

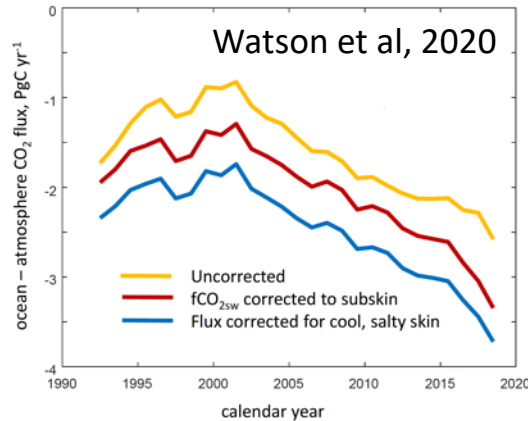
Gas Exchange Uncertainties

Working group member

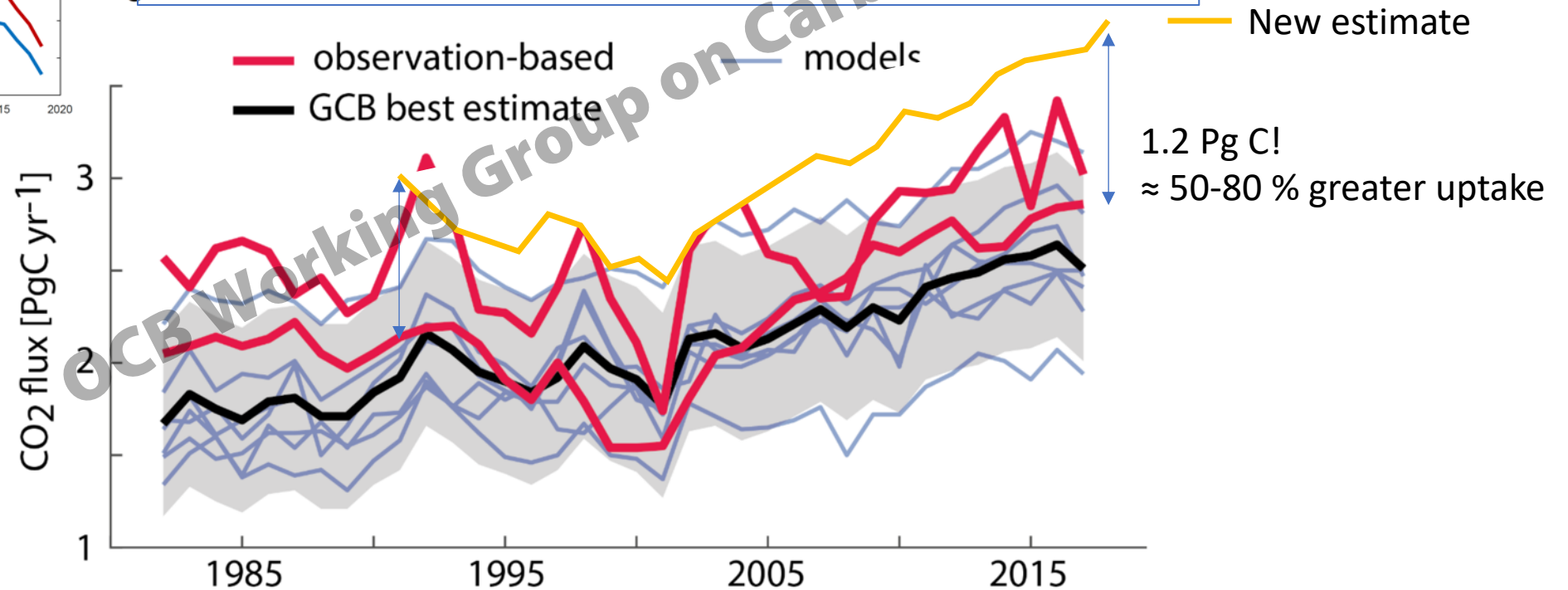
Rik Wanninkhof, NOAA/AOML, Miami



OCB Working Group: Filling the gaps in observation-based estimates of air-sea carbon fluxes



Including chemical and thermal gradients in the surface boundary (≈ 100 micron) could increase the $\Delta p\text{CO}_2$ by $\approx 6 \mu\text{atm}$ and thereby the air-sea flux by 1.2 Pg C [Watson, Woolf, Shutler et al.)



Its not who is right or wrong- but rather that there are unresolved issues.

Direct Flux measurements : $F = w'c'$

are challenging:

- Experimental
- Need large concentration gradient ($\Delta p\text{CO}_2 \approx 40 \mu\text{atm}$)
- Use to understand the forcing of transfer
- Interpolating Flux on large scales very tricky

Bulk Formulation: a physical and an operational definition

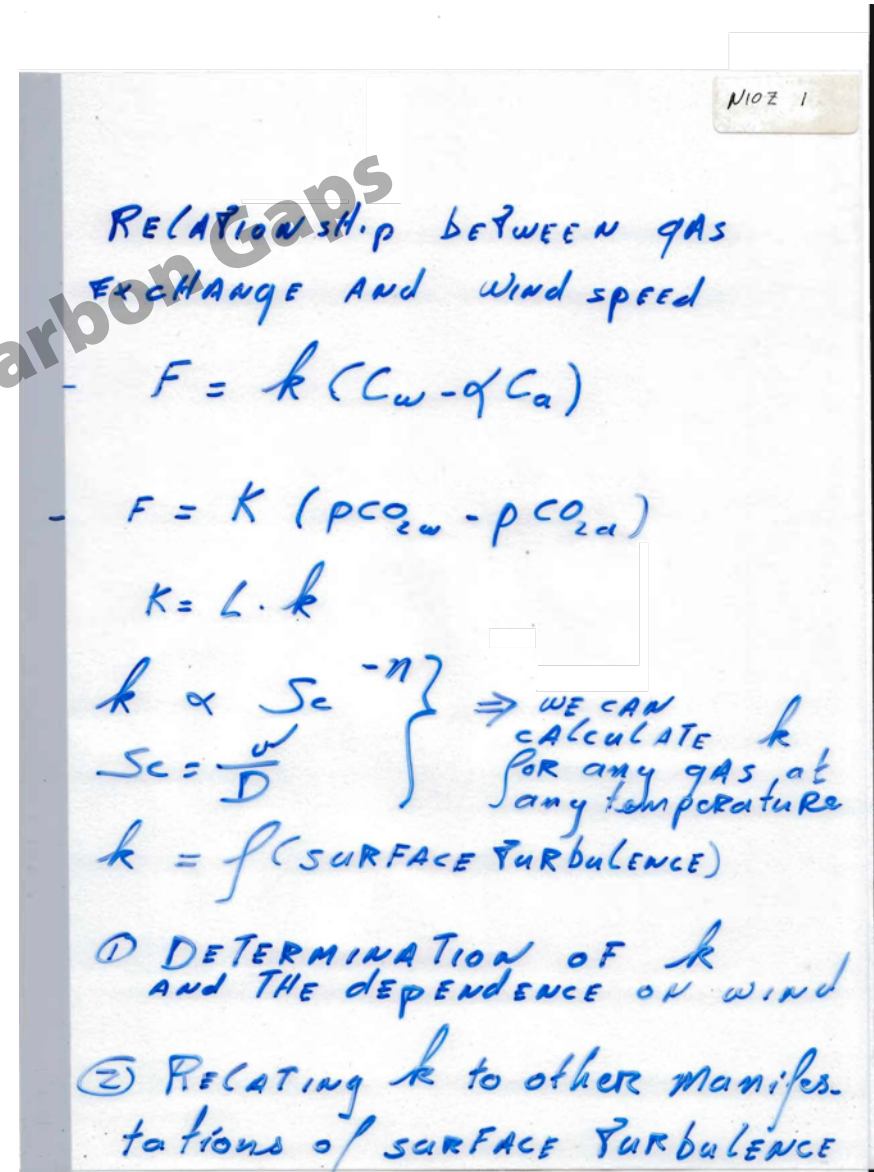
$$F = k (C_w - K_0 C_a)$$

$K_0 C_a = C_0$ right at the interface

C_w = right below the mass (diffusive) boundary layer
retarding the flux

$$F = k K_0 (p\text{CO}_{2w} - p\text{CO}_{2a})$$

Measurement in bulk phases (air $\approx 10\text{-m}$; water $\approx 1\text{-m}$)



RELATIONSHIP BETWEEN GAS EXCHANGE AND WIND SPEED

$$F = k (C_w - K_0 C_a)$$
$$F = K (p\text{CO}_{2w} - p\text{CO}_{2a})$$
$$K = L \cdot k$$

$k \propto Sc^{-n}$
 $Sc = \frac{w}{D}$

\Rightarrow WE CAN CALCULATE k FOR ANY GAS AT ANY TEMPERATURE

$$k = f(\text{SURFACE TURBULENCE})$$

- ① DETERMINATION OF k AND THE DEPENDENCE ON WIND
- ② RELATING k TO OTHER MANIFESTATIONS OF SURFACE TURBULENCE

Responsive to SOLAS themes:

- Theme 1. Greenhouse gases in the ocean (Tish, Greg Cutter)
- Theme 2. Air-sea interface and fluxes of mass and energy (Rachel/Penny, Haus, Potter)
- Theme 3. Atmospheric deposition and ocean biogeochemistry (Kate/Yuan/Bill Landing/Santiago/Greg Cutter)
- Theme 4. Interconnections between aerosols, clouds, and marine ecosystems (Cassie/Nicholas/Santiago)
- Theme 5. Ocean biogeochemical controls on atmospheric chemistry (Dave)

Workshop report chapter:

Air-sea exchange (Ho lead, Bell, Stanley, Potter pull together)

- Uncertainties and biases in CO₂ flux calculations (Palter, Cai)
- Medium (5-15 m s⁻¹) wind speed regimes (Emerson, Bell, Laxague)
- High (>15 m s⁻¹) wind speed regimes (Vlahos, Ortiz-Suslow, Stanley, Laxague, Haus, Potter)
- Coastal regimes (Ho, Najjar, Ortiz-Suslow)
- Sea spray aerosol fluxes (Henry, Allie?)
- Impact of sea ice on air-sea fluxes of gases and aerosols (Matrai, Schultz)

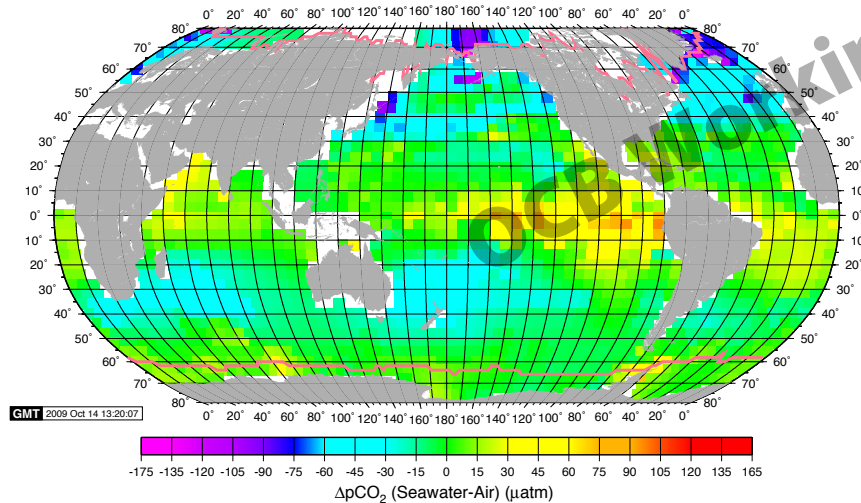
Bulk Flux equation for sea-air CO₂ exchange

$$F = k K_o \Delta p\text{CO}_2$$

Thermodynamic Forcing

$\Delta p\text{CO}_2$

$\Delta p\text{CO}_2$ (Seawater-Air) (Rev Oct 09) for June 2000



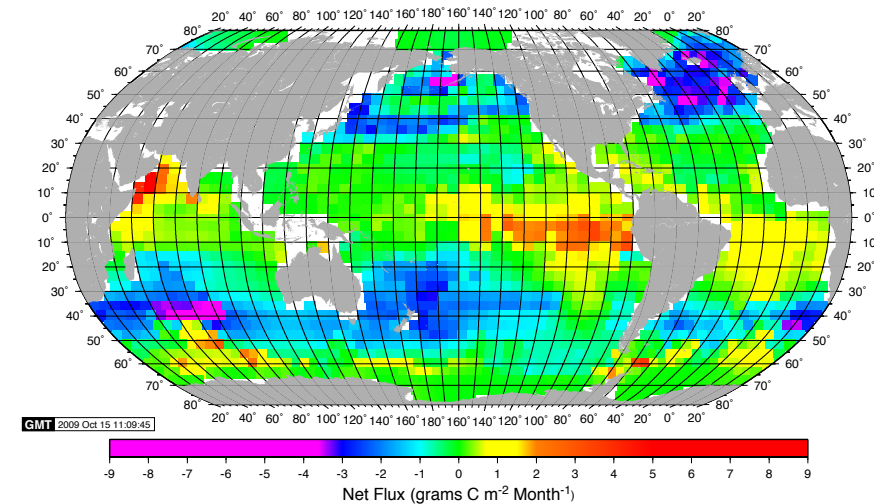
Concentration gradient

Kinetic forcing

$k = f(U_{10})$

Flux

MONTHLY Flux for June 2000 [Rev Oct 09] NCEP II Wind, 3,040K (U^2 wind, $\Gamma=.26$)



Fluxes

The importance of sea –air gas transfer in carbon cycle research

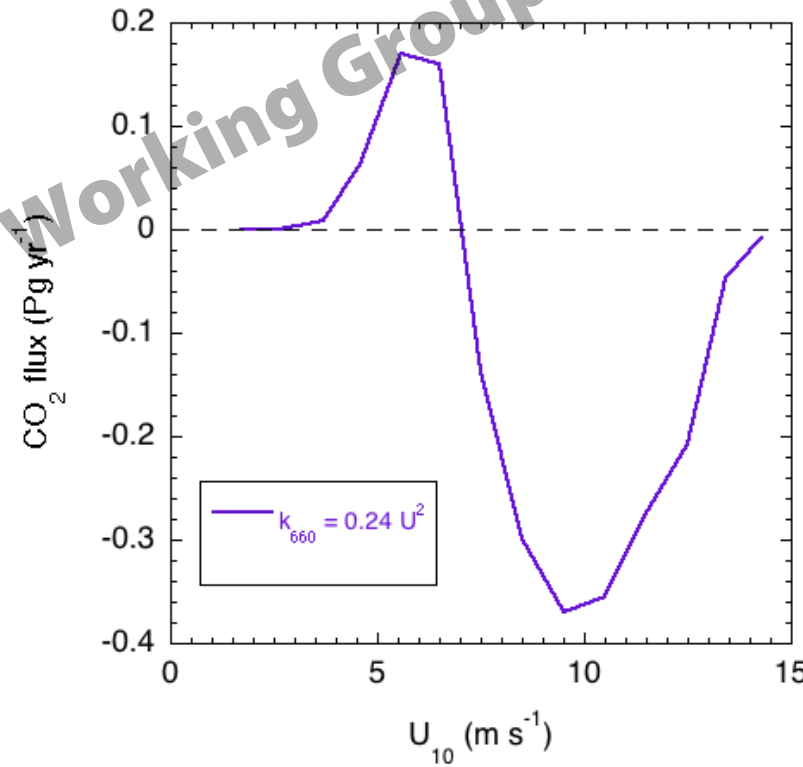
Global *net* Sea-Air CO₂ flux with wind- k parameterization

= -1.38 Pg C yr⁻¹

Global *net* Sea-Air CO₂ flux with globally average k everywhere
(global average winds)

= -0.3 Pg C yr⁻¹

There is a correlation between areas of out/in gassing and low/high winds



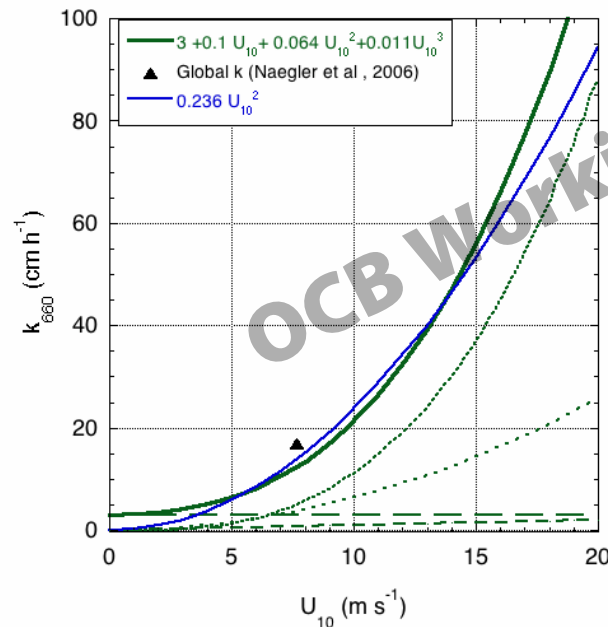
Wanninkhof et al., 2009

Reconciliation of current knowledge into a simple relationship

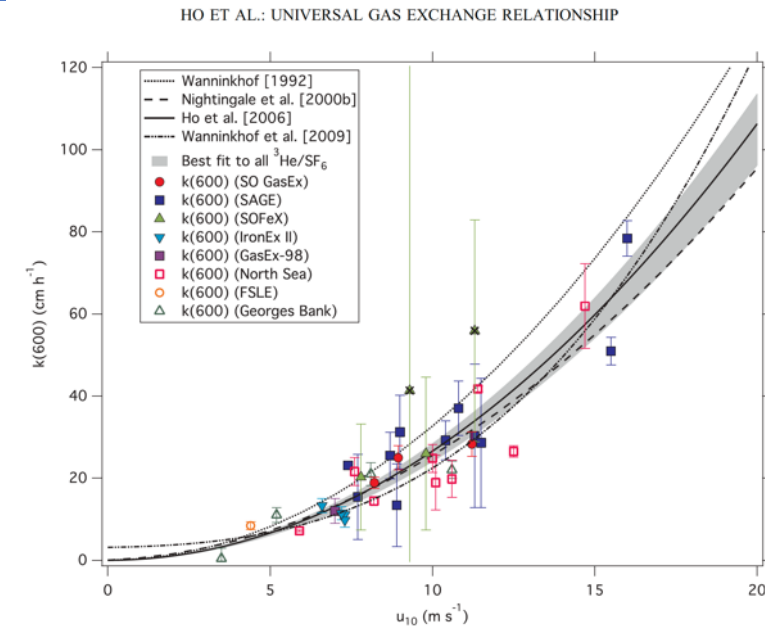
The hybrid model: What if all processes play a role?

1. Chemical enhancement and non-wind induced turbulence : Constant
2. Transfer across rigid wall: U
3. Wind stress: U^2
4. Energy dissipation/bubbles: U^3

$$k_{660} = 3 + 0.1 \langle U \rangle + 0.26 \langle U^2 \rangle + 0.011 \langle U^3 \rangle \approx k_{660} = 0.24 \langle U^2 \rangle \text{ for } 4\text{--}15 \text{ m s}^{-1} \pm 20 \%$$



(Wanninkhof et al., 2009)



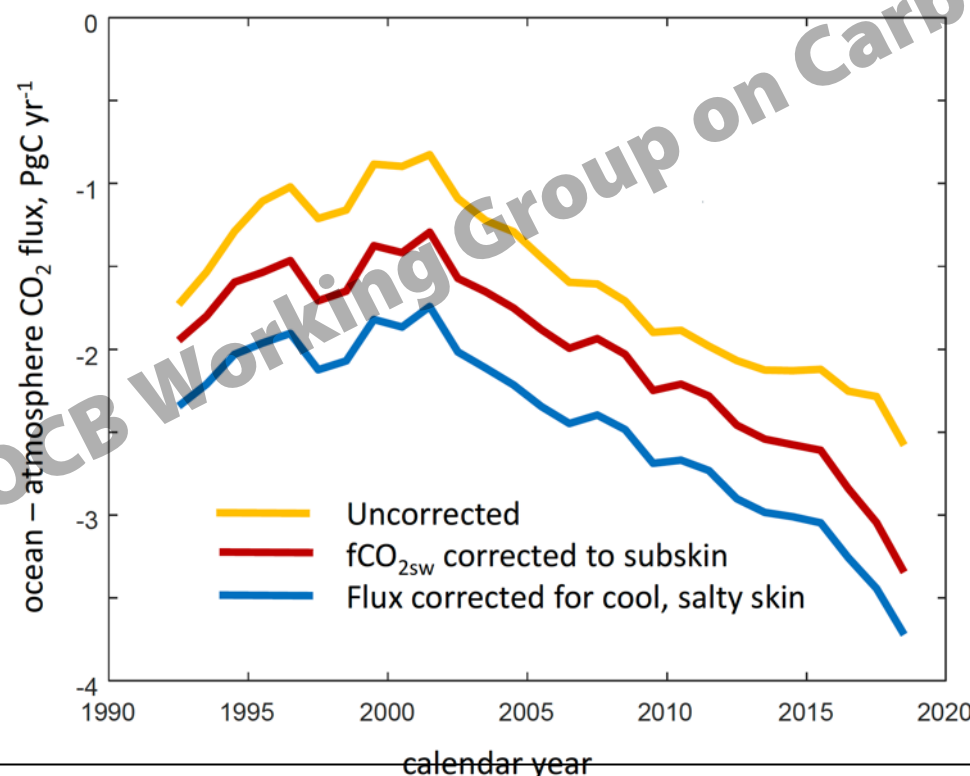
[Until we have a better knowledge of other factors and these factors can be measured globally], a quadratic dependency seems to work

Now what is the fuss?

Earth's radiation and water balance dictates a "cool and salty skin"

Having an accurate $\Delta p\text{CO}_2$ is critical (at all scales)

- $\Delta\Delta p\text{CO}_2$ of $1 \mu\text{atm} \approx 0.25 \text{ Pg C}$
- $\Delta p\text{CO}_2$ is small ≈ -50 - $50 \mu\text{atm}$ global mean $\approx <8 \mu\text{atm}$
- $p\text{CO}_2$ is very temperature dependent ($\approx 4 \% \text{ } ^\circ\text{C}^{-1}$)



Thermal $\approx 0.7 \text{ Pg C}$

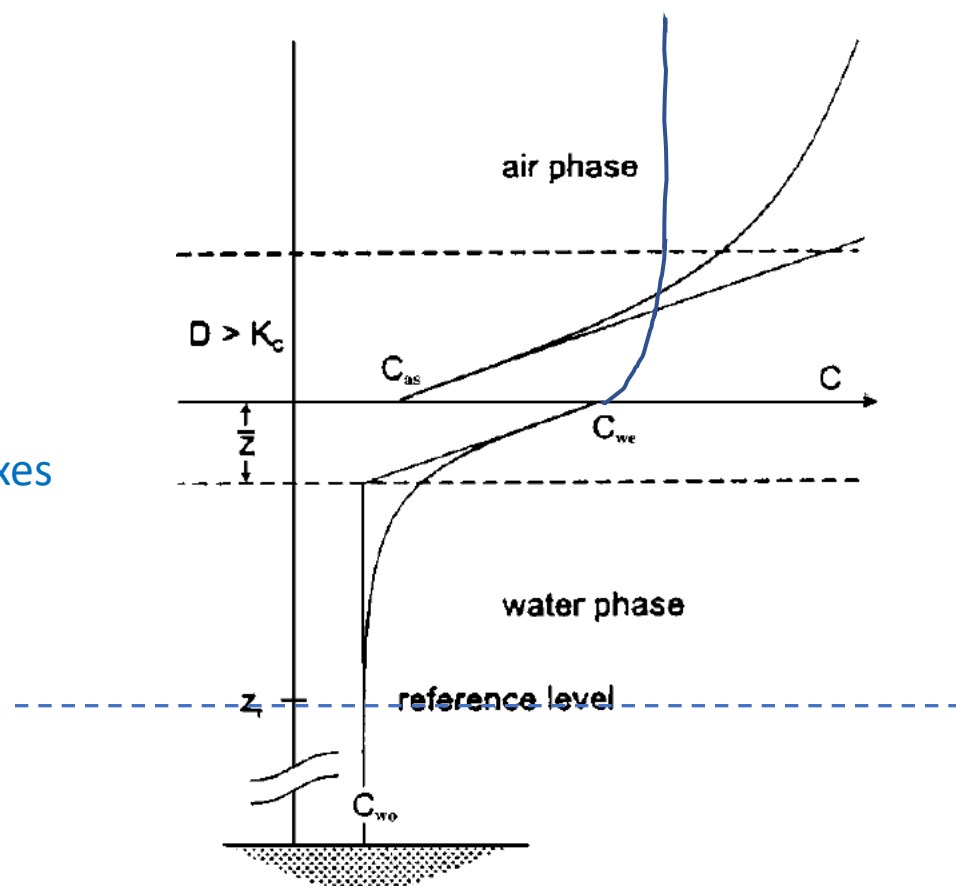
Chemical $\approx 0.4 \text{ Pg C}$

Larger estimates of ocean-atmosphere CO_2 flux are consistent with ocean carbon inventory

Andrew J Watson, Ute Schuster, Jamie. D Shutler, Thomas Holding, Ian G. C. Ashton, Peter Landschützer David K. Woolf, Lonneke Goddijn-Murphy

Mass and thermal boundary layers

Sensible and latent heat fluxes
Cause a cool salty skin
($\Delta T = 0.2\text{ }^{\circ}\text{C}$ and $\Delta S = 0.1$)



Mass boundary layer

Thermal boundary Layer

$\approx 5\text{-}10$ times the mass boundary layer

- Where is the gas transfer impeded?
- What are the conditions across the relevant boundary layer

What do we use for concentration gradient to get air-sea CO₂ flux ?

$$F = k (K_{0w} p\text{CO}_{2w} - K_{0s} p\text{CO}_{2a})$$

$$[\text{CO}_2]_{\text{aq},0} = K_0 (T_s, S_s) p\text{CO}_{2a}$$

$$[\text{CO}_2]_{\text{aq},d} = p\text{CO}_{2w} * \text{?????}$$

Questions:

- What is the appropriate concentration gradient?
- What is the [CO₂] at different locations through the mass and thermal boundary layers

Challenges:

- Very difficult to verify this issue and to address problem [on small scales]
- k and $\Delta p\text{CO}_2$ determined independently [cannot do direct flux measurement and bulk flux estimate at same time]
- Adjusting k based on operational definition using bulk phases cannot be done simply as its $\Delta p\text{CO}_2$ that drives the flux

Approaches:

- LES modeling could offer some insights
- Global mass balances could offer constraints.

Fate of anthropogenic CO₂ emissions (2007–2016)

Sources = Sinks



9.4 Pg C /yr
88%



12%
1.3 Pg C /yr

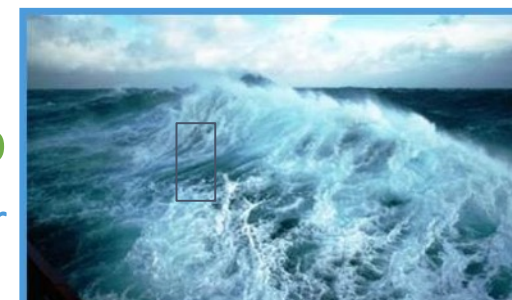
4.7 Pg C /yr
44%



28%
3.0 Pg C /yr



33 %
2.4 Pg C /yr
3.6 Pg C /yr??



Yellow:
Estimates
presented
here

How well
do we really
know the
carbon
uptake by
the ocean?

Budget Imbalance:

(the difference between estimated sources & sinks)

6%
0.6 Pg C /yr

- 6 % ?????

SOCONET, SOCAT, ICOS-OTC, ISOOS

(Wanninkhof, 8 min)



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Providing an operational foundation for sustained surface operations from a diverse community – Bottom up efforts

ISOOS/OASIS: Integrated Surface Ocean Observing system: A “Big Idea” from Ocean Obs-19 for a sustained surface ocean observing system of all physical, chemical, and biological parameters

The Decade of Ocean Science for Sustainable Development;
GOOS The Science We Need For The Ocean We Want]
Step one: Observing Air-Sea Interaction Strategy (OASIS), SCOR WG

SOCAT: Collate, “homogenize” and higher level quality control of surface $p\text{CO}_2$ data; advocate for Best practices; assess overall quality of datasets
(($\pm 2 \mu\text{atm}$ (A,B); $\pm 5 \mu\text{atm}$ (C,D); $> \pm 5 \mu\text{atm}$ (E))
Good data ingestion tools
Aspirations: other surface water measurements; water column

SOCONET: Surface CO_2 reference network – Sustain and improve surface water CO_2 measurement good enough for air-sea CO_2 flux determinations ($\pm 2 \mu\text{atm}$ water; 0.2 ppm air)
Regional network: ICOS-OTC: Integrated carbon observing system-Ocean Thematic Center A European Infrastructure project.
OTC lead Richard Sanders- looking at ways to sustain (fund) global networks

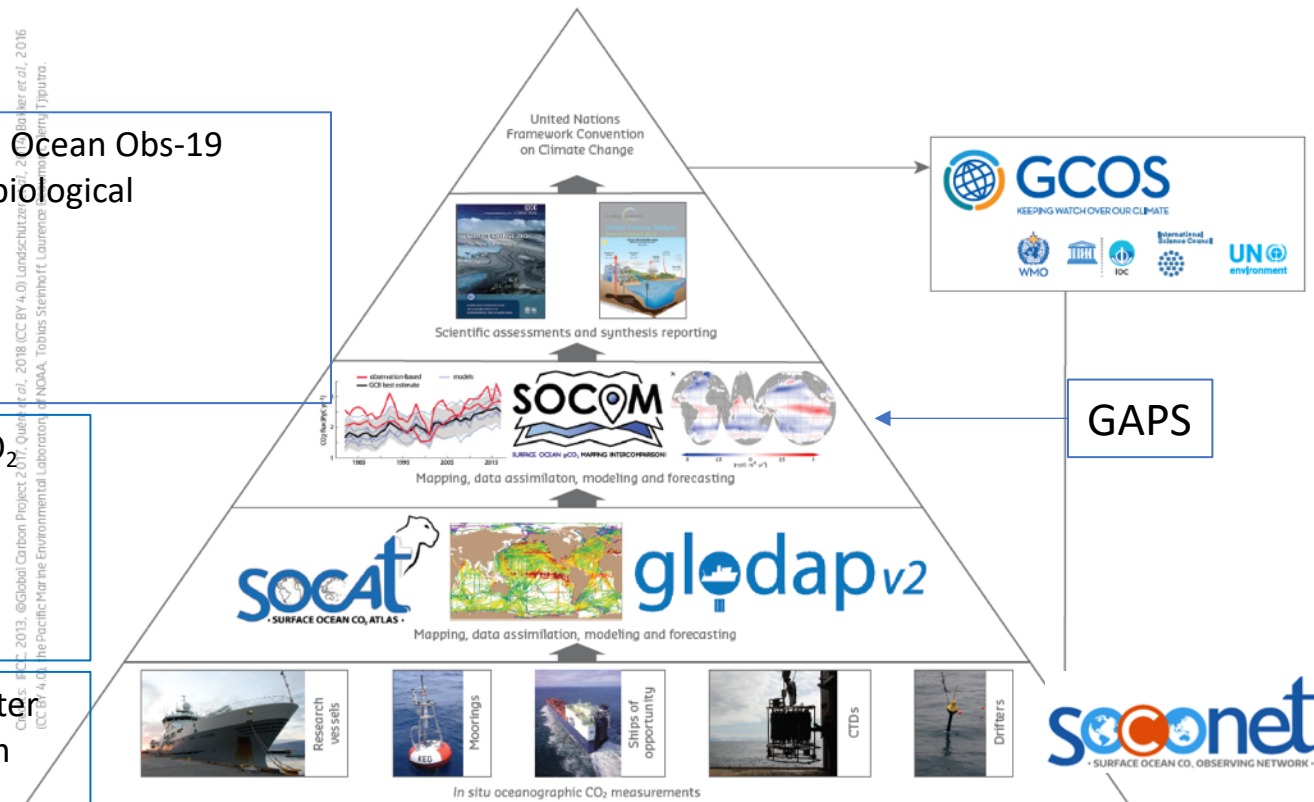


Figure 2.2. The value chain that connects *in situ* oceanographic measurements of carbonate chemistry variables to climate negotiations.