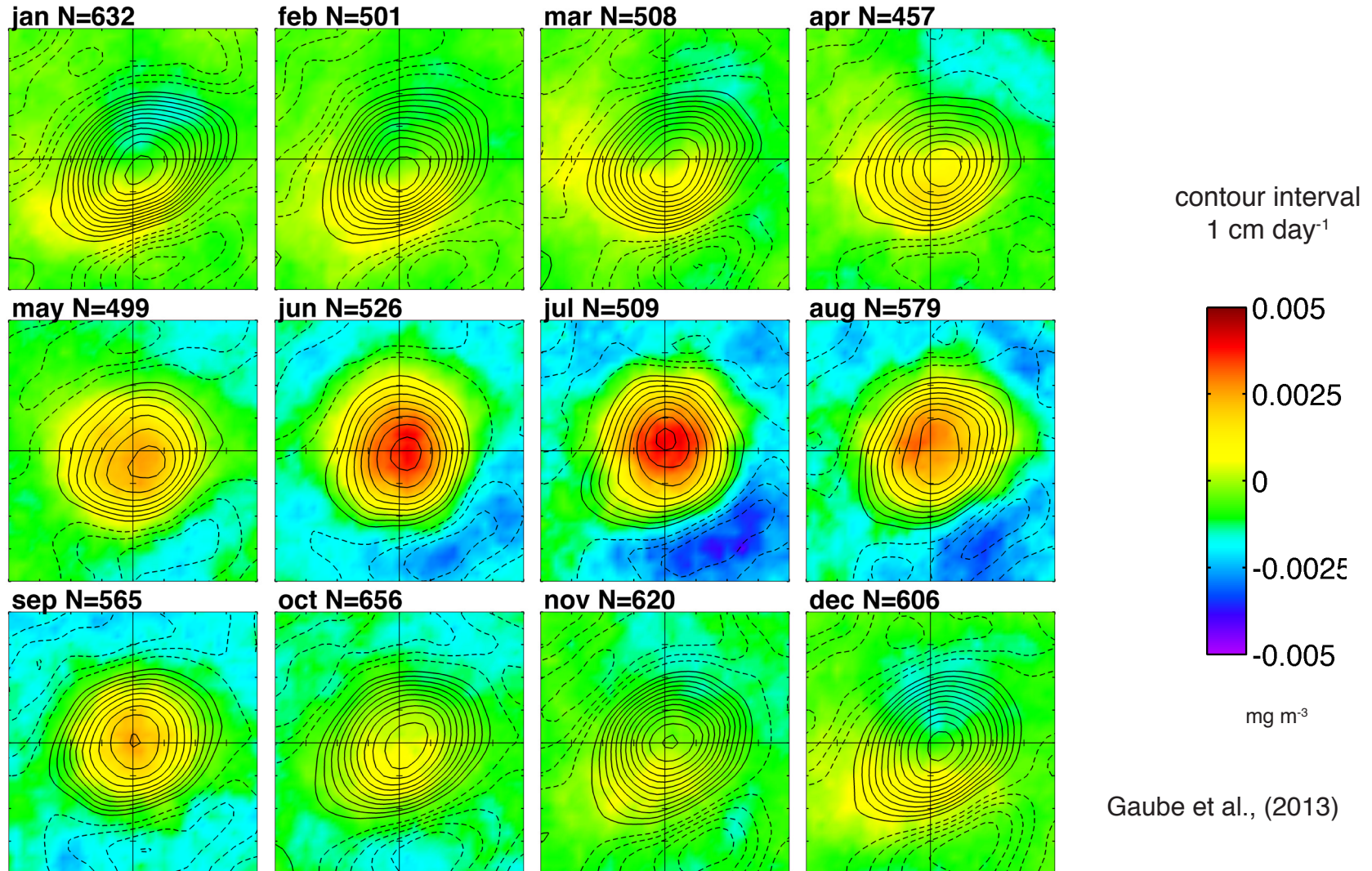
- Upwelling of NO_3 begins at week 7 and becomes significant at week 8 in model that includes eddy-induced Ekman pumping.
- The CHL response starts at week 6-7, but does not become significant until week 11 (15 days later).

Eddy-mediated MLD modulations



Composite Averaged Chlorophyll Anomalies by Month

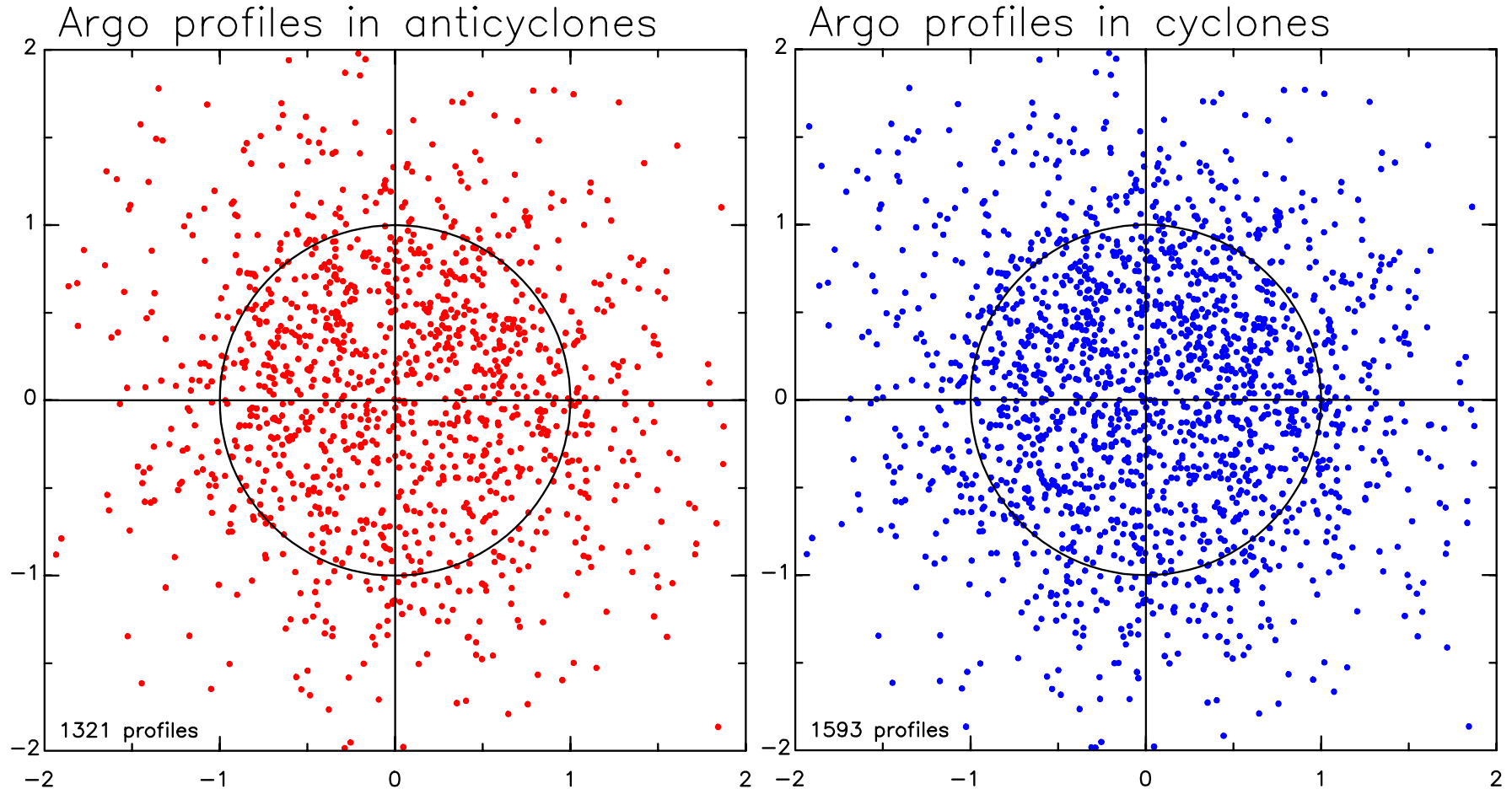
filtered SeaWiFS chlorophyll with contours of QuikSCAT Ekman pumping (anticyclones)



What is driving the seasonality of CHL anomalies in the interiors of South Indian Ocean anticyclones?

Argo Float Profiles within South Indian Ocean Eddies

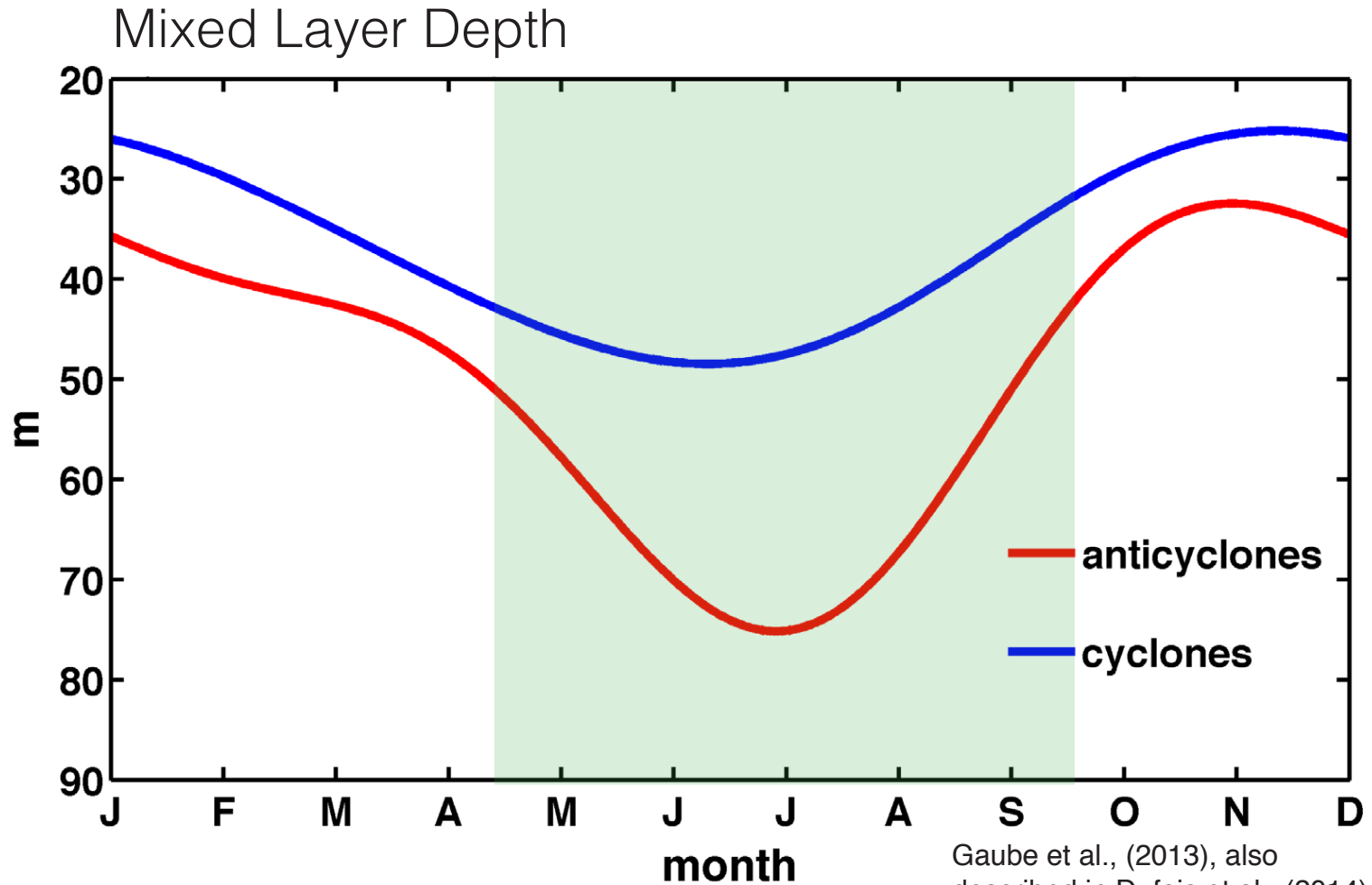
Argo profiles from which MLD seasonal cycles are calculated



Seasonal cycles of MLD are calculated from the above profiles as a function of eddy polarity.

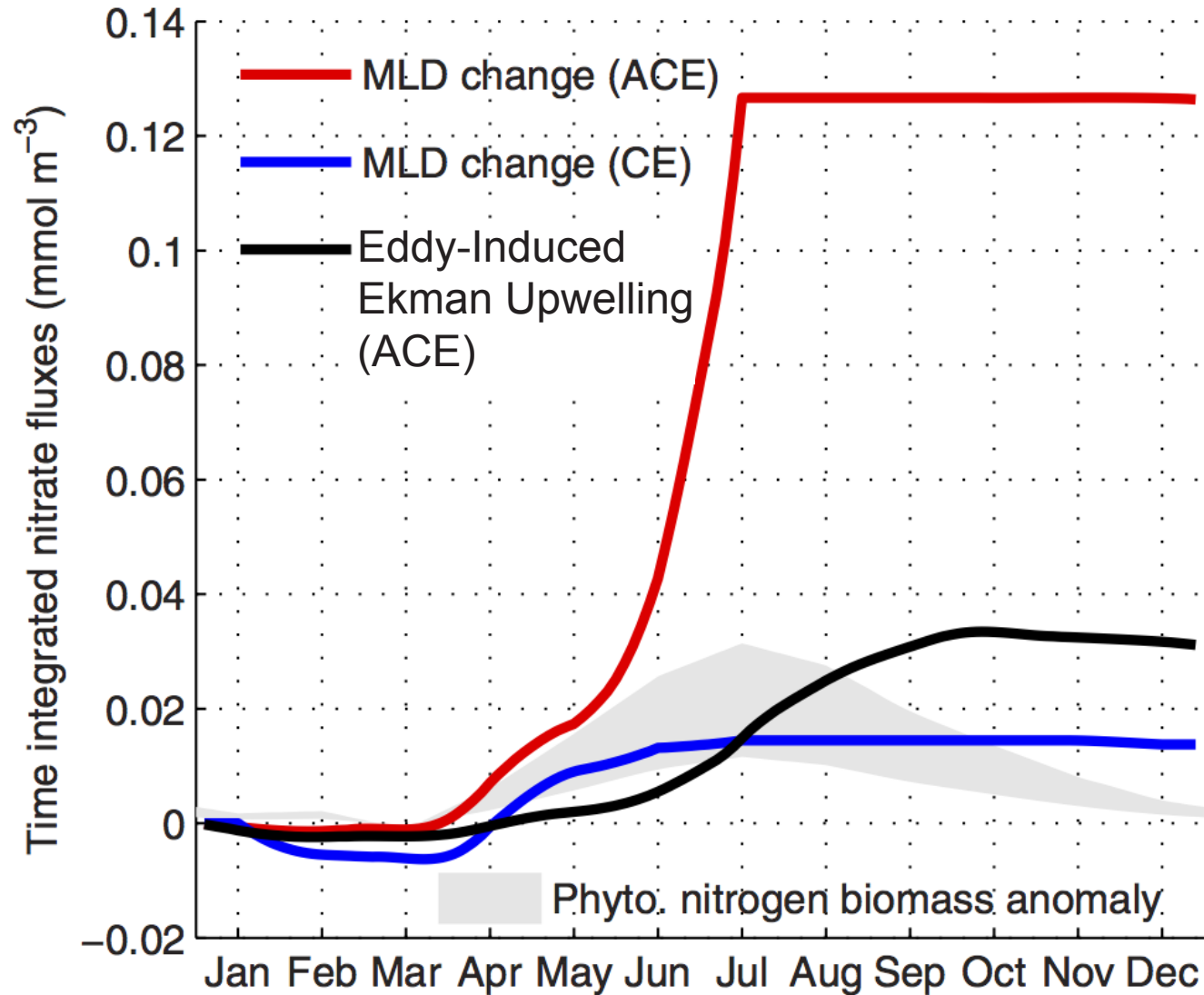
Seasonal Cycle of Mixed Layer Depth in Eddies

MLD from Argo float profiles
(Holte and Talley, 2009, <http://mixedlayer.ucsd.edu/>)



Gaube et al., (2013), also
described in Dufois et al., (2014)

Integrated Nitrate Flux Resulting from Deeper Mixing in Anticyclones

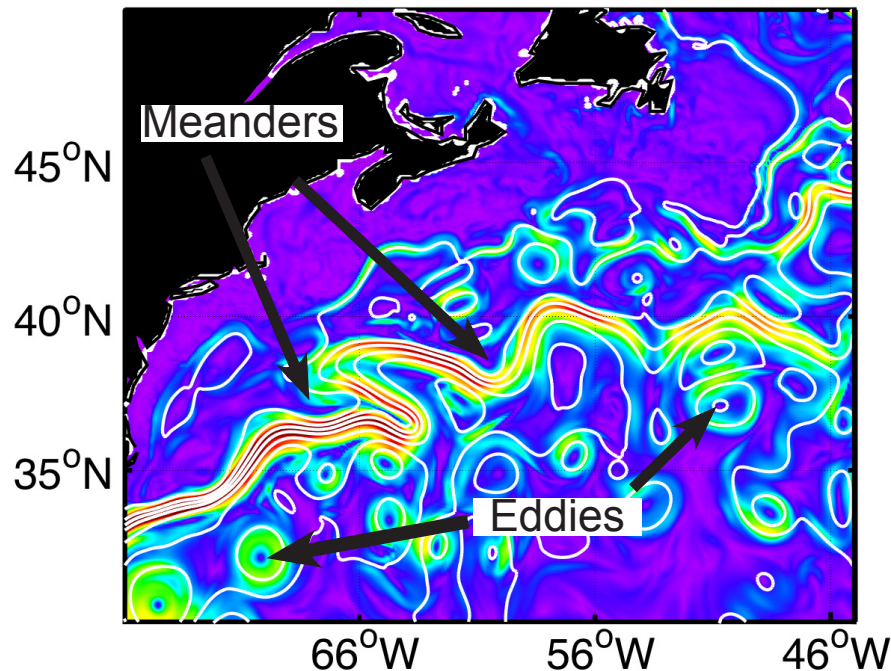


Summary and Conclusions

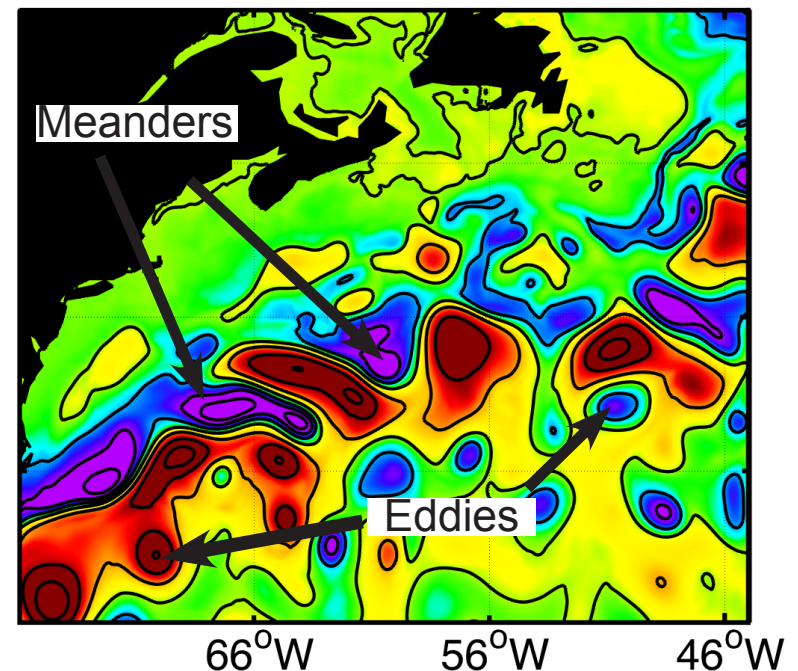
- Mechanism of mesoscale physical/biological interaction can be classified as:
 1. The advection ecosystems and nutrients horizontally
 2. Vertical fluxes of ecosystems and nutrients
 3. Eddy-induced perturbation of near-surface mixing
-
- Globally, eddy stirring dominates the observed impact of eddies on CHL
- Our analysis reveals that multiple mesoscale mechanisms influence the observed response of CHL to eddies in a given region
- By combining satellite and in situ observations, we are now able to estimate the eddy-impact on near-surface mixing

Differentiating Eddies and Meanders in the Gulf Stream Region

Current Speed and
Total Sea Surface Height



Sea Level Anomaly



- Coherent mesoscale structures are identified and tracked in fields of sea level anomaly.
- Meanders appear as closed contours in sea level anomaly
- The propagation direction of meanders and eddies are used to differentiate them.